

Two lessons from past warm climates

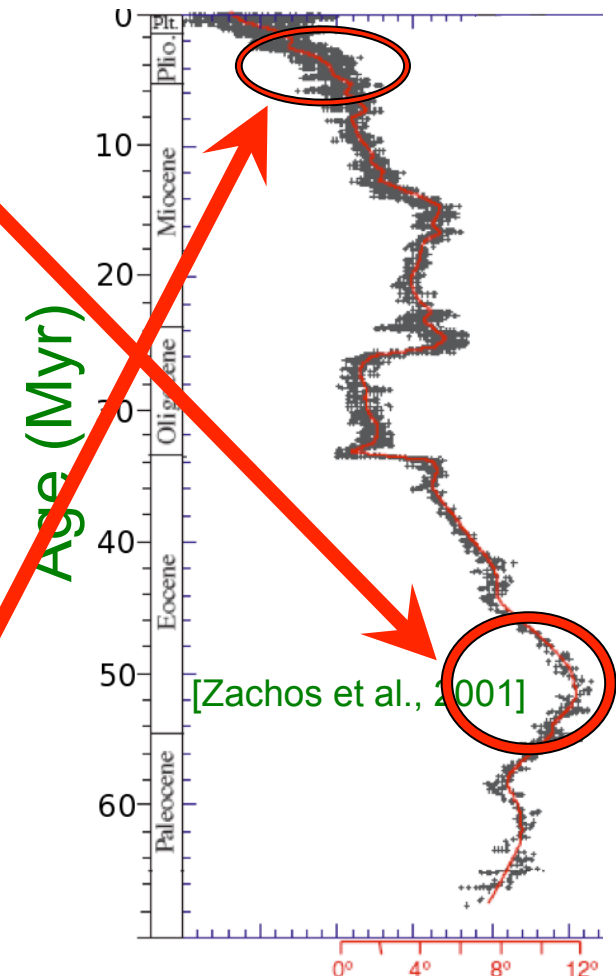
1. 34-146 Myr ago: warming of the high latitudes by a convective cloud feedback.

Work of **Dorian Abbot**; recent updates using SPCAM: **David Randall, Mark Branson**.

1. 2-5 Myr ago: *Permanent El Nino: due to atmospheric superrotation?*

Work of **Nathan Arnold**, with **Brian Farrell**; Rossby wave resonance and SPCAM results

Gradual cooling over past 55Myr



Deep ocean
Temperature

A High Latitude Convective Cloud Feedback and Equable Climates

Dorian S. Abbot

Recent SPCAM updates: with
David Randall, Mark Branson



LESSON 1

Hadrosaurus – Cretaceous [Karen Carr]

EON	ERA	PERIOD	EPOCH	Ma
Phanerozoic	Cenozoic	Quaternary	Holocene	0.01 –
			Pleistocene	0.8 –
		Tertiary	Early	1.8 –
			Late	3.6 –
			Pliocene	5.3 –
			Early	11.2 –
			Late	16.4 –
			Miocene	23.7 –
			Early	28.5 –
			Late	33.7 –
	Mesozoic	Cretaceous	Late	41.3 –
			Early	49.0 –
		Jurassic	Middle	54.8 –
			Early	61.0 –
		Triassic	Late	65.0 –
			Early	99.0 –
	Paleozoic	Permian	Late	144 –
			Early	159 –
		Pennsylvanian	Middle	180 –
			Early	206 –
		Mississippian	Late	227 –
			Early	242 –
		Devonian	Late	248 –
			Early	256 –
		Silurian	Late	290 –
			Early	323 –
		Ordovician	Late	354 –
			Early	370 –
		Cambrian	Middle	391 –
			Early	417 –
Precambrian	Proterozoic	Archean	Late	423 –
			Early	443 –
			Middle	458 –
	Archean	Proterozoic	Early	470 –
			Late	490 –
			Middle	500 –
			Early	512 –

[Abbot & Tziperman, 2008,9: QJRM, GRL, JAS, J. Climate]

Outline: Eocene (50 Myr)

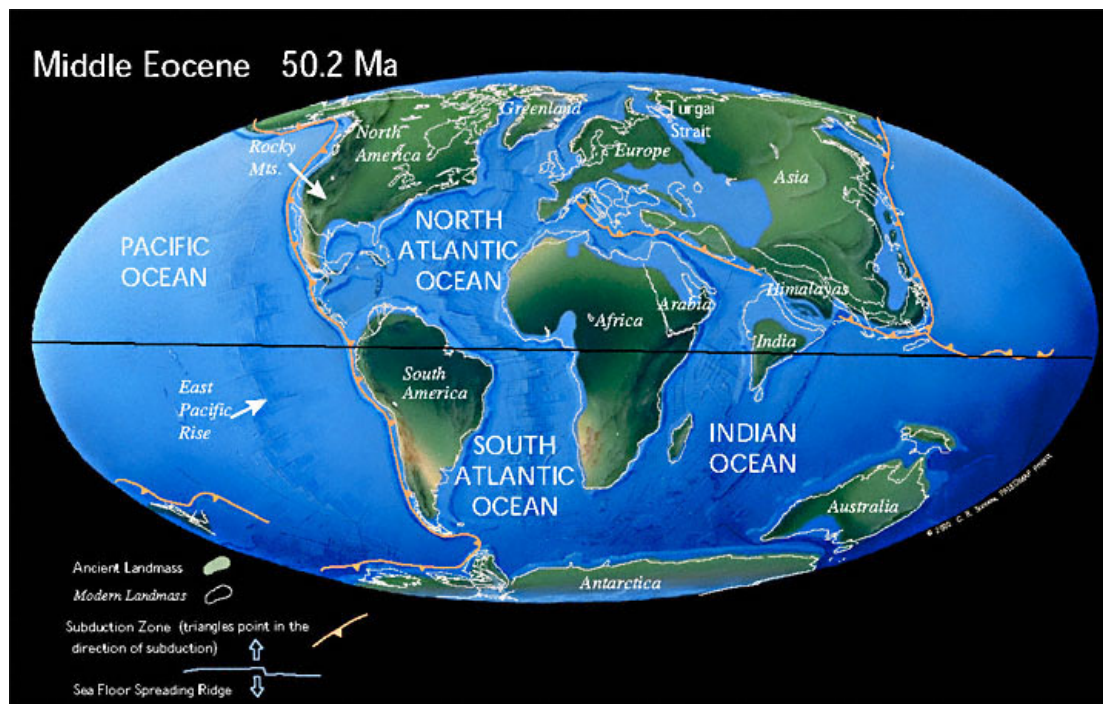
warmth & a convective-cloud feedback

- 1) **Observations:** very warm climate 146-34 Myr ago; what's the mechanism..??
- 2) **Previous explanations...**
- 3) **Our mechanism:** a qualitatively different state of the atmosphere, with tropical-like deep atmospheric convection and high tropospheric clouds at mid- to high-latitudes providing a strong greenhouse effect.
- 4) **Why should you care**

Observations (1st/3): warm climate ~146-34 Ma

- High global mean temperature
- Low Equator-pole temperature difference: $\sim 25^{\circ}\text{C}$ (now $\sim 45^{\circ}\text{C}$).
- above freezing winter temperatures @ 60N, interior of N. America (now -30°C);
- Weak high-lat seasonality
- No significant ice
- Tropical SSTs $>\approx$ modern
- Warm deep ocean: 15°C
- $\text{CO}_2=500\text{-}5,000$ ppm(??)

