Greenhouse

Global Warming Science, EPS101

Camille Hankel and Eli Tziperman

https://courses.seas.harvard.edu/climate/eli/Courses/EPS101/

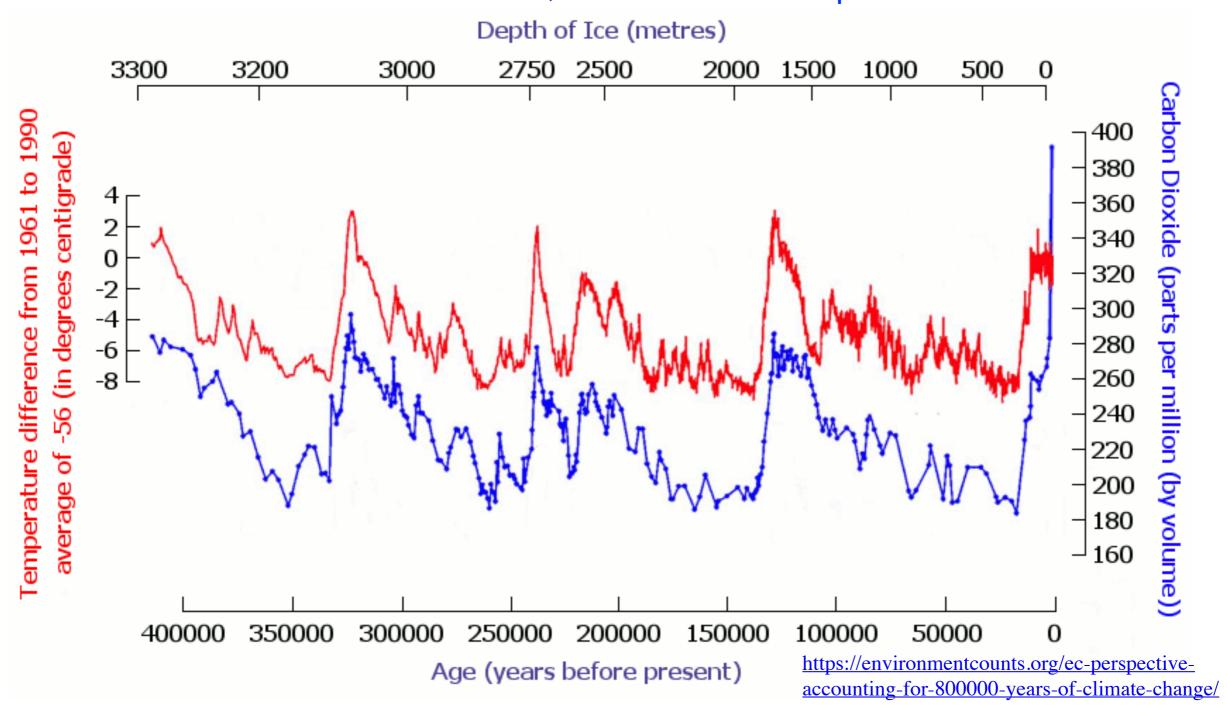
An inconvenient truth



Start at beginning for ice core collection
Start at 2:10 for intro to temperature/CO2 curve, play until 5:30 or less
https://www.youtube.com/watch?v=-JluKjaY3r4

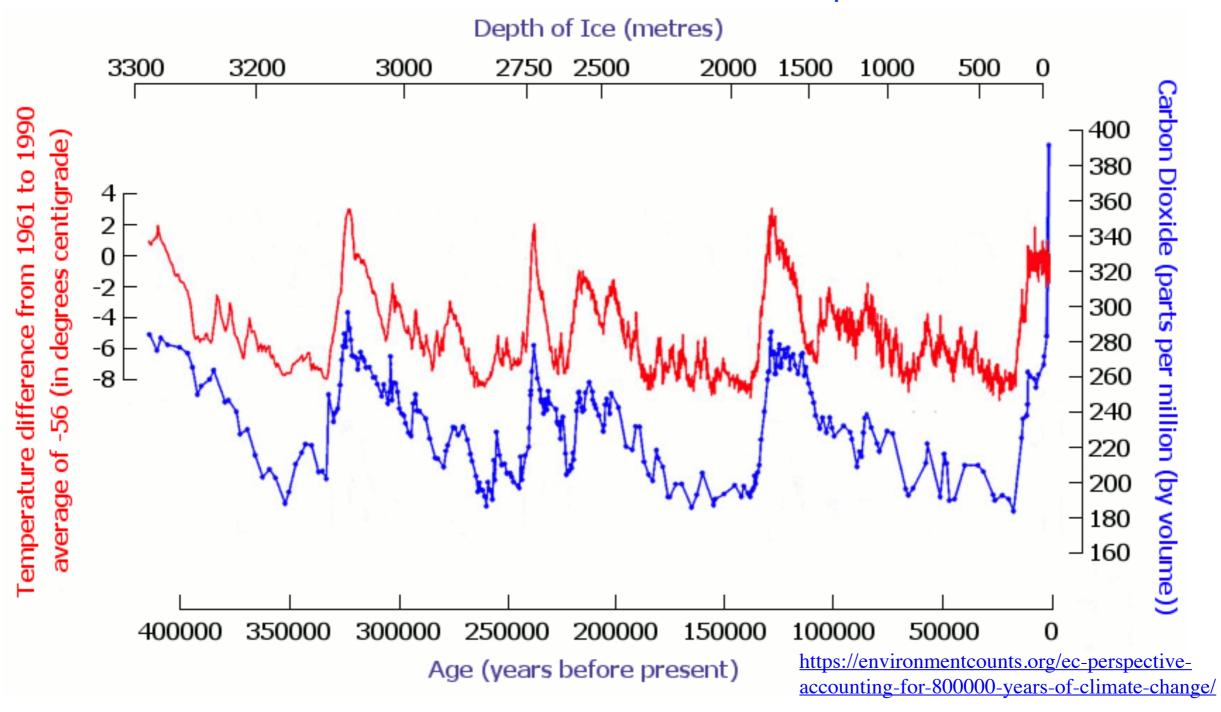
lce ages: CO₂ vs "temperature"

Vostok Antarctica, CO₂ and temperature



Ice ages: CO₂ vs "temperature"

Vostok Antarctica, CO₂ and temperature



Why are CO₂ and temperature correlated?

(IPCC AR5, 2013)

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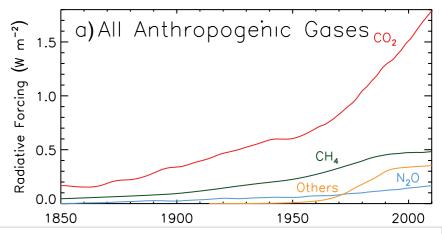
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Figure 8.6 I (a) Radiative forcing (RF) from the major well-mixed greenhouse gases (WMGHGs) and groups of halocarbons, 1850-2011, (b) as (a) but with a log scale, (c) RF from minor WMGHGs, 1850-2011 (log scale). (d) Rate of change in forcing from the major WMGHGs and groups of halocarbons, 1850-2011.



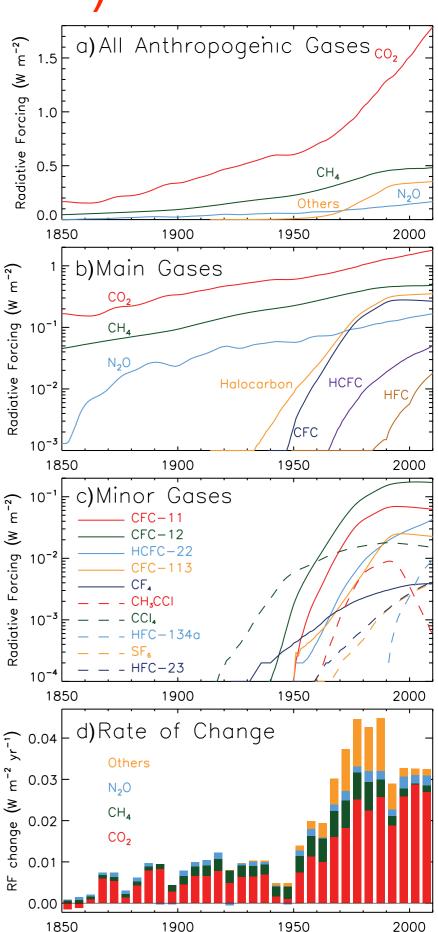
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Representative Concentration Pathway (RCP)

Future scenarios.

RCP8.5: worst-case scenario/ business as usual

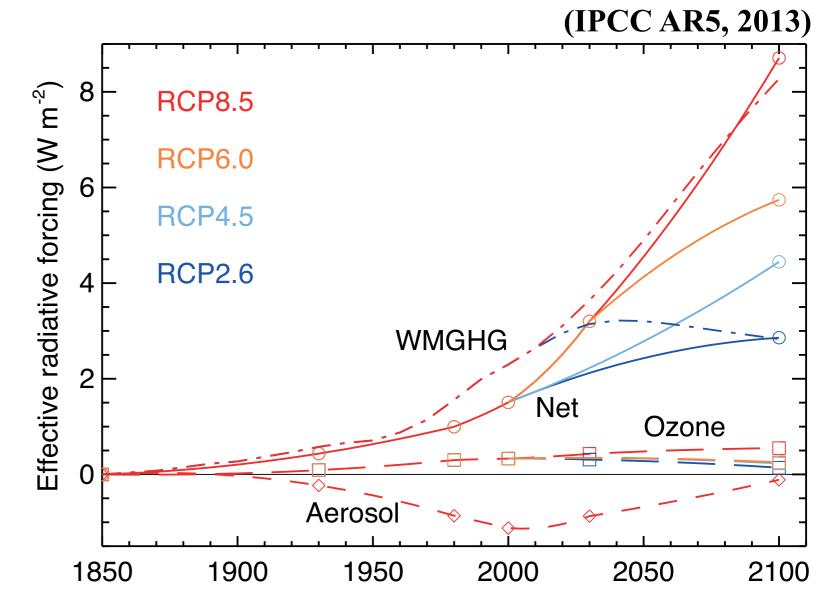


Figure 8.22 I Global mean anthropogenic forcing with symbols indicating the times at which ACCMIP simulations were performed (solid lines with circles are net; long dashes with squares are ozone; short dashes with diamonds are aerosol; dash-dot are WMGHG; colors: RCPs with **red for RCP8.5**, **orange RCP6.0**, **light blue RCP4.5**, **and dark blue RCP2.6**). RCPs 2.6, 4.5 and 6.0 net forcings at 2100 are approximate values using aerosol ERF projected for RCP8.5 (modified from Shindell et al., 2013c). Some individual components are omitted for some RCPs for visual clarity.

• How do greenhouse gases warm the planet?

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- How can such small changes in CO₂ concentration (ppm) make such a big difference?

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- What is the role of water vapor in the greenhouse effect?

Workshop #1:

Observed and projected increase in greenhouse gasses

Observations & projections of CO₂ concentration

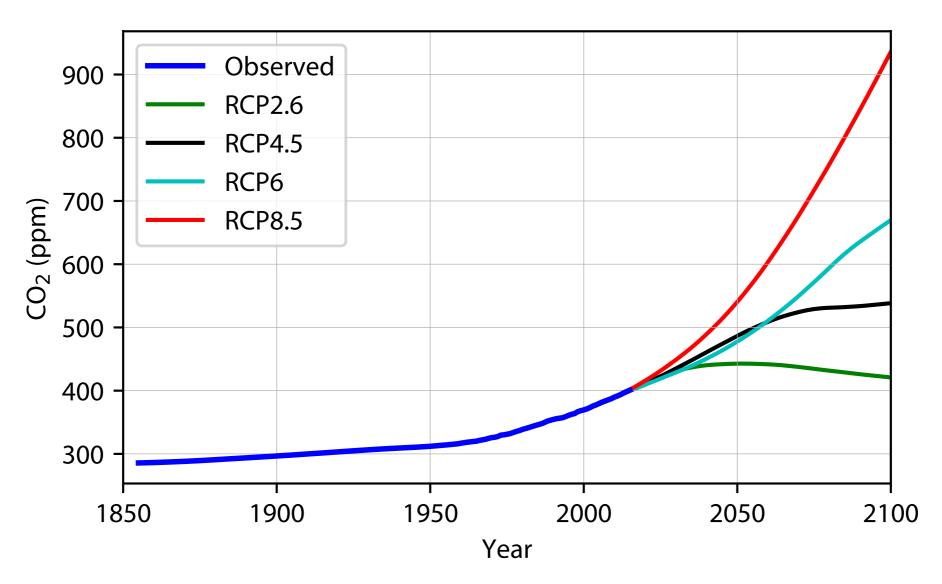


Figure 2.1:CO₂ time series.