“Equable” climate \equiv warm poles, mild winters



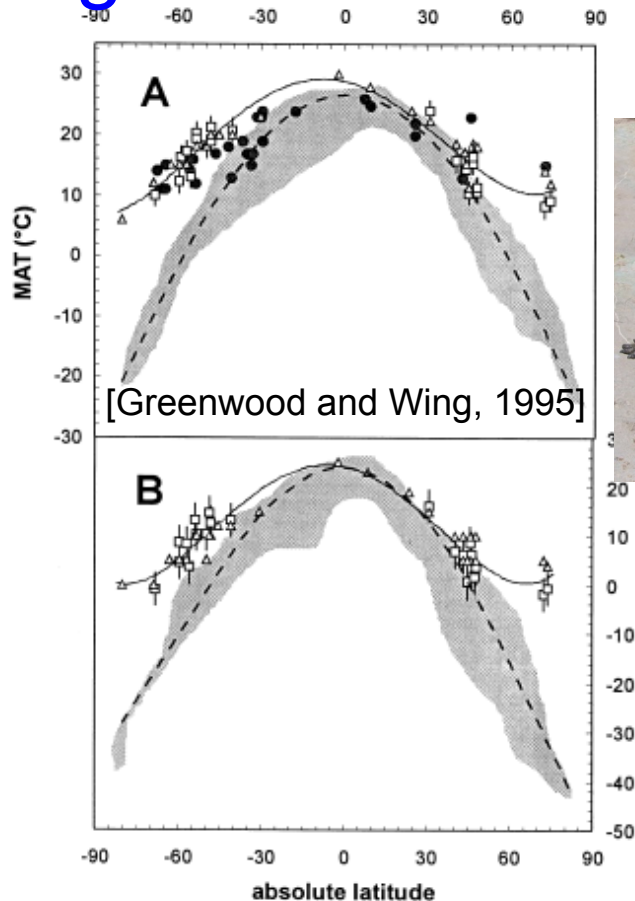
[146-34 Ma ago:
Cretaceous -
Paleocene -
Eocene]

Eocene: modern-like
continental configuration.

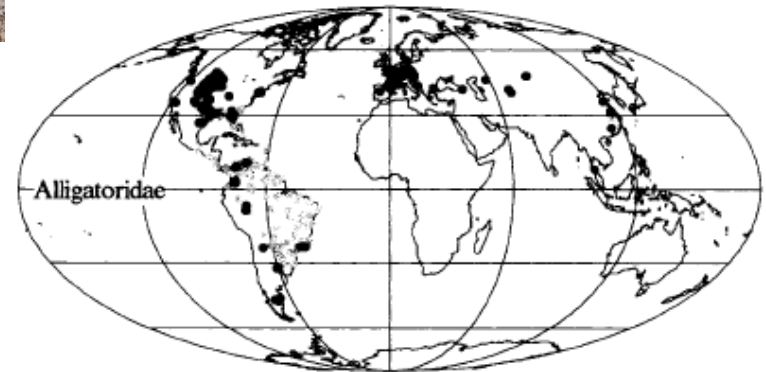
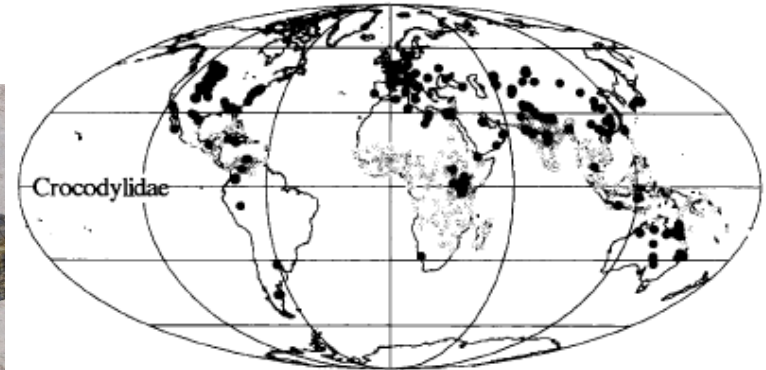
Observations (2nd/3): warm climate ~146-~34Ma

Cool tropics, warm high-latitudes

Crocodiles in Greenland,
Palm trees in Wyoming!



■ - Modern land temp.
● △ □ - Eocene SST;
NLR&LMA, CLAMP



Crocodiles need: Mean annual $T > 14.2^{\circ}\text{C}$ & Cold month mean $> 5.5^{\circ}\text{C}$ [Markwick, 1998]

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Previously proposed mechanisms

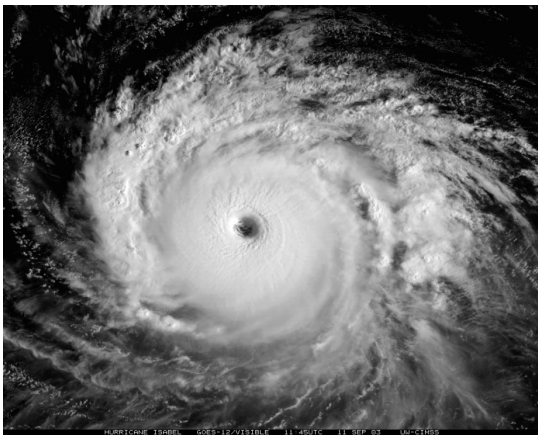
Ocean

Atmosphere

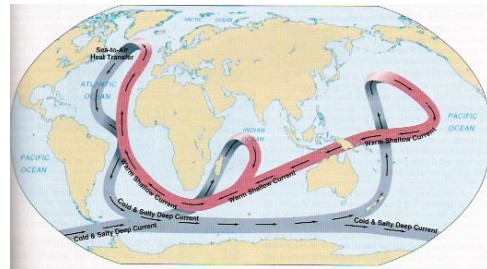
Stratos. clouds

Warm climate → stronger Hurricanes → stronger oceanic thermohaline circulation → more heat transport to the pole → warmer poles [Emanuel 2002]:

Stronger hurricanes

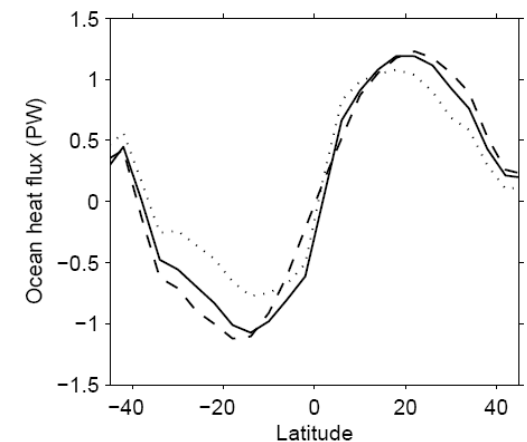
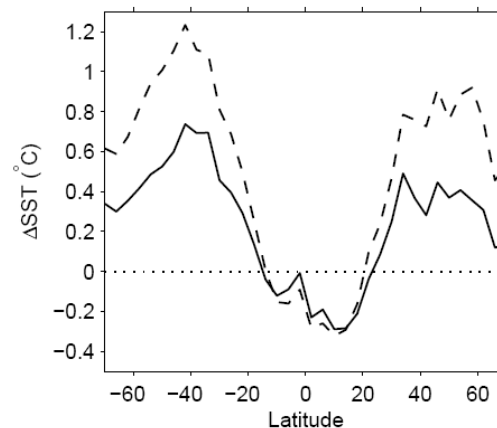


Thermohaline circulation



Warmer high latitudes

However, this feedback may cool the tropics better than warm the high latitudes [Korty & Emanuel 2007]



Previously proposed mechanisms

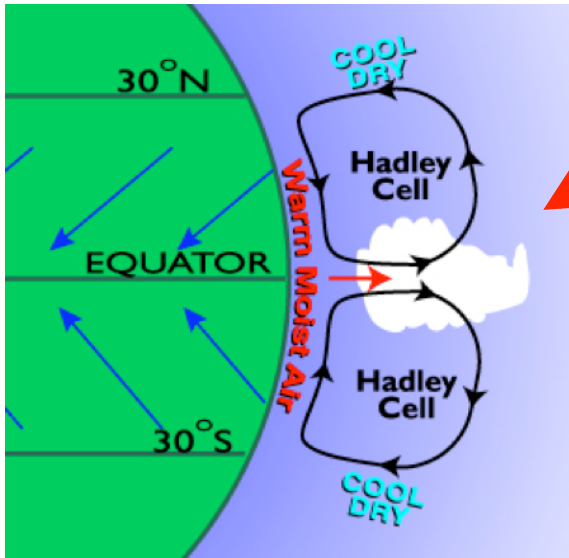
Ocean

Atmosphere

Stratos. clouds

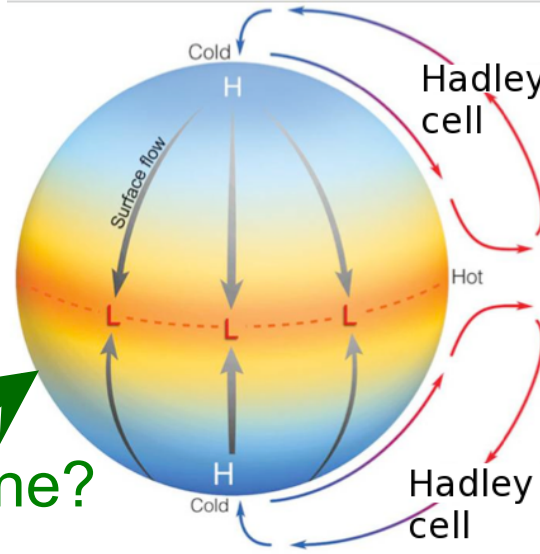
Equator-to-pole Hadley cell:

B. Farrell, 1990



Today

Eocene?



Giant Pterosaur: strong Easterly zone could help them fly... ☺?

Like Venus: large Hadley cell due to slow rotation; Poles warm as Equator



But: requires X8 angular momentum dissipation; Based on now challenged theory of [Schneider 77; Held Hou 80]

Back of the envelope...

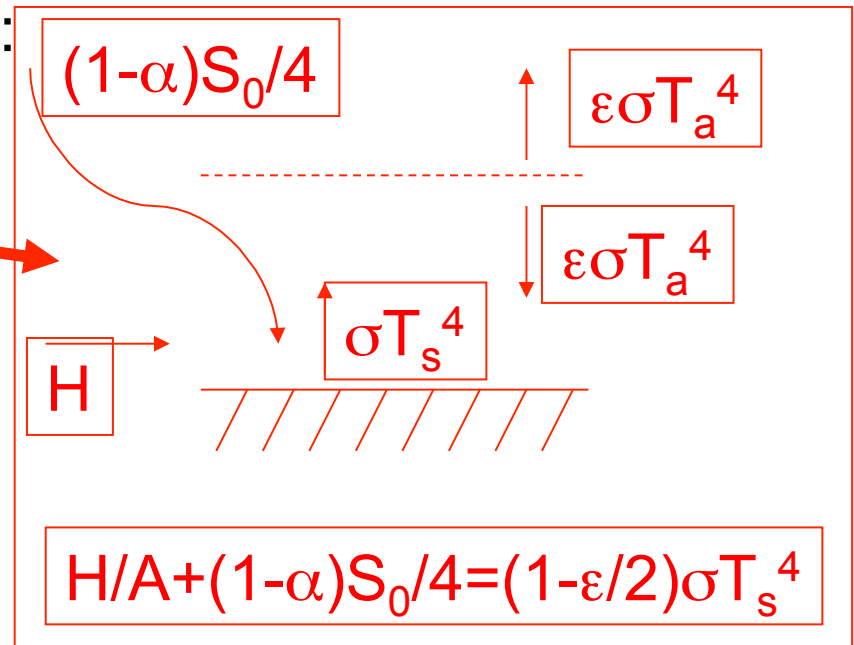
It is very difficult to reduce the equator-to-pole temperature gradient by increasing CO₂ or meridional heat fluxes; far easier using long-wave emissivity (**clouds!**):



Modern climate:

T [°C]	H [PW]	ϵ []	α []
-8.0	3.5	0.60	0.55

Energy-balance
model



Increasing high latitude temperature:

ΔT [°C]	ΔH [PW]	$\Delta \epsilon$ []	[CO ₂] _{dry} [ppm]	[CO ₂] _{wet} [ppm]	$\Delta \alpha$ []
10.0	1.1	0.20	$x2^5 \approx 9 \times 10^3$	$x2^{2.5} \approx 2 \times 10^3$	-0.15
15.0	1.7	0.28	$x2^{7.5} \approx 5 \times 10^4$	$x2^{3.75} \approx 4 \times 10^3$	-0.23

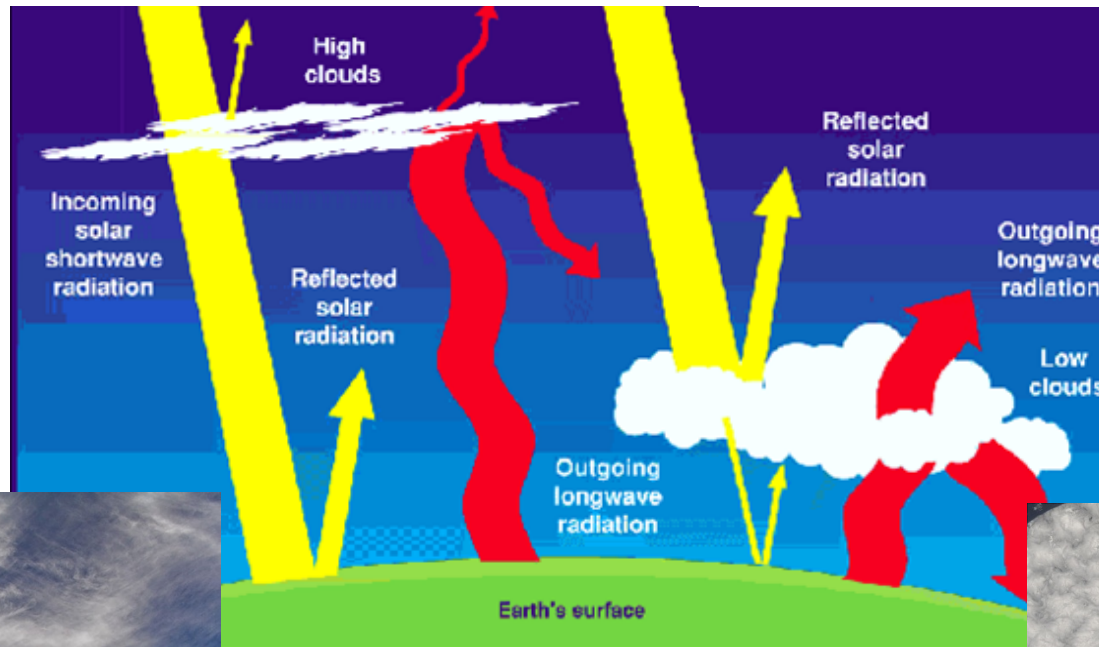
A reminder: Cloud feedbacks

High clouds (cirrus)

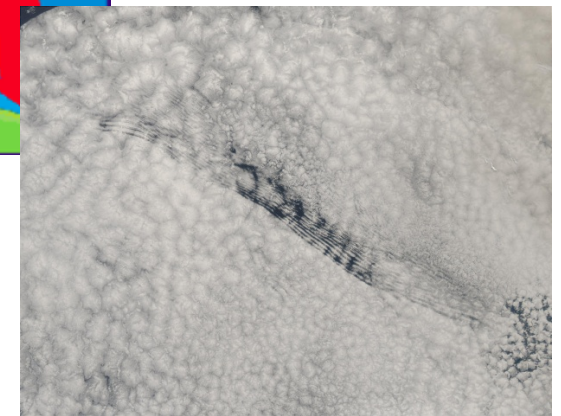
- low albedo, high emissivity
- High altitude (>8 km)
- Warming effect on climate

Low clouds (marine stratus)

- High albedo
- Low altitude (<1km)
- Cooling effect on climate



➔ high clouds may help explain equable climate



[End of] Previously proposed mechanisms

Ocean

Atmosphere

Stratos. clouds

Polar Stratospheric Clouds (PSCs), at 15-25km, have a strong greenhouse effect! Formed via methane-moistening of stratos.

- Eocene PSCs due to methane [Sloan'92];

BUT: methane source not clear

- PSCs due cooling & *weakening* of Brewer-Dobson stratospheric circulation [Kirk-Davidoff et al 2002] : **BUT stratospheric circulation may *increase* in warm climate** [Korty & Emanuel 2007]



[PSCs at dusk over the Arctic region of Sweden]

http://www.nasa.gov/images/content/65932main_sageii_psc_640x480.jpg

An interim summary...