Annually averaged CO2 concentration, observed and projected according to different RCP scenarios.

Historical & projected concentrations

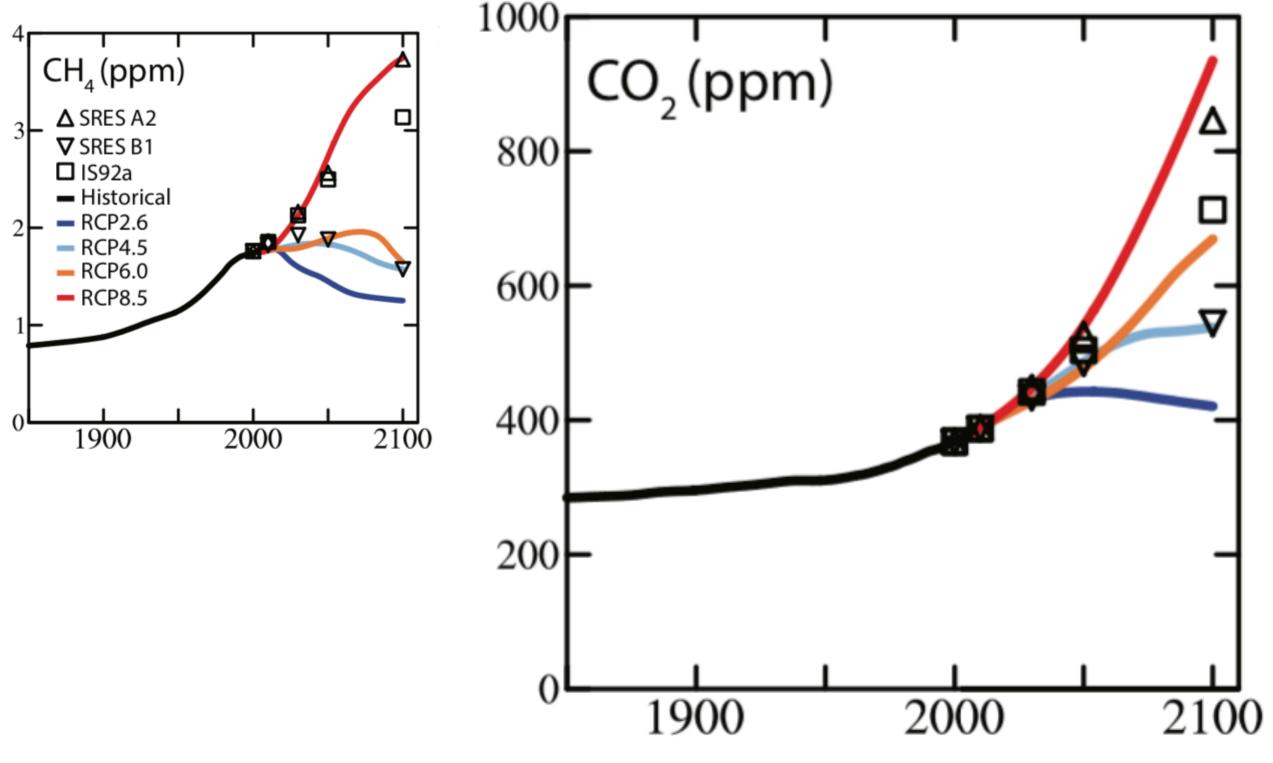


Figure 8.5 (above) Time evolution of global-averaged mixing ratio of long-lived species 1850–2100 following each RCP; blue (RCP2.6), light blue (RCP4.5), orange (RCP6.0) and red (RCP8.5). (Meinshausen et al., 2011b)

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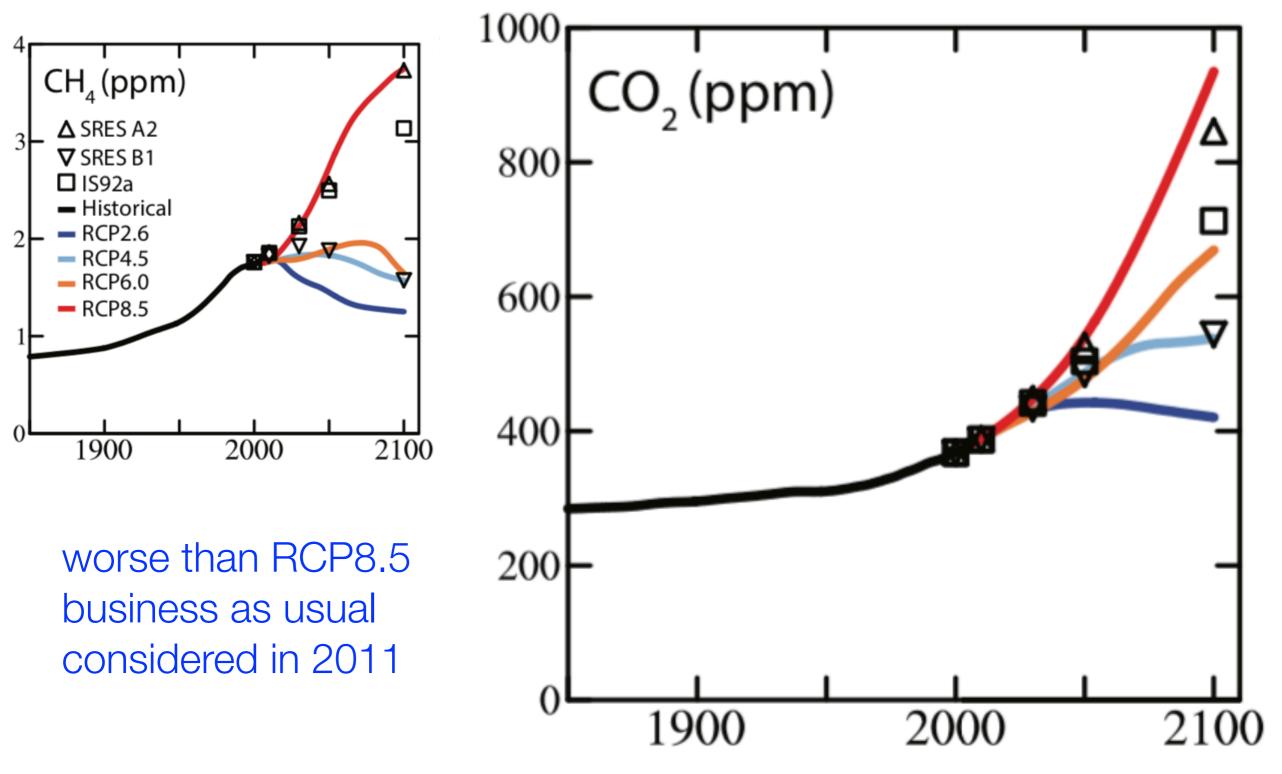


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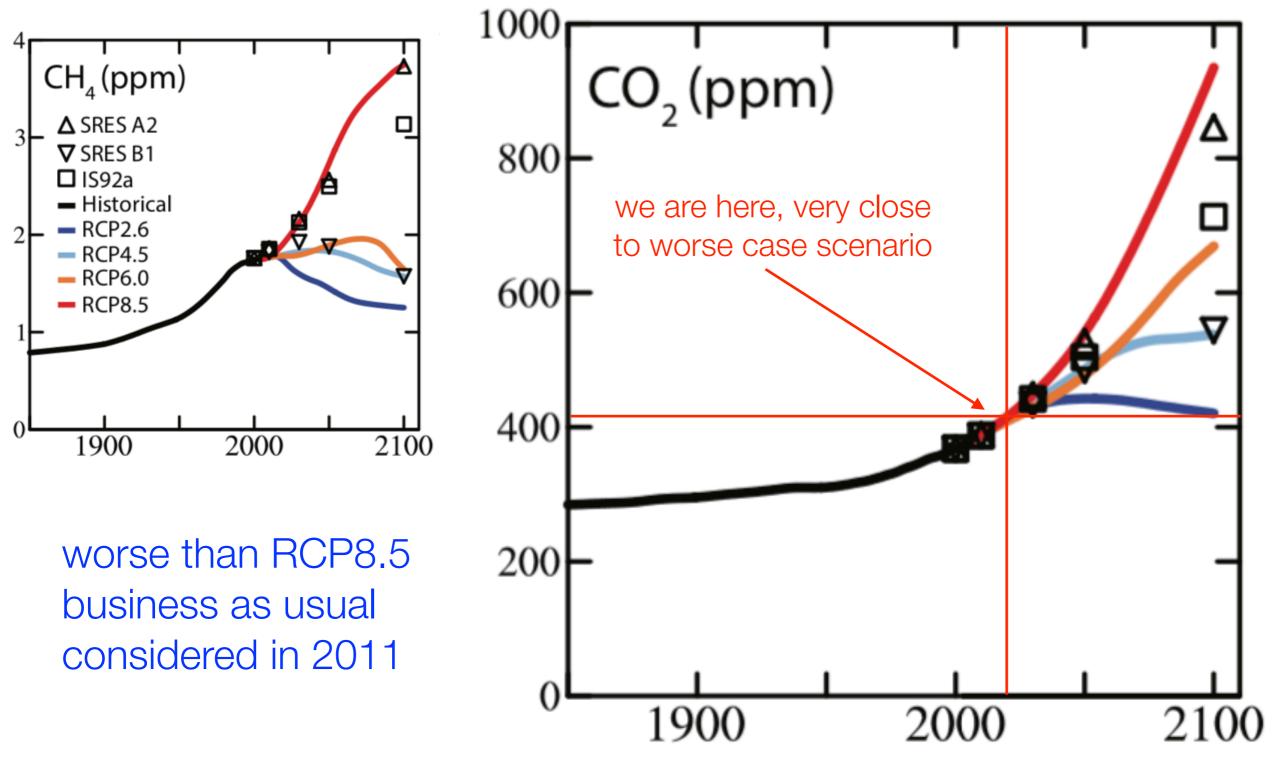


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Energy balance: Albedo, greenhouse

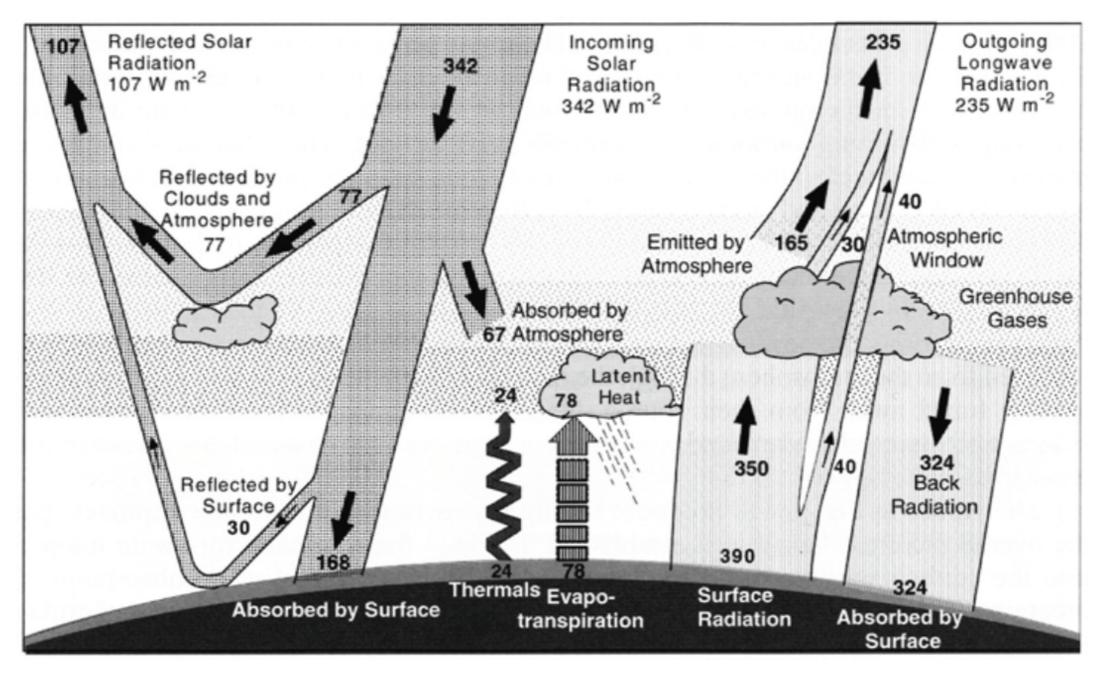


FIGURE 9.4. Earth's energy balance (from Trenberth, K.E., and D.P. Stepaniak, 2004: The flow of energy through the Earth's climate system. **Q.J.R.Meteorol.Soc.**, **130**, 2677-2701).