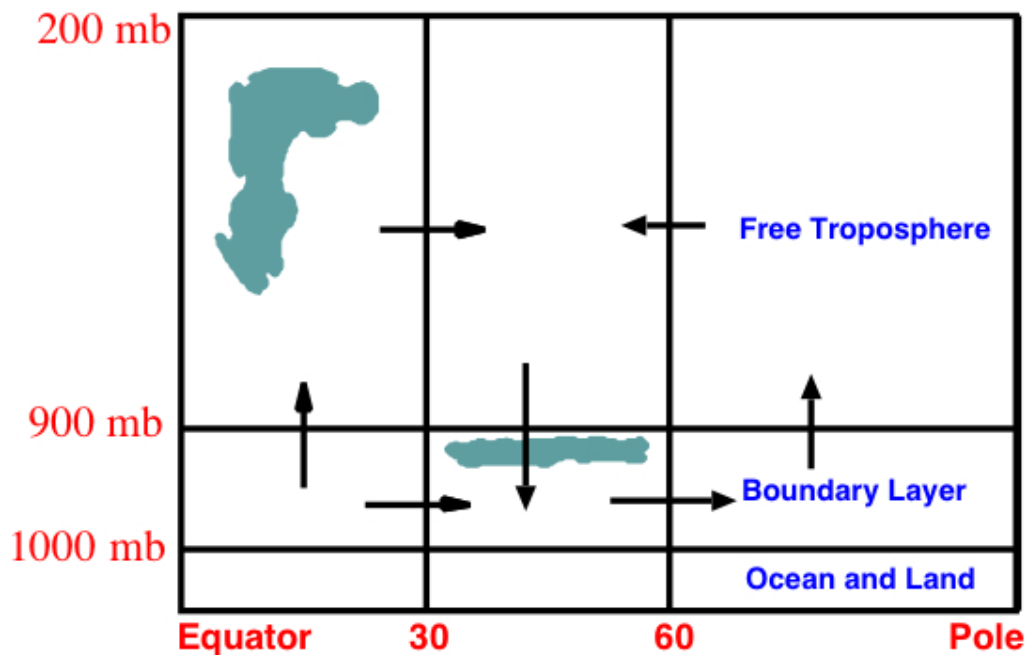
- High CO₂, water vapor, increased poleward heat flux by ocean/atmosphere cannot explain equable climates (Eocene, 50Myr)
- Perhaps clouds? Polar *Stratospheric* clouds not so simple...

Outline: Eocene (50 Myr) warmth & a convective-cloud feedback

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First step: a “toy” model (1st / 2)

- Zonally averaged
- Equator to pole
- Two levels: Boundary Layer + Free Troposphere
- Mixed layer ocean
- Non-linear momentum eqns
- Merid resolution: 3 columns
- Prognostic dry static energy & water vapor
- Simple land surface
- Advection
- Diffusive eddies
- Convection + Precipitation
- Clouds: convective and large-scale
- Radiation: SW, LW, CO₂, water vapor, clouds...
- Surface fluxes



The toy model (2nd / 2): equations

Dry Static Energy and Moisture Equations

$$\begin{aligned}\frac{\partial DSE}{\partial t} &= \frac{\omega}{\Delta P} DSE + \frac{v}{\Delta y} DSE + D(\theta) \frac{\partial DSE}{\partial \theta} + L_v \frac{q - \tilde{q}}{\tau} + k_{con}(DSE_1 - DSE_2) \\ &\quad - \delta_{k2} \epsilon_{re} L_v \frac{q_1 - \tilde{q}}{\tau} + \epsilon A (F_{in} - 2\sigma T^4) + \delta_{k2} \rho_2 C_{SH} (T_S - \theta_2) \\ \frac{\partial q}{\partial t} &= \frac{\omega}{\Delta P} q + \frac{v}{\Delta y} q + D(\theta) \frac{\partial q}{\partial \theta} - \frac{q - \tilde{q}}{\tau} + k_{con}(q_1 - q_2) \\ &\quad + \delta_{k2} \epsilon_{re} \frac{q_1 - \tilde{q}}{\tau} + \delta_{k2} \rho_2 C_{LH} (q^*(T_S) - q_2)\end{aligned}$$

Equations of Motion: angular momentum conservation included

$$\begin{aligned}u_t + \frac{1}{a \cos(\theta)} \frac{\partial}{\partial \theta} (uv \cos(\theta)) - uv \tan(\theta)/a + \frac{\partial}{\partial p} (\omega u) - 2\Omega \sin(\theta) v &= v \frac{\partial^2 u}{\partial \theta^2} - \delta_{k2} r u \\ v_t + 2\Omega \sin(\theta) u &= -\frac{1}{a} \frac{\partial \phi}{\partial \theta} + v \frac{\partial^2 v}{\partial \theta^2} - \delta_{k2} r v \\ \phi_p &= -\alpha \\ \frac{1}{a \cos(\theta)} \frac{\partial}{\partial \theta} (v \cos(\theta)) + \omega_p &= 0 \\ p\alpha &= R^* T\end{aligned}$$

Model experiments & results: summary

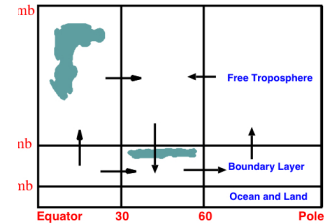
Model experiments:

Slowly increase CO₂ to extreme values & then decrease it

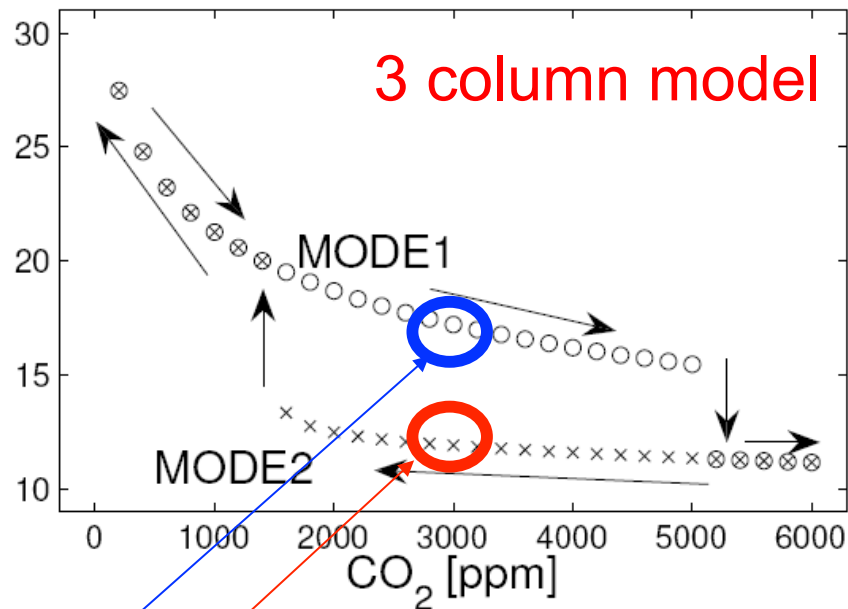
Results:

A ***qualitatively*** different climate regime at sufficiently high CO₂, warm high latitudes and low equator-to pole temperature difference.

Results: 2 modes of atmospheric dynamics; [& multiple equilibria at a given CO₂, hysteresis]



Equator to pole temperature difference (EPTD, °K).



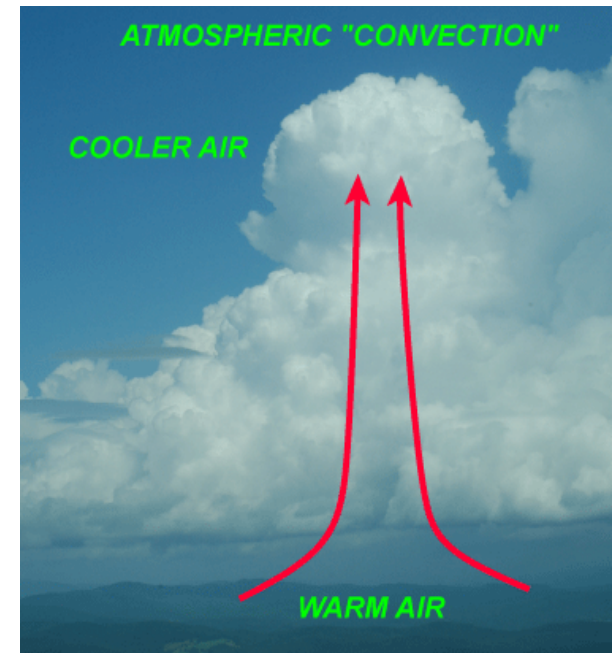
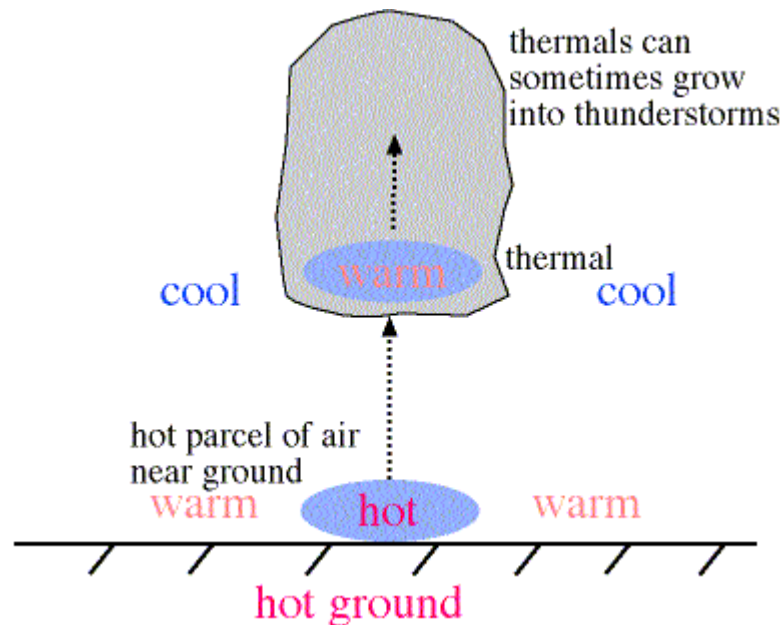
“Present-day” solution: high EPTD, colder; non convecting

“Equable” solution: low EPTD, warm; convecting, high clouds

Arrows: path of solution if CO₂ slowly increased then decreased.