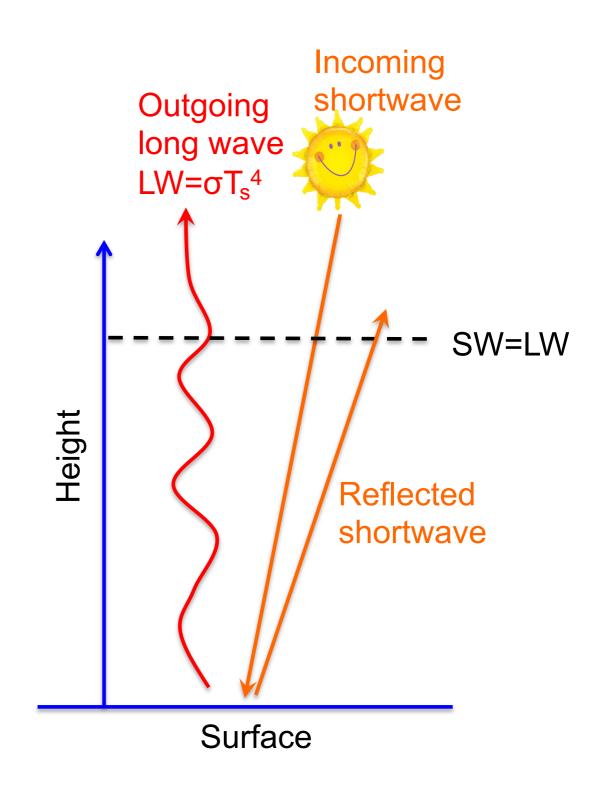
Notes sections 2.1.1, 2.1.2, 2.1.3: Energy balance, 2-layer model, continuous temperature profile and level of last absorption

(use next three slides)

Energy balance of the Earth

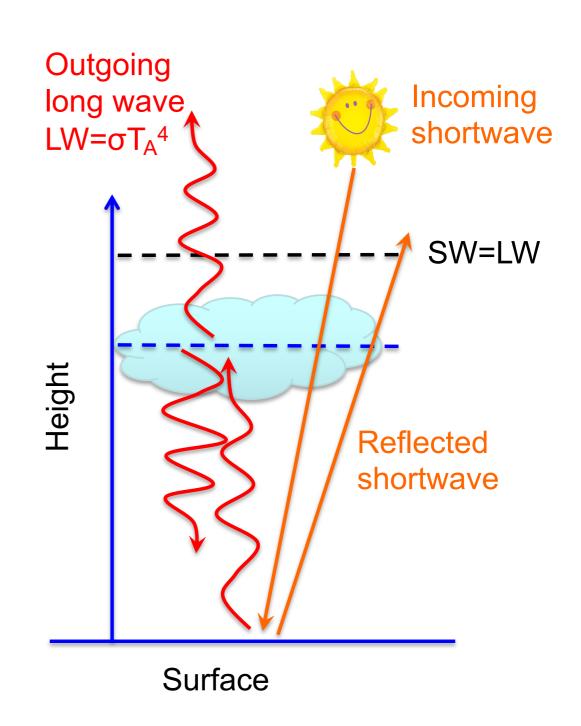
Step 1: no atmosphere!!

- Energy conservation: incoming SW radiation to Earth = outgoing LW radiation to space
- Incoming $\mathbf{SW} = \frac{S_o}{4}(1 \alpha)$
- α = albedo = proportion SW reflected
- Outgoing **LW** = σT^4
- Set incoming = outgoing -> solve for T

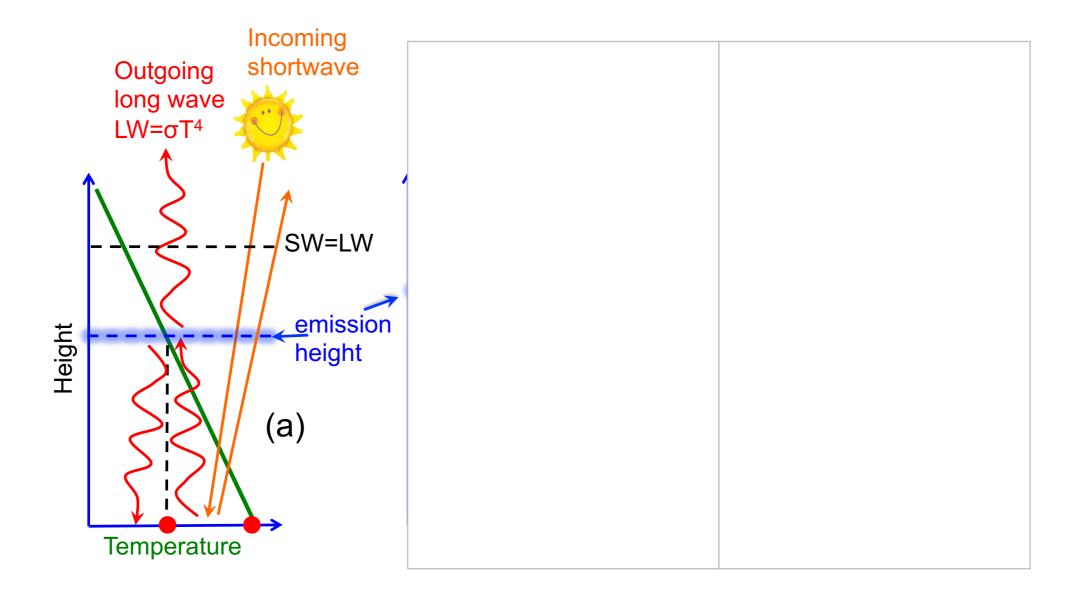


Step 2: add a 1-layer atmosphere

- Add an atmospheric layer (because gases in the real atmosphere absorb radiation): transparent to SW, absorbs/emits LW
- LW radiation emitted from surface is "trapped" (absorbed and re-emitted) by atmosphere
- Two unknowns: surface temperature T and (mid) atmospheric temperature θ . Two equations (energy balance at surface, and at mid-atmosphere)
- Do the calculations (see notes) and result: surface temperature increases!
- This is the "greenhouse effect"

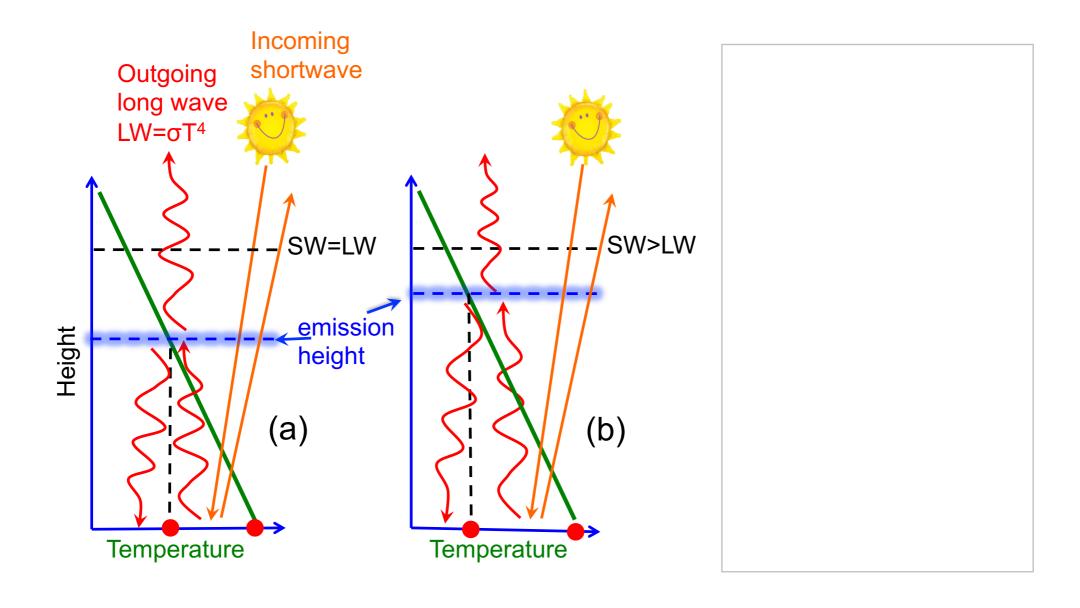


Step 3: add a continuous atmospheric temperature profile



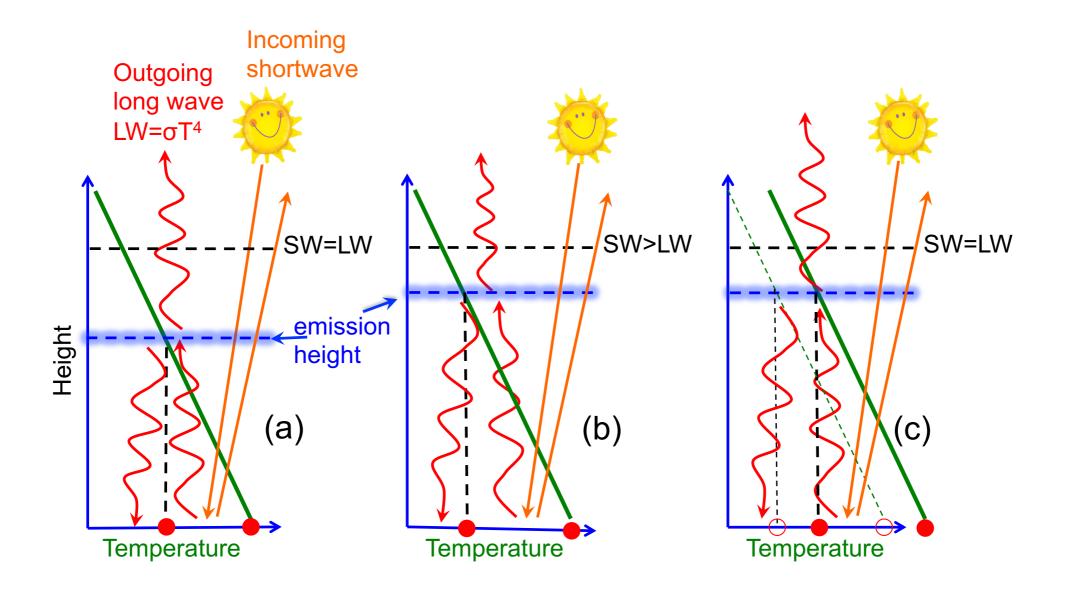
- Level of last absorption: where most of the radiation emitted upward escapes to space, without getting absorbed again
- Increasing greenhouse gas —> raising level of last absorption —> Earth radiates from a colder temperature —> Energy balance is broken: LW < SW —> temperature must adjust