A reminder: Atmospheric convection

1. Air parcel in lower atmosphere rises up
2. It expands, cools, and water vapor condenses
3. Condensation leads to latent heat release, air parcel heats
4. Parcel becomes warmer, lighter, more buoyant, and rises even more → positive feedback, instability
5. Condensation creates clouds, rain



<http://apollo.lsc.vsc.edu/classes/met130/>

Why does convection start at high latitudes @ high CO₂?

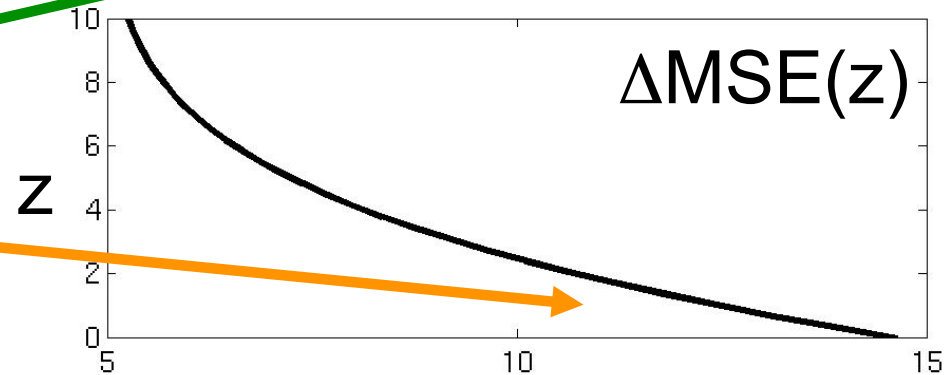
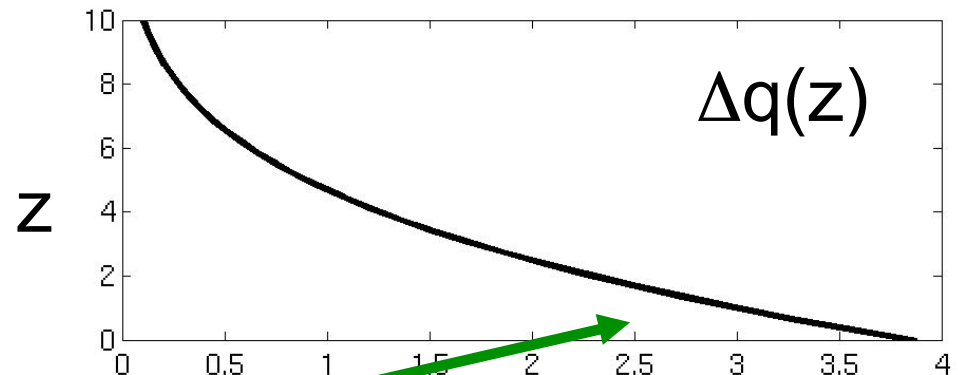
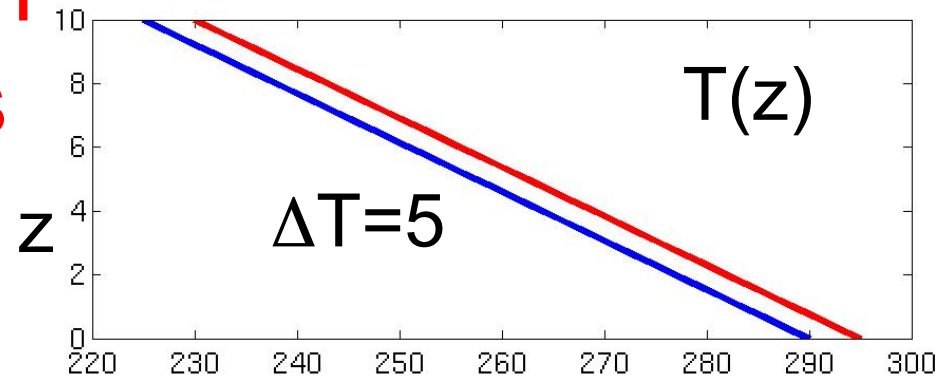
→ as CO₂ increases, moist stability decreases & eventually leads to convection.

why: Clausius-Clapeyron

→ increasing CO₂ leads to larger moisture

& Moist Static Energy increases in lower atmos,

→ convection

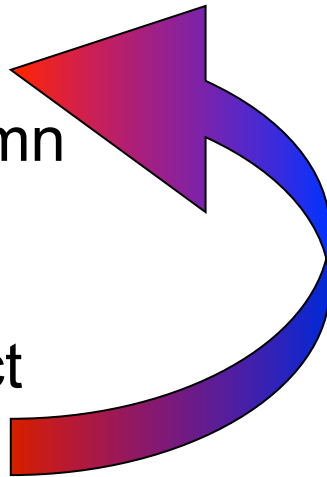


A summary of the proposed mechanism for equable-climate via high latitude convection*

warmer surface

- unstable air column
- deep convection
- high clouds
- greenhouse effect

warmer surface



This positive feedback supports 2 states:

(1) Equable (high lat deep convection, high clouds & warm)

(2) present-day-like: deep convection only at equator

→ **Positive feedback!**

Low CO₂: only present-like state;

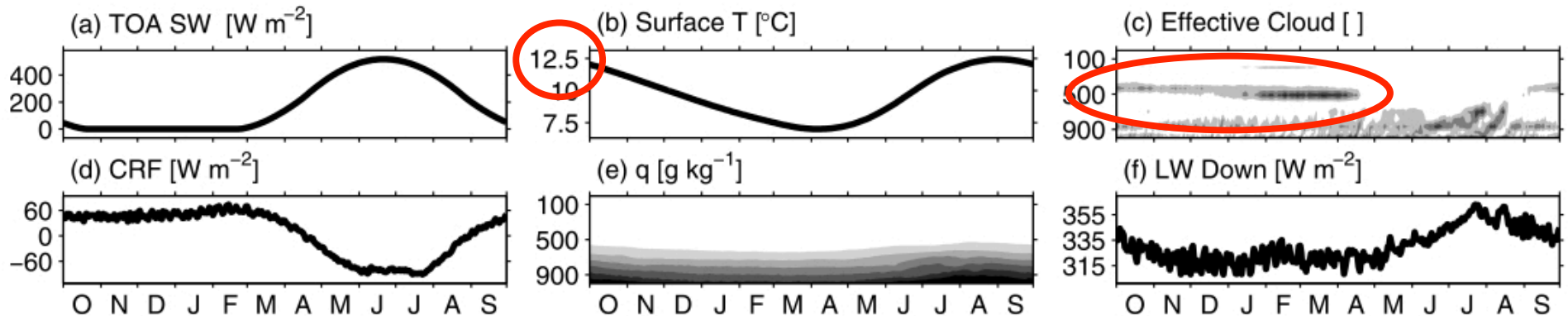
High CO₂: equable only

Intermediate CO₂ : both (may be sensitive to model details...)