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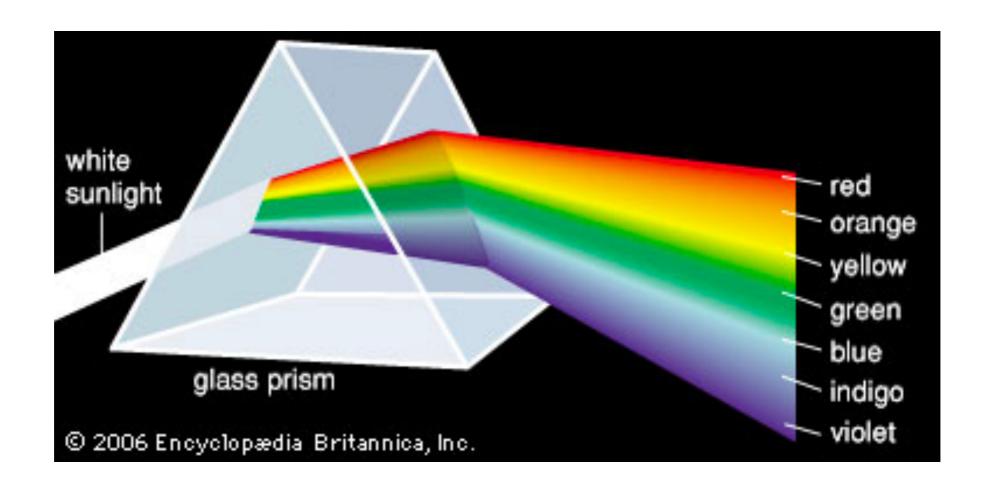


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Workshop #2: 2-layer energy balance model

How greenhouse gasses work: wavelength-dependent radiation & molecular energy levels

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Black body radiation

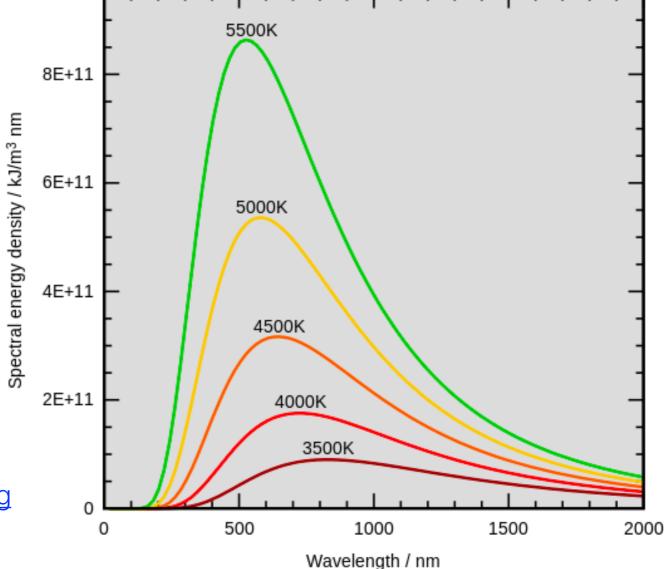
Planck's law of black-body radiation:
$$B(\lambda,T)=\frac{2hc^2}{\lambda^5}\frac{1}{e^{\frac{hc}{\lambda k_BT}}-1}$$

 $B(\lambda,T)d\lambda$ is the energy per area/ time/ angle emitted between wavelengths λ &

 $\lambda + d\lambda$; T=temperature; h=Planck's const; c=speed of light; k=Boltzmann's const.

Total emitted radiation per area/ time: σT4 Stefan-Boltzmann constant:

 $\sigma = 5.670367 \times 10^{-8} \text{ W} \cdot \text{m}^{-2} \cdot \text{K}^{-4}$



https://en.wikipedia.org/wiki/File:Wiens law.svg

Shortwave vs longwave radiation

Earth's surface and the sun both emit blackbody radiation according to Planck's function — they radiate over the full spectrum

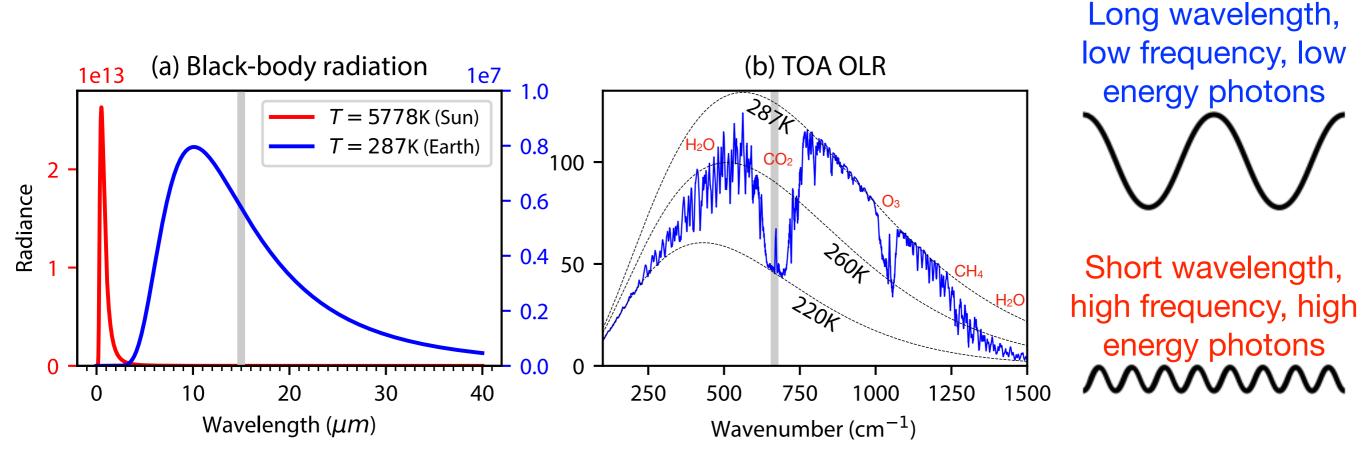
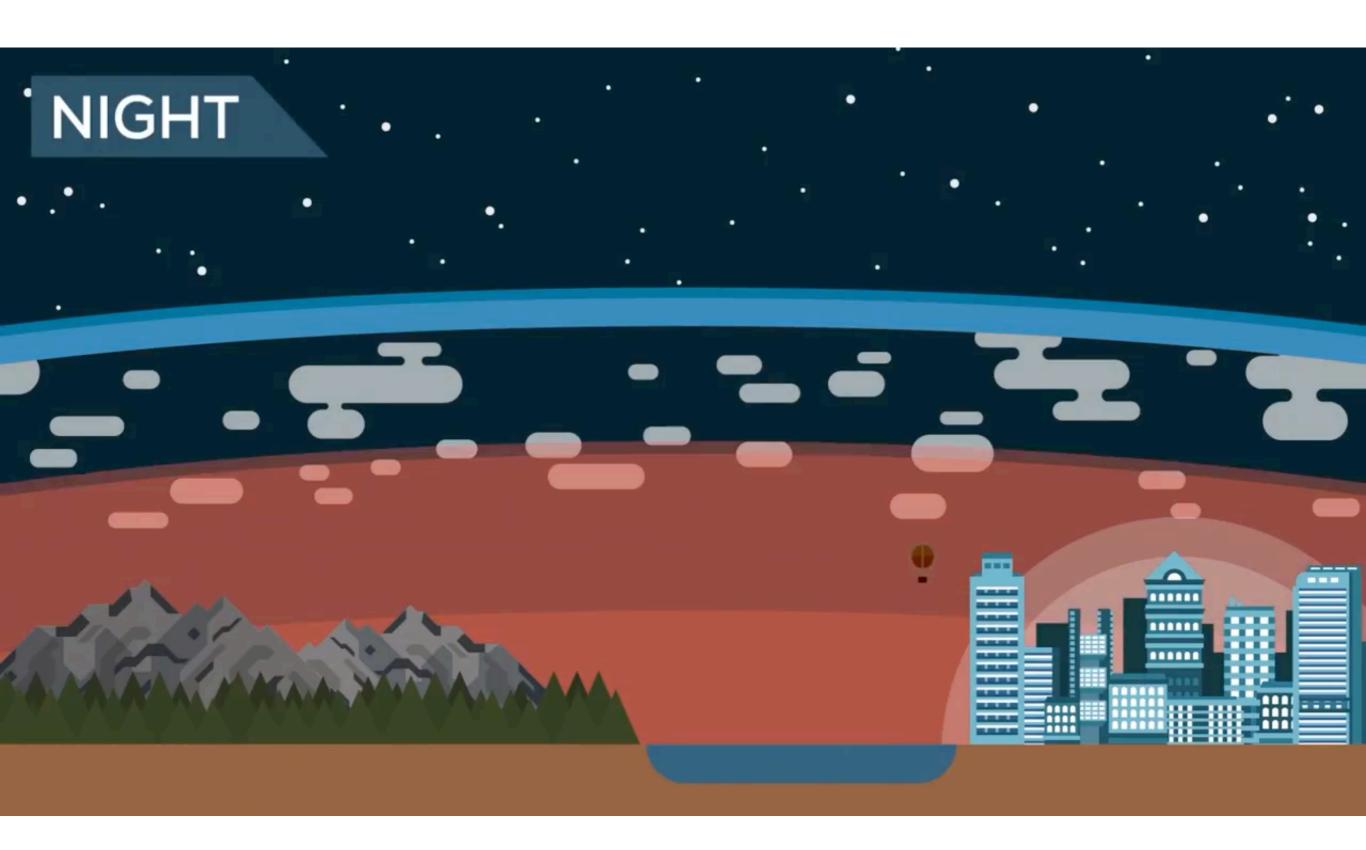


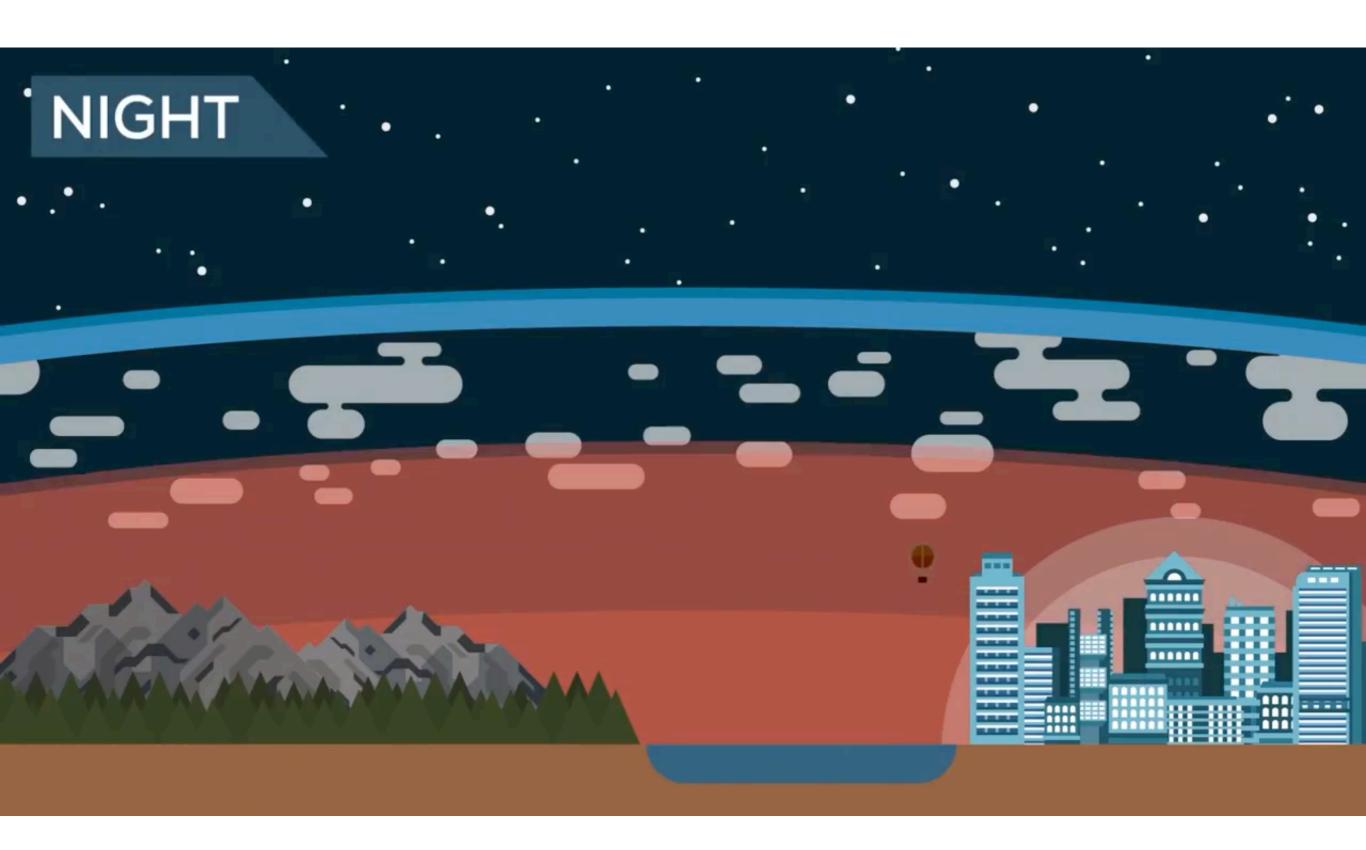
Figure 2.5: Blackbody radiation.

(a) Planck's black-body spectral radiance for the emission temperatures of the Earth (blue) and Sun (red), as a function of wavelength. (b) blue curve shows estimated outgoing longwave radiation at the top of the atmosphere as a function of wavenumber, with black-body radiation curves at different temperatures shown by dashed lines. The deviations from the 287 K black body radiation curve indicate absorption bands due to CO2, CH4, H2O, and O3; the central CO2 absorption line is shown on both panels as a vertical gray bar.

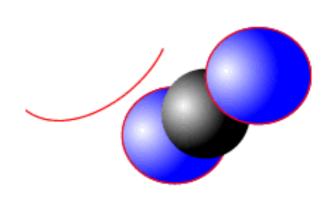
How do greenhouse gases work?

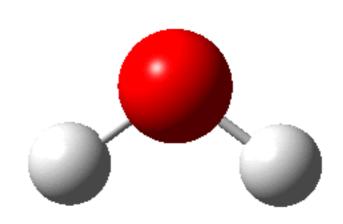


How do greenhouse gases work?



Vibration energy levels of CO₂ & H₂O





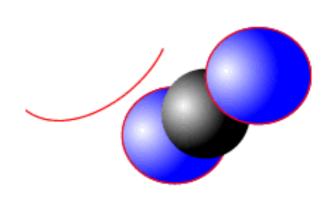
 CO_2

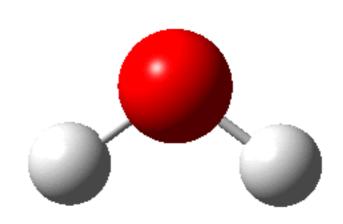
 H_2O

http://www.dynamicscience.com.au/tester/solutions1/ chemistry/greenhouse/co2andtheghe.htm http://davidobru.blogspot.com/2017/01/some-animations.html

The vibration energy levels determine the frequency of absorption

Vibration energy levels of CO₂ & H₂O





 CO_2

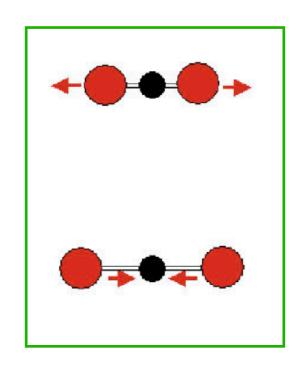
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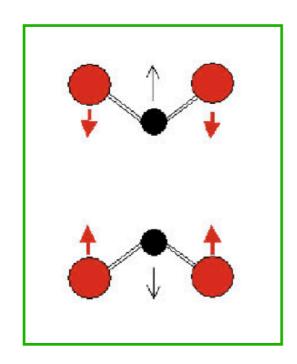
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Greenhouse gasses vibration modes

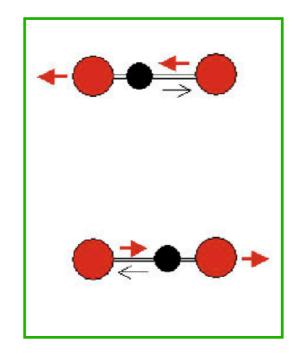
- Discrete wavelengths excite molecular transitions in greenhouse gas molecules
- This results in absorption and re-emission of radiation of that wavelength
- The "transition wavelengths" for various greenhouse gases have been measured in lab experiments



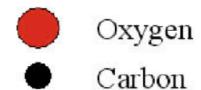
Symmetric Stretch 1366 cm⁻¹



Bending Mode 667 cm⁻¹



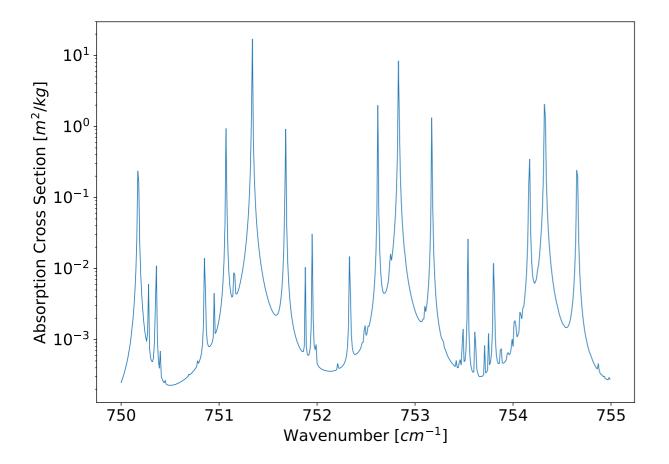
Asymmetric Stretch 2349 cm⁻¹



http://butane.chem.uiuc.edu/pshapley/GenChem1/L15/2.html

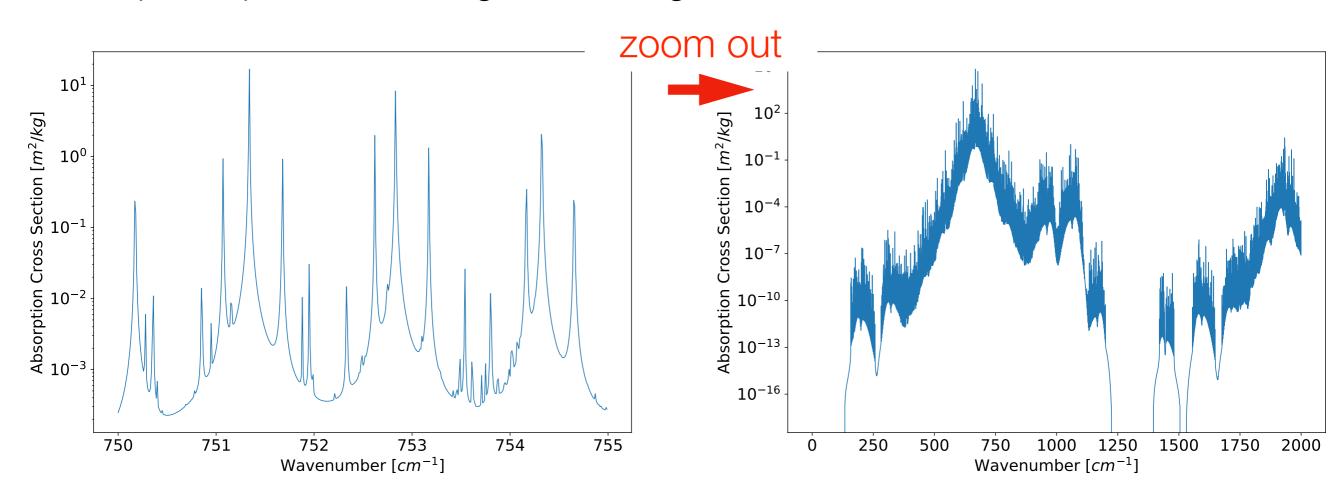
Absorption cross-section

- Spike in absorption "efficiency" i.e. cross-section at wavenumber that excites a molecular transition: spikes called <u>absorption lines</u>
- Absorption "lines" look like Gaussian/Lorentzian $(\alpha_L/\pi)/((\nu-\nu_0)^2+\alpha_L^2)$ profiles due to "line broadening"
 - Pressure broadening due to molecular collisions
 - Doppler broadening due to gaussian distribution of particle velocities
- All individual lines calculated and broadened in radiative transfer code to give "absorption spectrum" for a greenhouse gas



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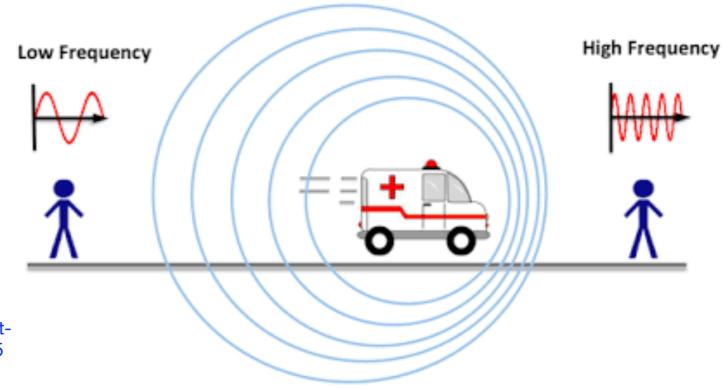


Notes section 2.2.3: Broadening

(use next three slides)

Doppler Broadening

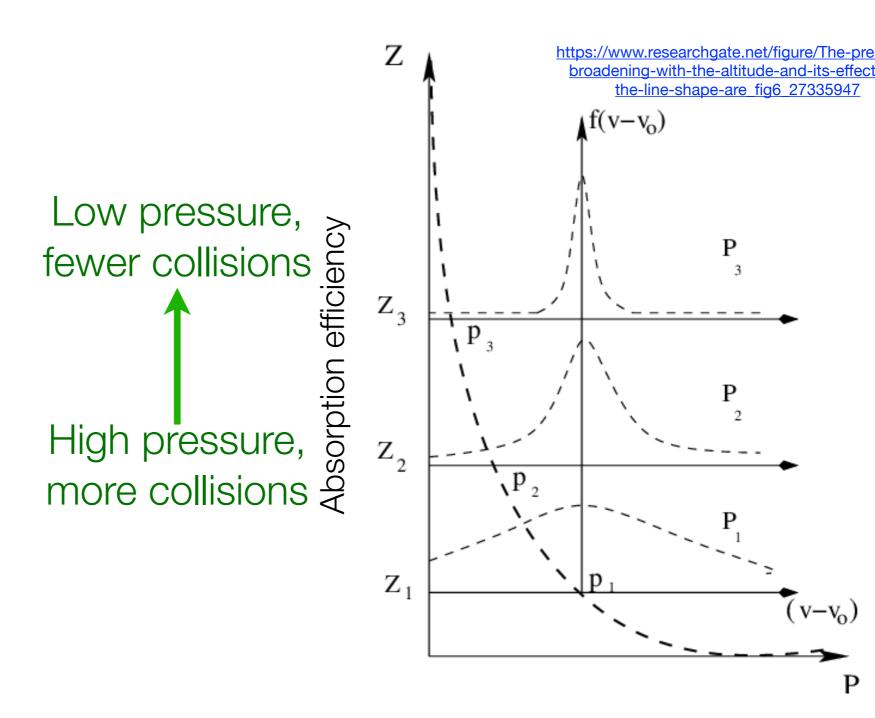
Doppler Effect



https://forum.huawei.com/enterprise/en/what-is-the-doppler-effect/thread/510221-100305

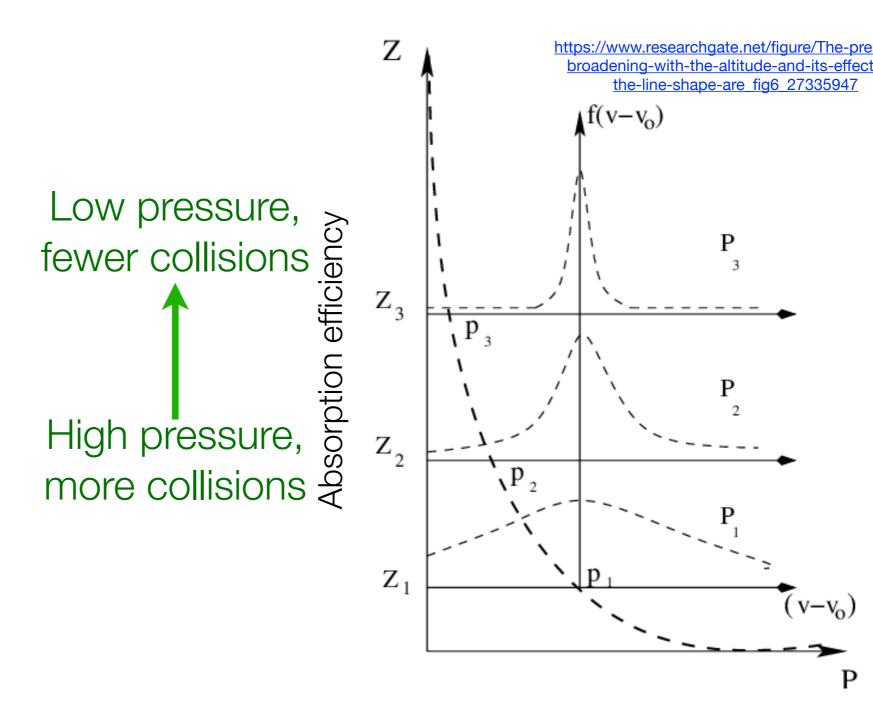
CO₂ molecule moves toward photon → photon seems at higher frequency → molecule absorbs a photon of lower frequency than that of absorption line. CO₂ molecule moves away from photon → opposite

The random motion of the gas molecules causes the widening of absorption lines because molecules that happen to be moving towards/away from the incoming photon will see it at a different frequency/ wavelengths.