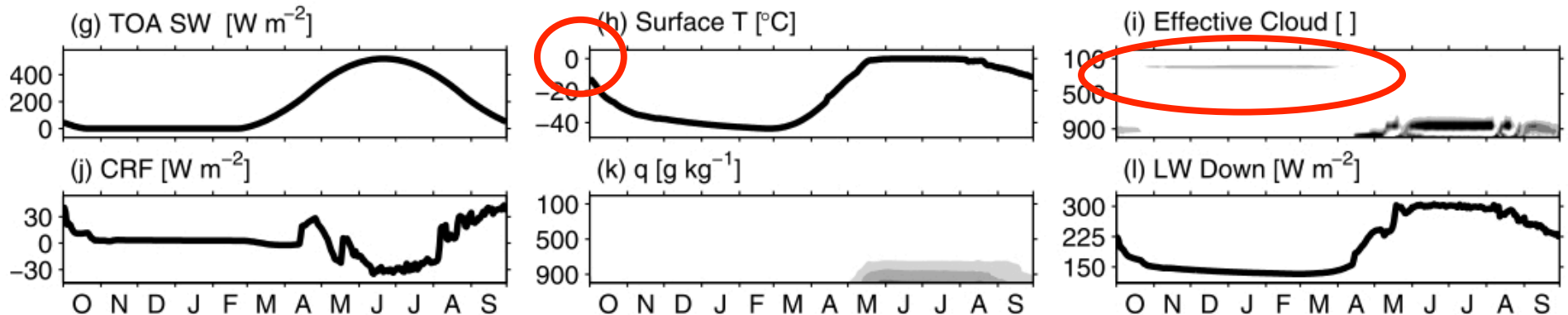
*(a related suggestion was made by Huber et al 1999)

SCAM supports TOY MODEL'S mechanism for high-latitude warmth during equable climates

Ice-Free State



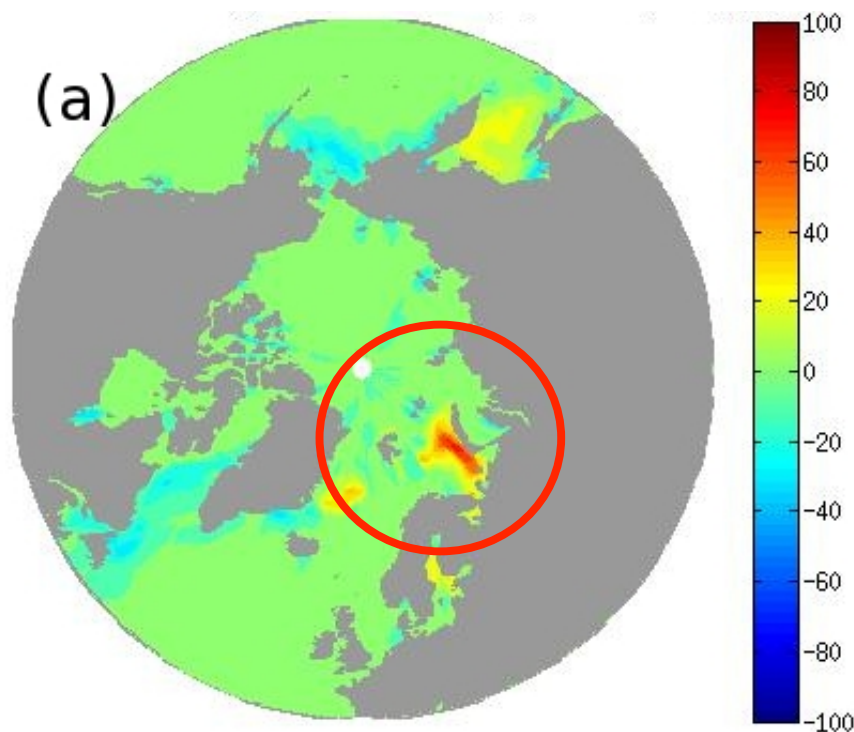
Ice State



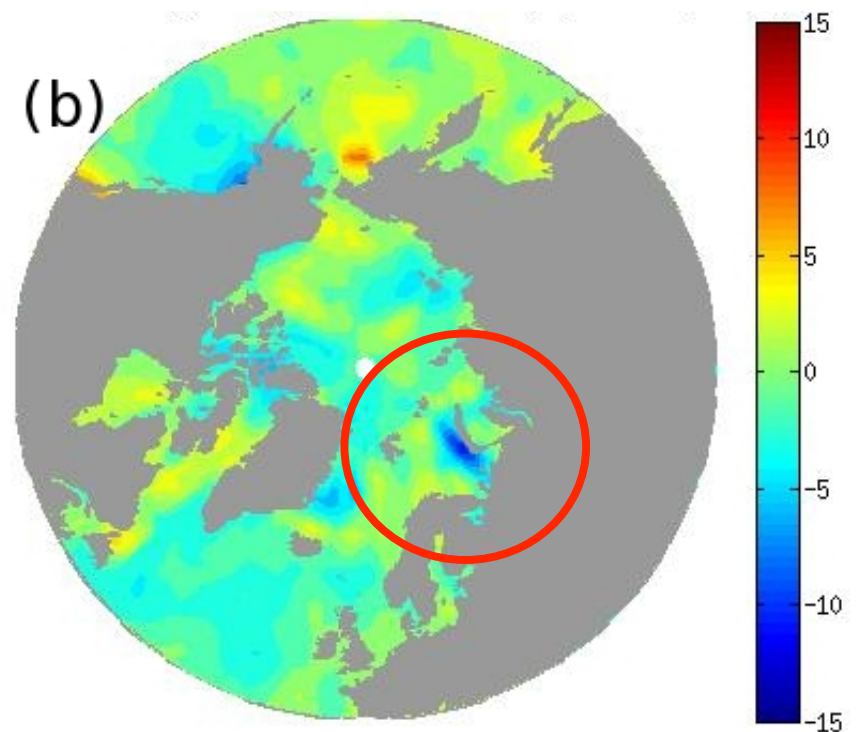
SCAM, $\text{CO}_2=1000\text{ppms}$, different I.C. \rightarrow 2 different *seasonal* states, with & w/o sea ice. Also: multiple equilibria, hysteresis

Additional evidence using reanalysis products with Kerry Emanuel, Ben Leibowicz

- Consider times with a high/ low sea ice cover and examine cloud radiative forcing then.
- ➔ Results indicate a clear correlation, so feedback seems active in today's atmosphere



High sea ice cover
anomaly during winter



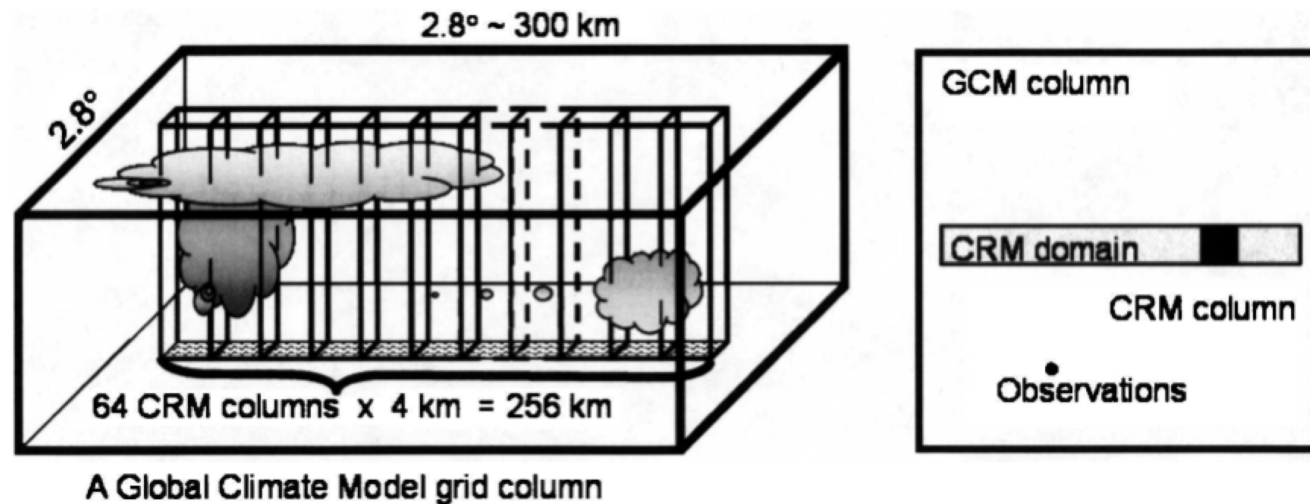
➔ Negative CRF anomaly
during same years

Moving (far) beyond simple models

“Super parameterized” community atmospheric general circulation model (= **SPCAM**, Grabowski, Randall & colleagues)

1 MAY 2006

OVTCHINNIKOV ET AL.