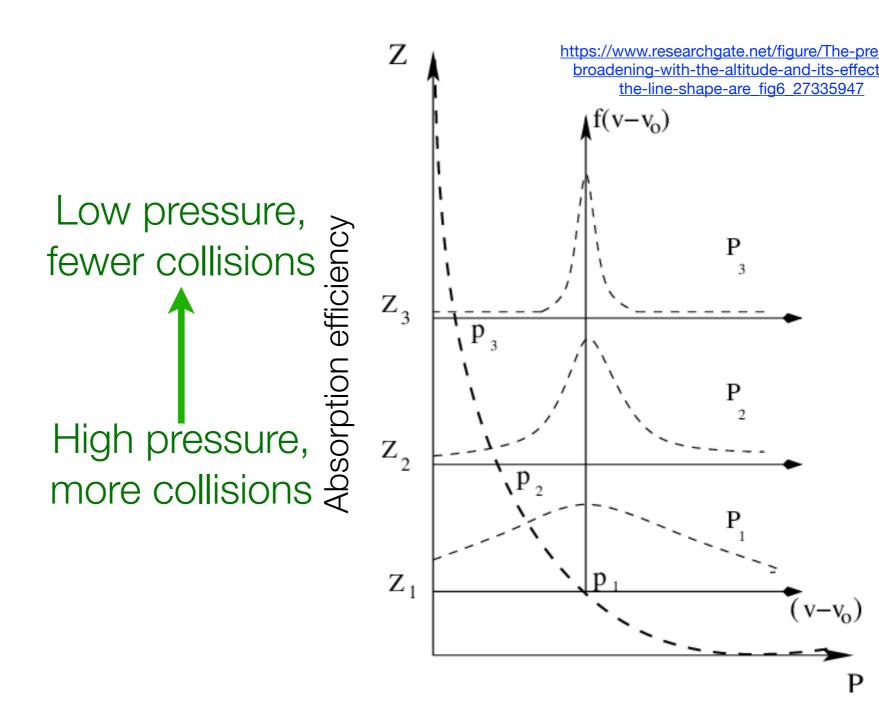


The collisions of gas molecules cause the widening of absorption lines:

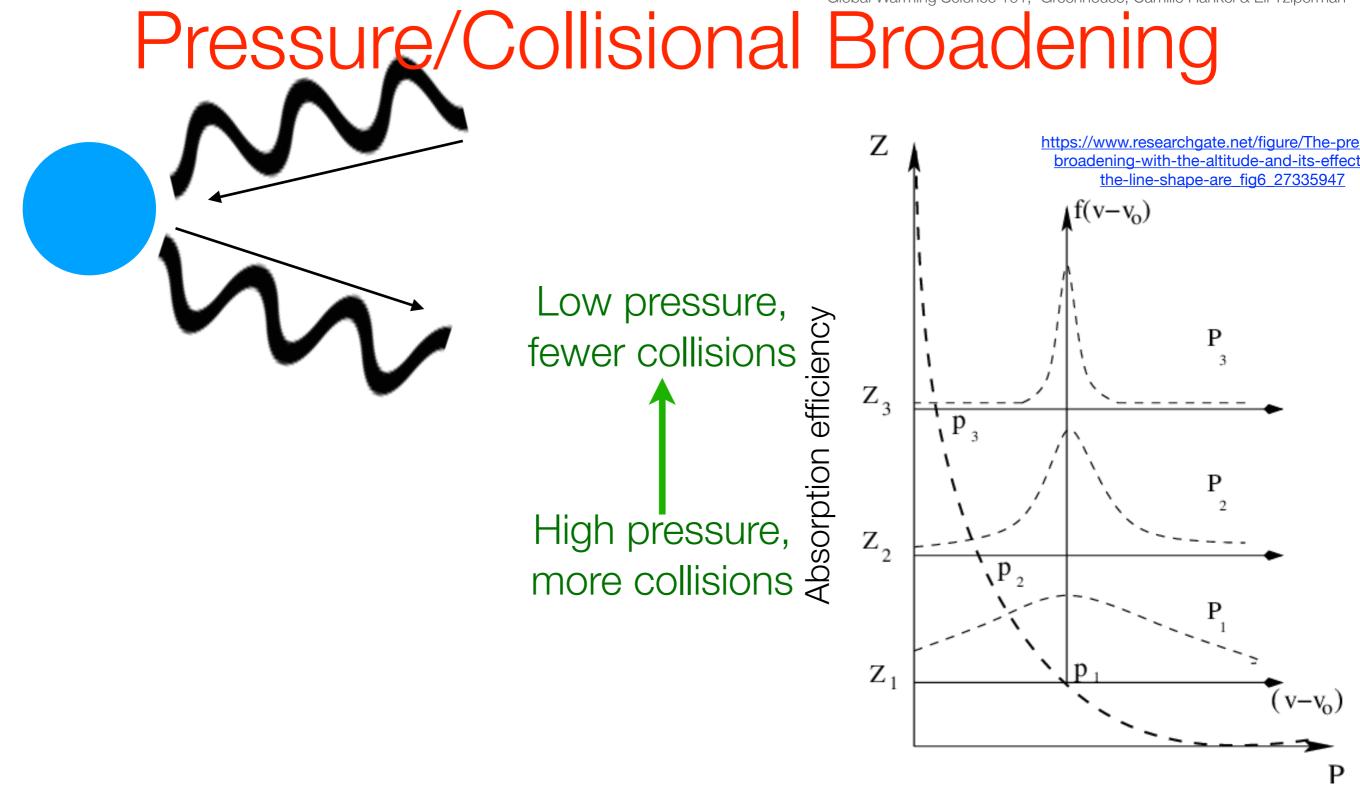




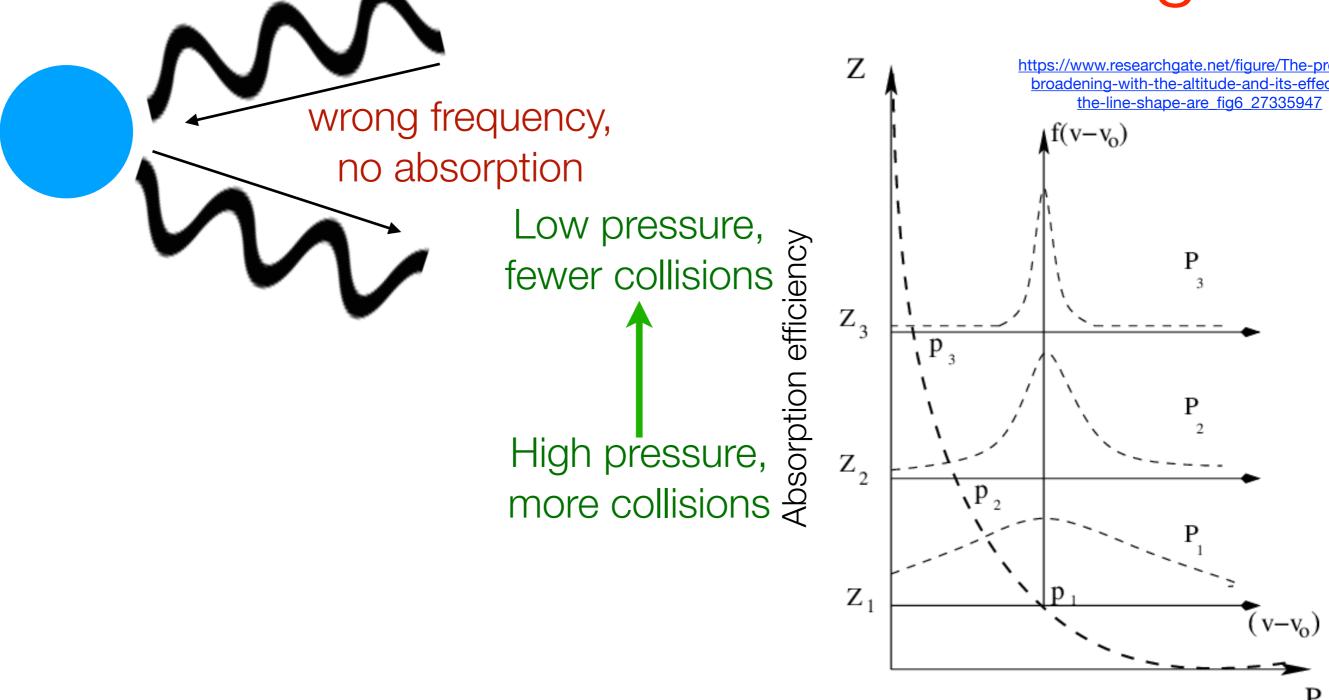
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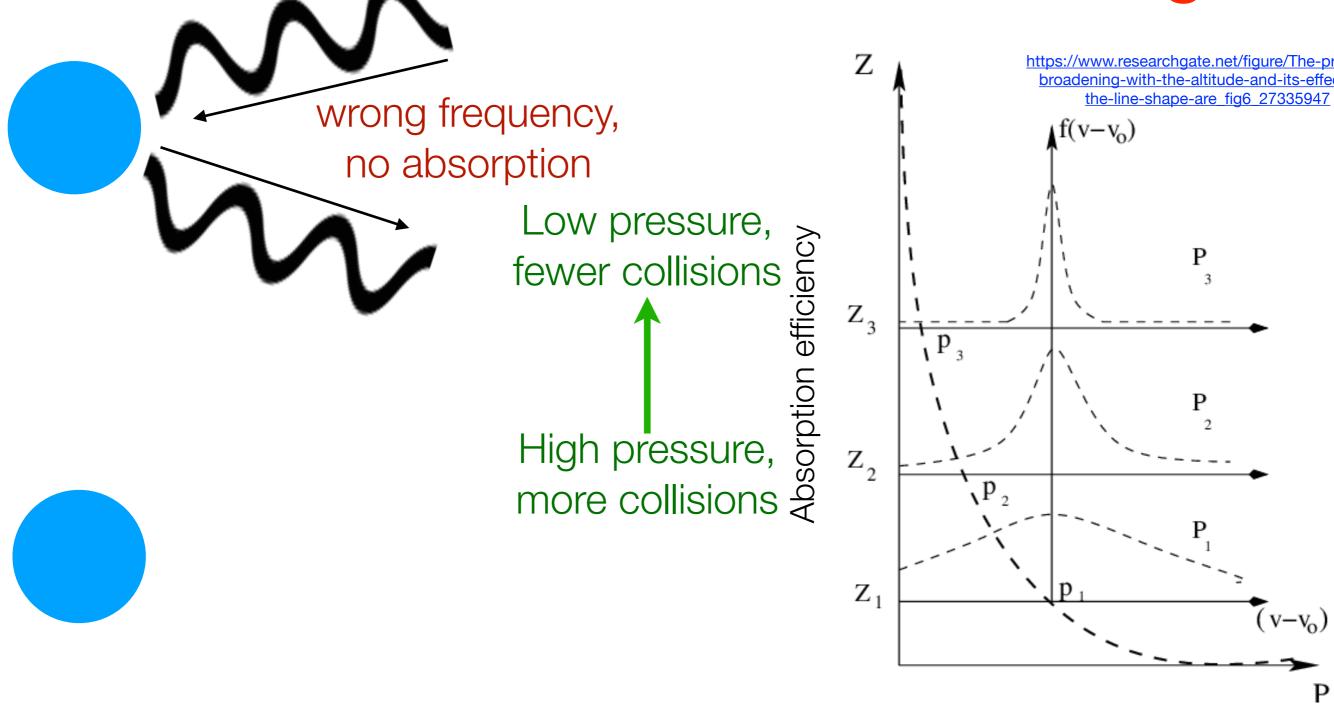


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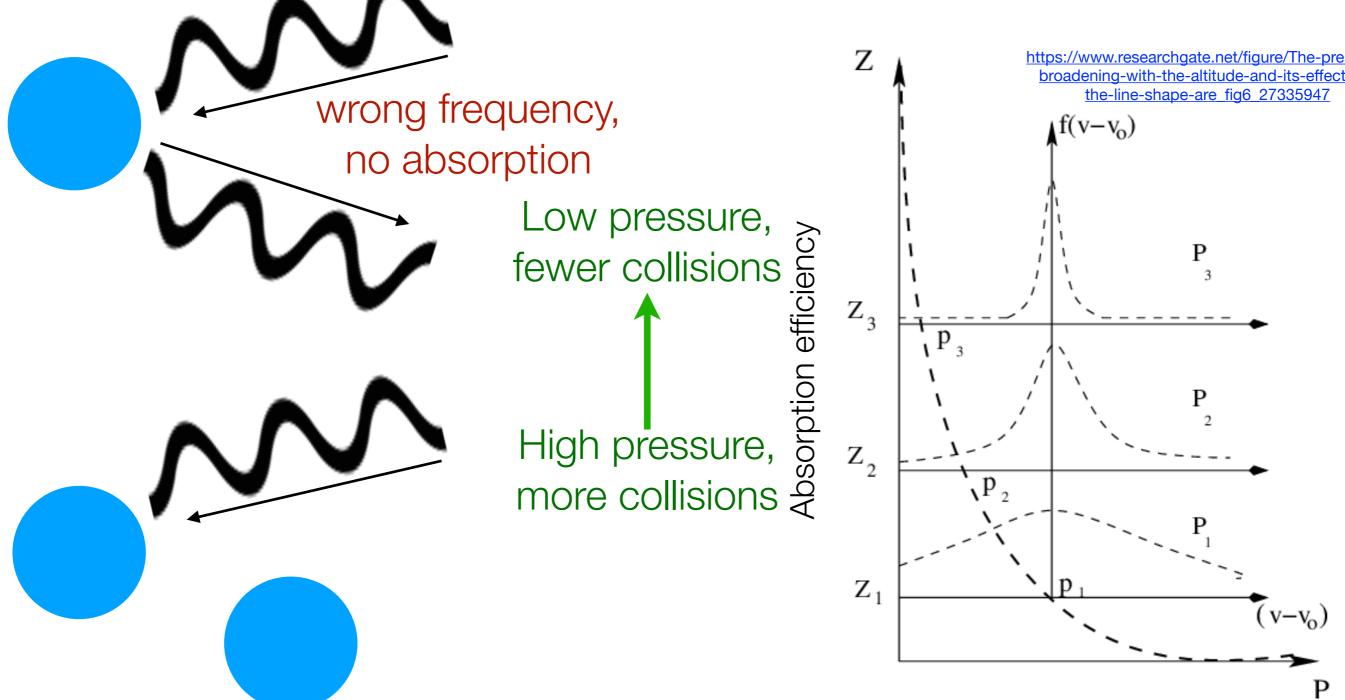


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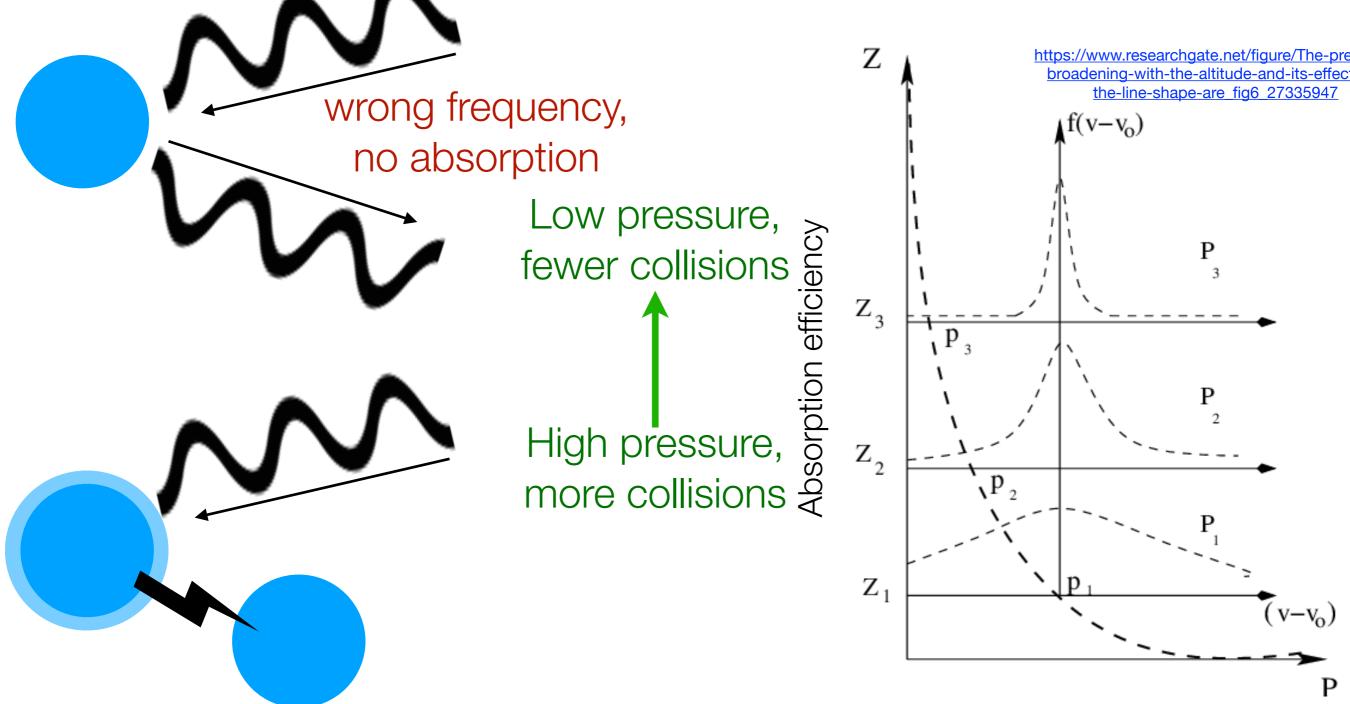
Pressure/Collisional Broadening https://www.researchgate.net/figure/The-pi broadening-with-the-altitude-and-it the-line-shape-are fig6 27335947 wrong frequency, $f(v-v_0)$ no absorption Low pressure, fewer collisions Z_3 High pressure, of more collisions

The collisions of gas molecules cause the widening of absorption lines:

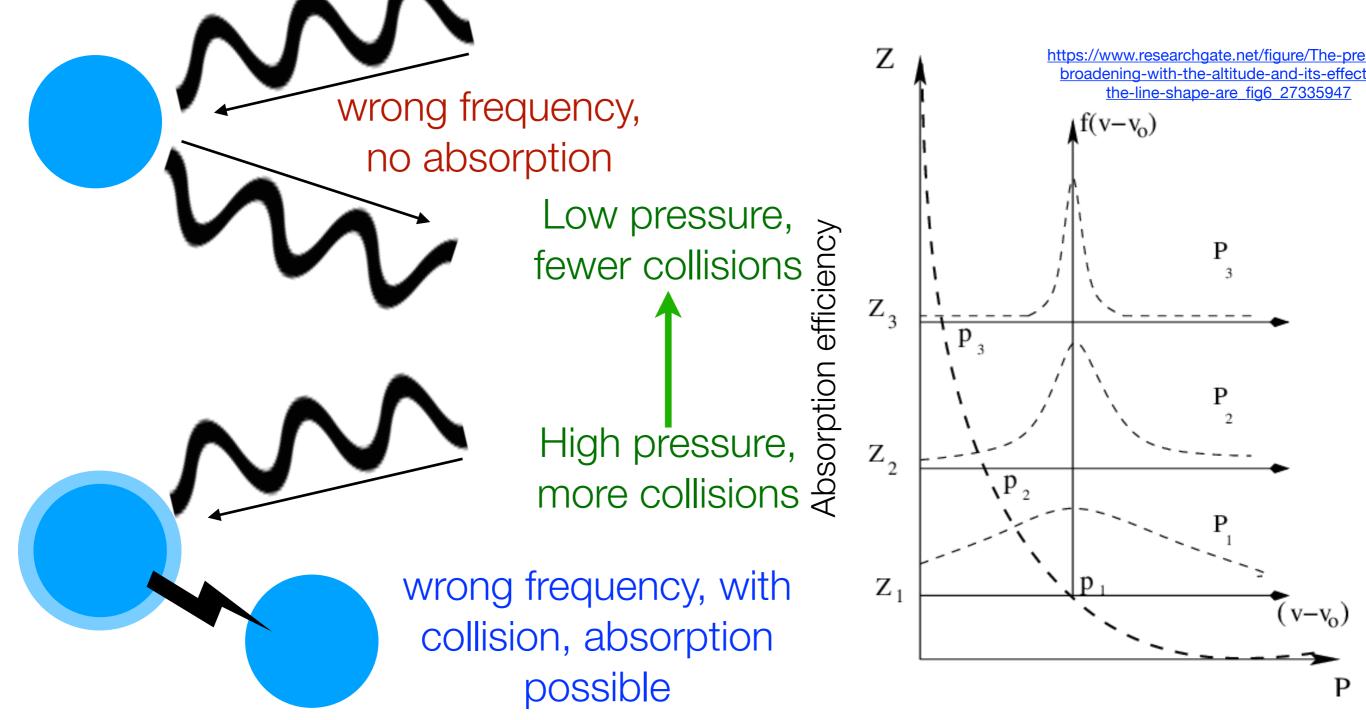
 Z_1



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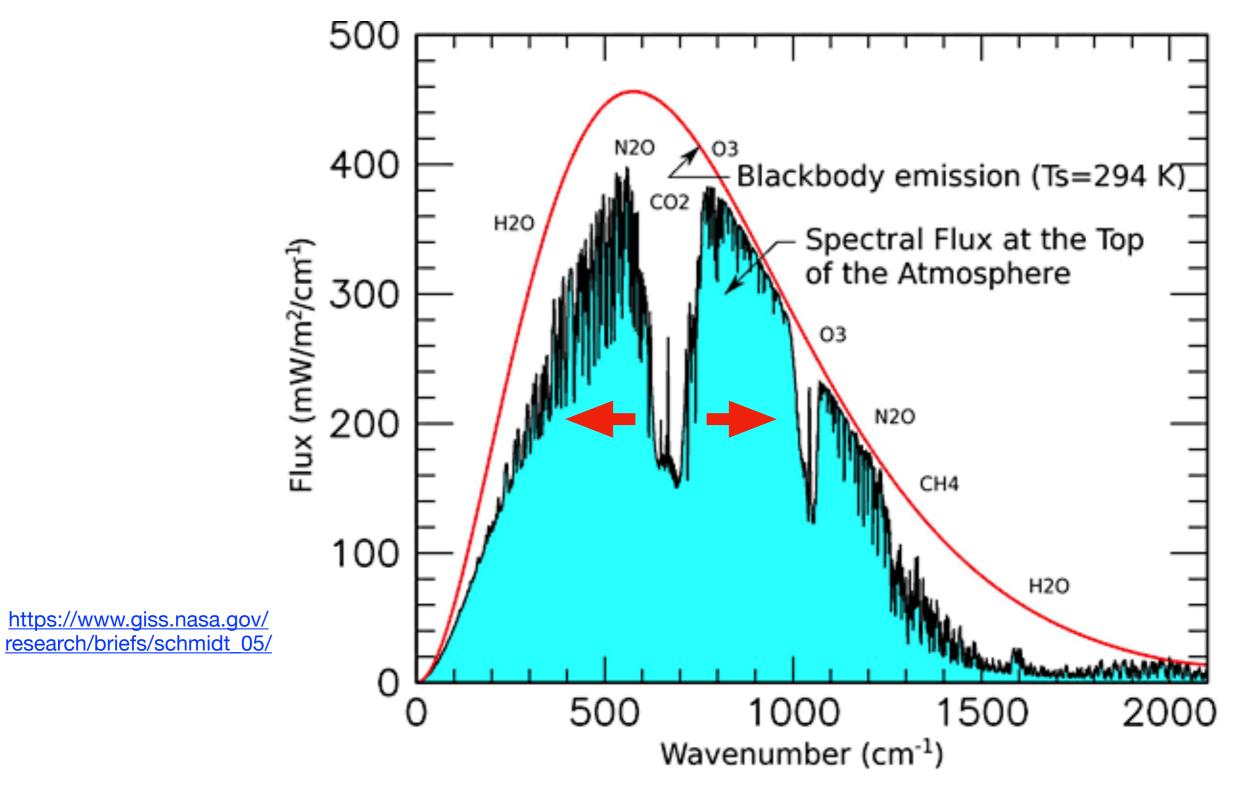