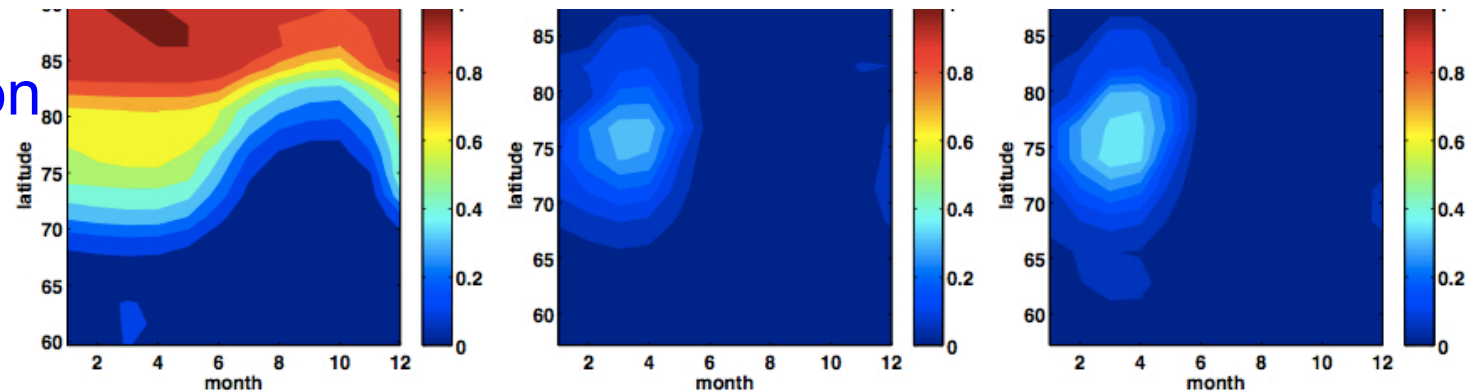


99% of computation time is spent on cloud resolving model which serves as the convection parameterization. → Cost is GCMx100

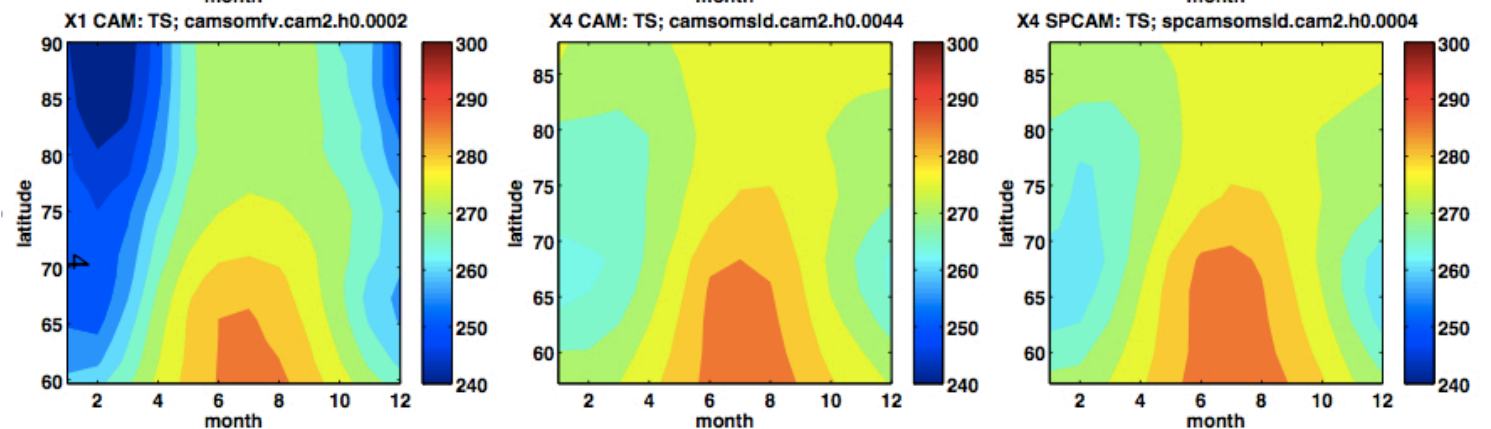
Preliminary SPCAM results: latitude (60N-90N) vs month