The collisions of gas molecules cause the widening of absorption lines:



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Outgoing Longwave Radiation (OLR)



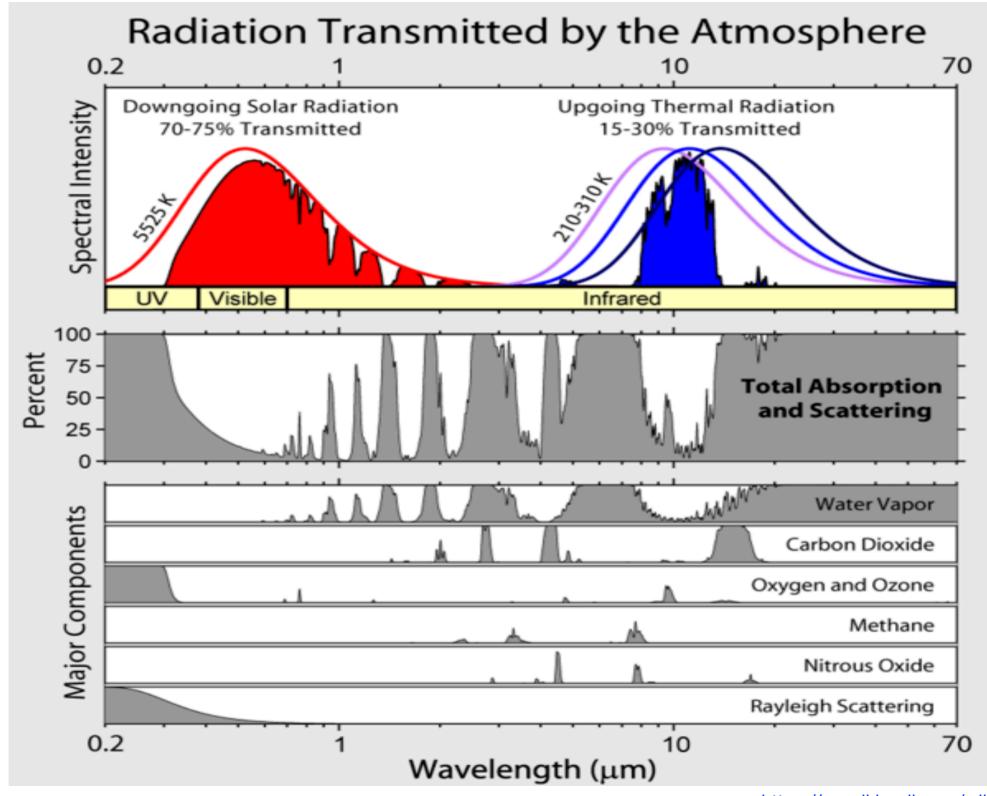
note absorption windows...

Workshop #3 radiative forcing

Notes sections 2.2.4, 2.2.5, 2.2.6: logarithmic dependence, global warming potential, water vapor feedback

(use next 4 slides)

IR absorption of the major greenhouse gases



https://en.wikipedia.org/wiki/Absorption_band

CO₂ & water vapor absorbs the most IR, at different wavelengths

GWP: the time-integrated RF due to a pulse emission of a GHG, relative to a pulse emission of an equal mass of CO2

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GWP of a greenhouse gas is its forcing effect relative to CO₂, considering its lifetime and its strength as an absorber, over a specified time horizon (TH).

GWP: the time-integrated RF due to a pulse emission of a

CO2, considering its lifetime and its strength as an absorber, over a specified time horizon (TH).

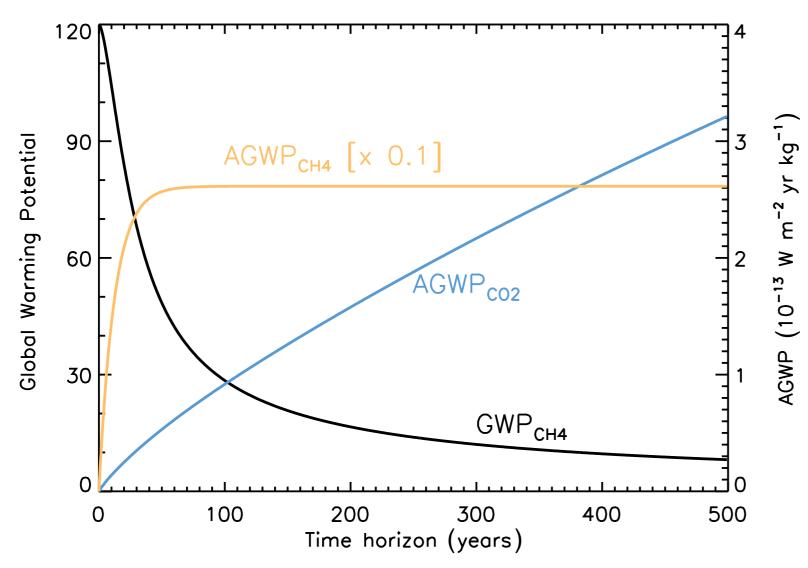
GWP: the time-integrated in GGO to a part of the foreign of solution of an equal mass of CO2 $\int_0^{TH} a_x \cdot \left[x(t)\right] dt$. GWP(x) = $\frac{\int_0^{TH} a_x \cdot \left[x(t)\right] dt}{\int_0^{TH} a_r \cdot \left[r(t)\right] dt}$.

x(t): time-dependent decay of a GHG; r(t) that of CO₂; ax: RF per 1 kg (W/m²/kg).

GWP: the time-integrated RF due to a pulse emission of a GWP: the time-integrated for our control paids of GHG, relative to a pulse emission of an equal mass of CO2 $GWP(x) = \frac{\int_0^{TH} a_x \cdot \left[x(t)\right] dt}{\int_0^{TH} a_r \cdot \left[r(t)\right] dt}.$

$$GWP(x) = \frac{1}{2}$$

CO2, considering its lifetime and its strength as an absorber, over a specified time horizon (TH).



x(t): time-dependent decay of a GHG; r(t) that of CO₂; ax: RF per 1 kg (W/m²/kg).

$$\Delta T(t) = \sum_{i} \int_{0}^{t} E_{i}(s) A(s)$$

IPCC, Climate Change 2013, Chapter 8

GWP: the time-integrated RF due to a pulse emission of a

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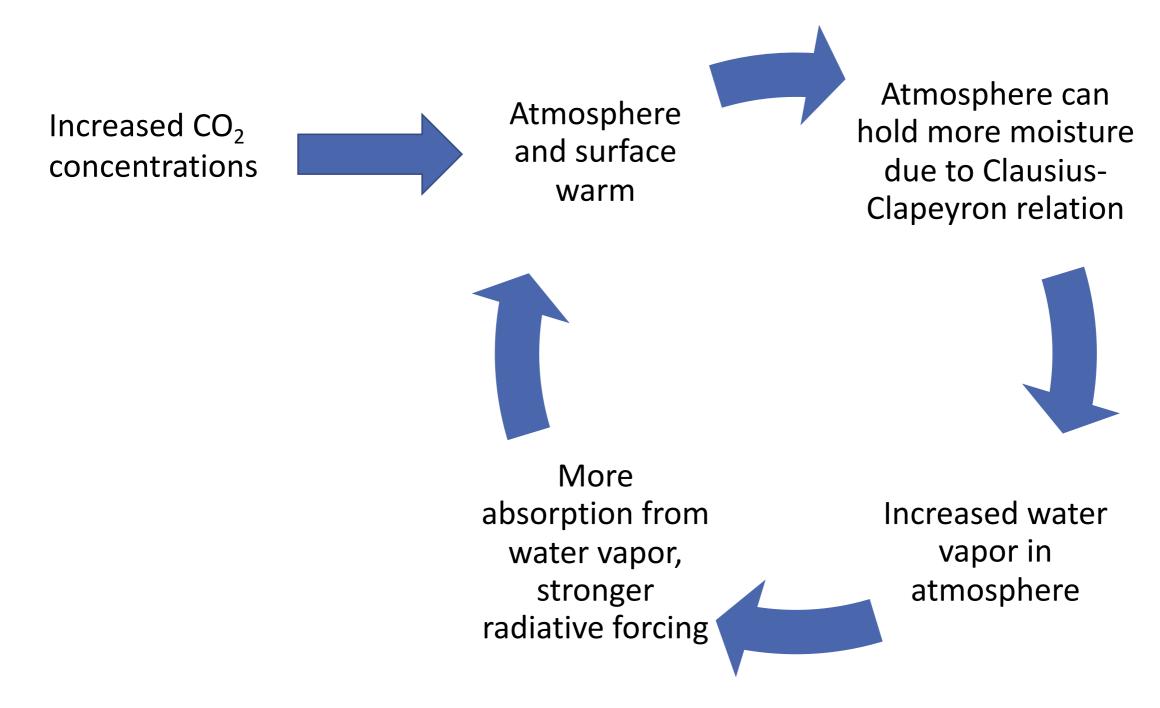
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ax: RF per 1 kg (W/m²/kg).

GWP values and lifetimes	Lifetime in years	Global Warming Potential (GWP)		
		20 years	100 years	500 years
Methane	12.4	86	34	
_				
Nitrous oxide (N ₂ O)	121.0	268	298	
Nitrous oxide (N ₂ O)	121.0	264	265	
HFC-134a (hydrofluorocarbon)	13.4	3790	1550	

wikipedia

Water vapor feedback



Direct radiative forcing of absorption by water vapor molecules reinforces that by CO₂ via the water vapor feedback

Workshops #4,5: logarithmic dependence global warming potential

Climate cools/warms to reach radiative equilibrium: incoming SW = outgoing LW

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- Surface temperature response scales (only/weakly) logarithmically with CO₂ concentration: warming due to each CO₂ doubling is approximately the same

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- Surface temperature response scales (only/weakly) logarithmically with CO₂ concentration: warming due to each CO₂ doubling is approximately the same
- Greenhouse effect of water vapor leads to a positive feedback on CO₂ increase

- Climate cools/warms to reach radiative equilibrium: incoming SW = outgoing LW
- Adding greenhouse gases to the atmosphere traps LW radiation, so that incoming SW > outgoing LW, the Earth must warm to get back to equilibrium
- Atmosphere is mostly saturated wrt absorption by CO₂. Anthropogenic CO₂ causes warming because temperature decreases with height (lapse rate), raising the level of last absorption; & bec of increased absorption in margins of absorption bands
- CO₂ is a particularly effective greenhouse gas because it absorbs wavelengths of radiation at which the Earth emits most strongly, & where H₂O is less effective
- Surface temperature response scales (only/weakly) logarithmically with CO₂ concentration: warming due to each CO₂ doubling is approximately the same
- Greenhouse effect of water vapor leads to a positive feedback on CO2 increase
- The Global warming potential of other GHGs depends on both their efficiency and life time, CO₂ has an especially long life time in the atmosphere

The End