X1CO2,CAM; X4CO2,CAM; X4CO2, SPCAM

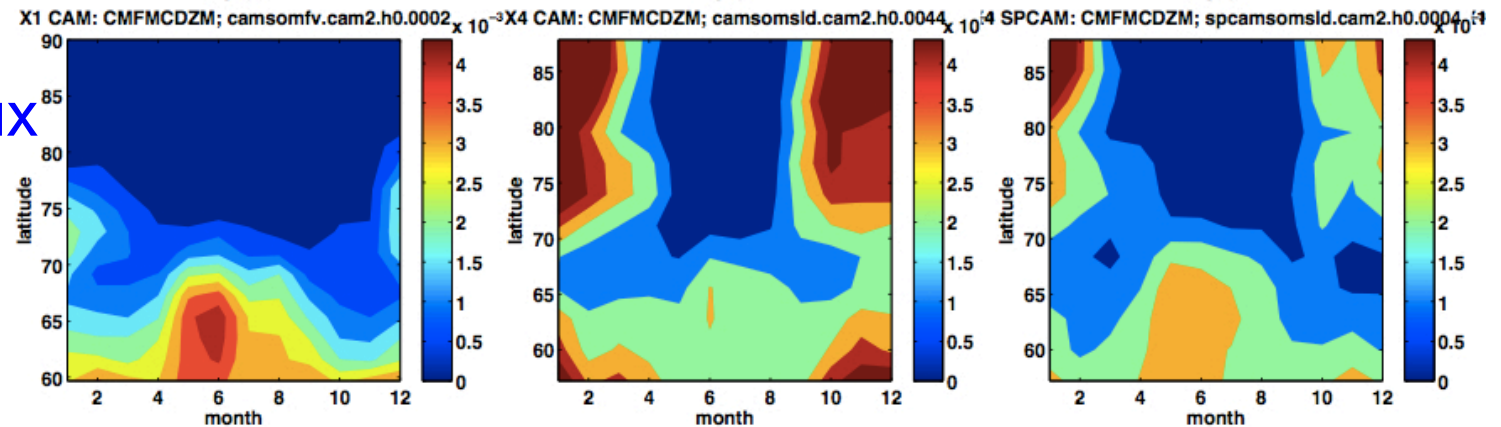
Sea ice fraction



TS



Conv Mass flux



with David
Randall, Mark
Branson

Preliminary SPCAM results: latitude (60N-90N) vs month

X1CO2,CAM; X4CO2,CAM; X1CO2, SPCAM

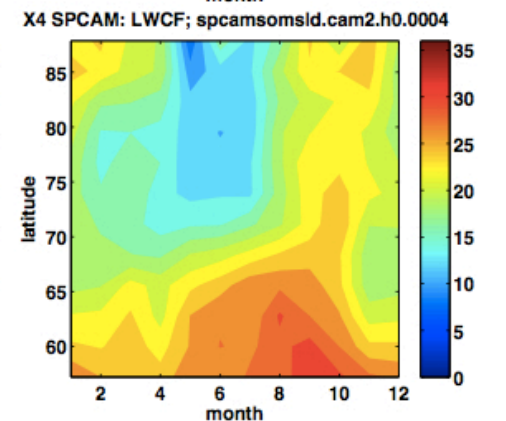
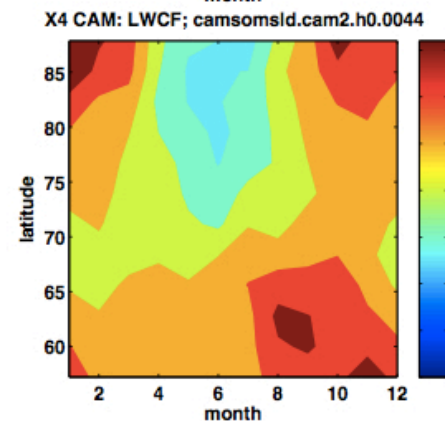
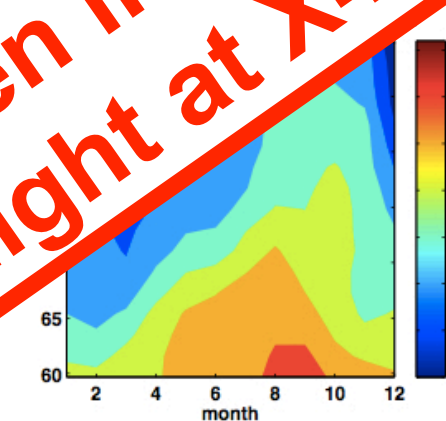
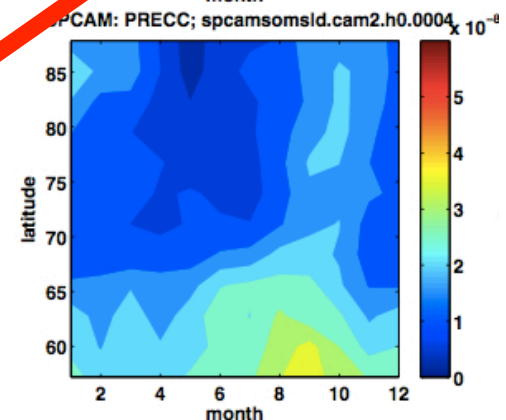
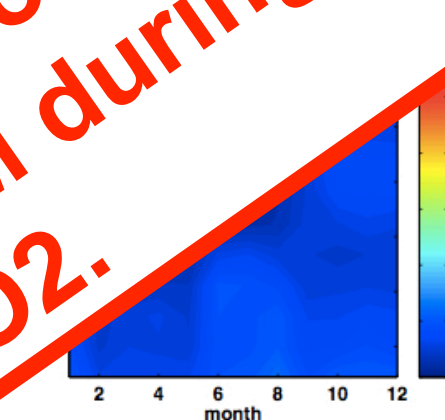
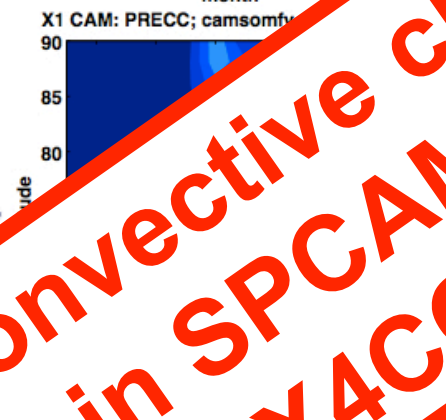
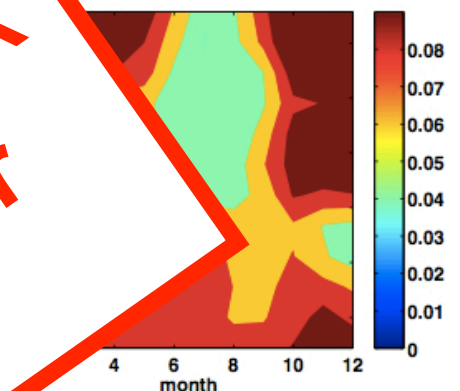
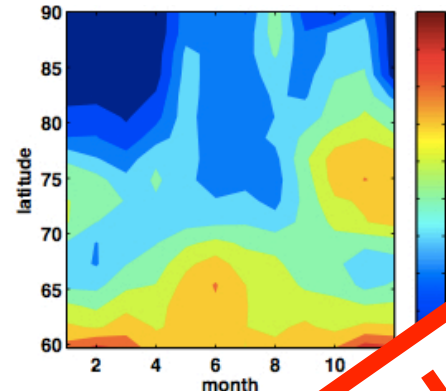
Conv clouds

Conv precip

LW CRI

with David
Randall, Mark
Branson

→ Convective cloud feedback
seen in SPCAM during polar
night at X4CO2.



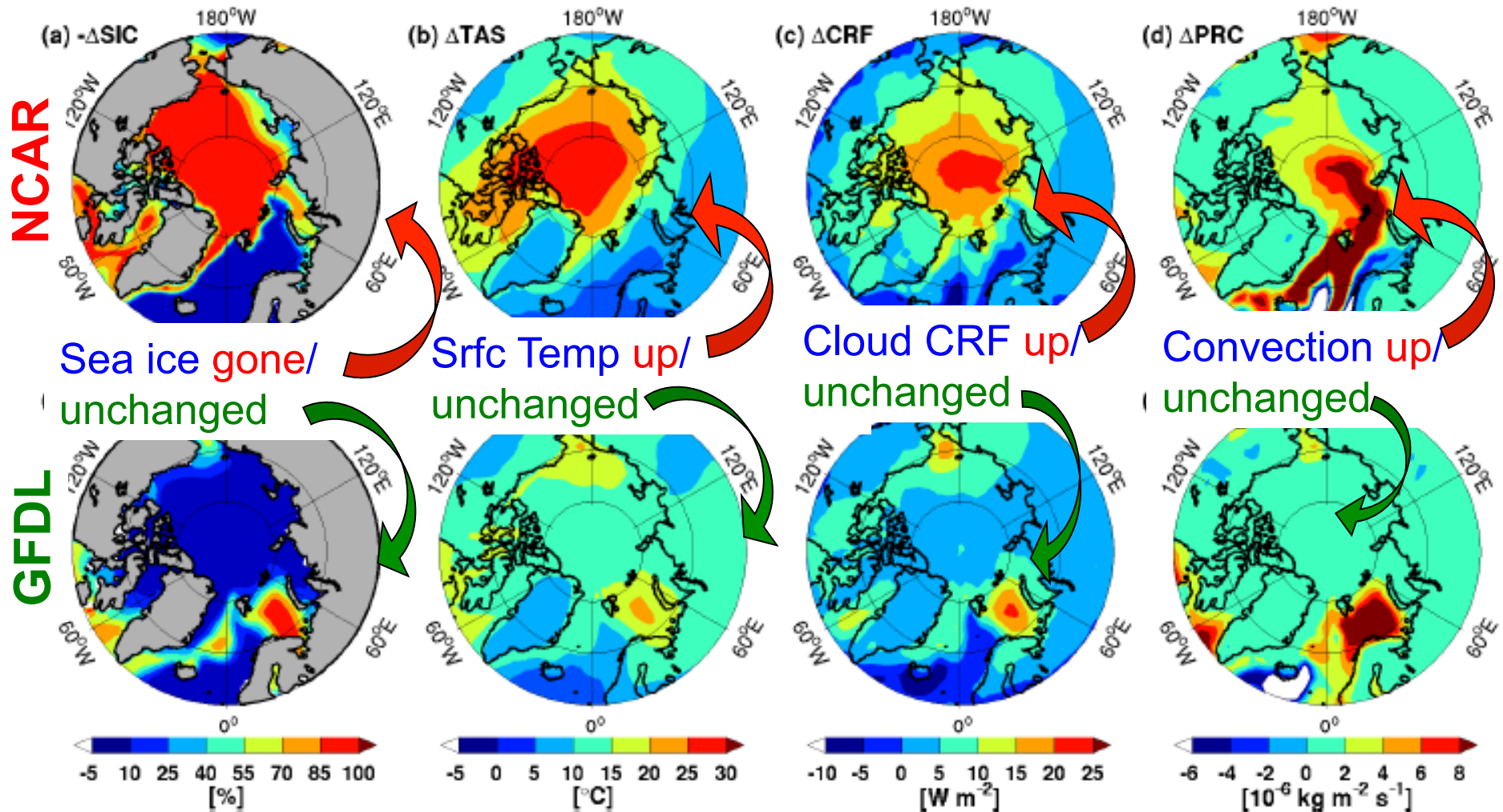
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warmth & a convective-cloud feedback

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Enticing 3D IPCC Model Simulations

Consider the solutions of **NCAR** & **GFDL** 3d coupled ocean-atm state-of-the-art models, at x4 CO₂; anomaly from pre-industrial:



➔ IPCC NCAR 3d model behaves like toy model & 1d SCAM!!

Conclusions: Eocene (50 Myr) warmth & convective-cloud feedback

Challenge: CO₂ insufficient to explaining Eocene warmth

Good news: Found a simple, interesting & unexpected climate state at high CO₂: high-latitude deep atmc convection & high tropospheric clouds result in an equable-like climate

- Solution is self-consistent, clouds and convection reinforce each other and don't need to be specified arbitrarily, confirmed in full complexity state-of-the-art atmospheric and climate models.
- Future? Arctic uncertainty may mean *more* warming... 14 IPCC models: feedback strength uncorrelated w/ climate sensitivity!

[Abbot & Tziperman, 2008,9: QJRMS, GRL, JAS, J. Climate]

Pliocene (1.8-5.3 Myr) “*permanent El Nino*” & atmospheric superrotation

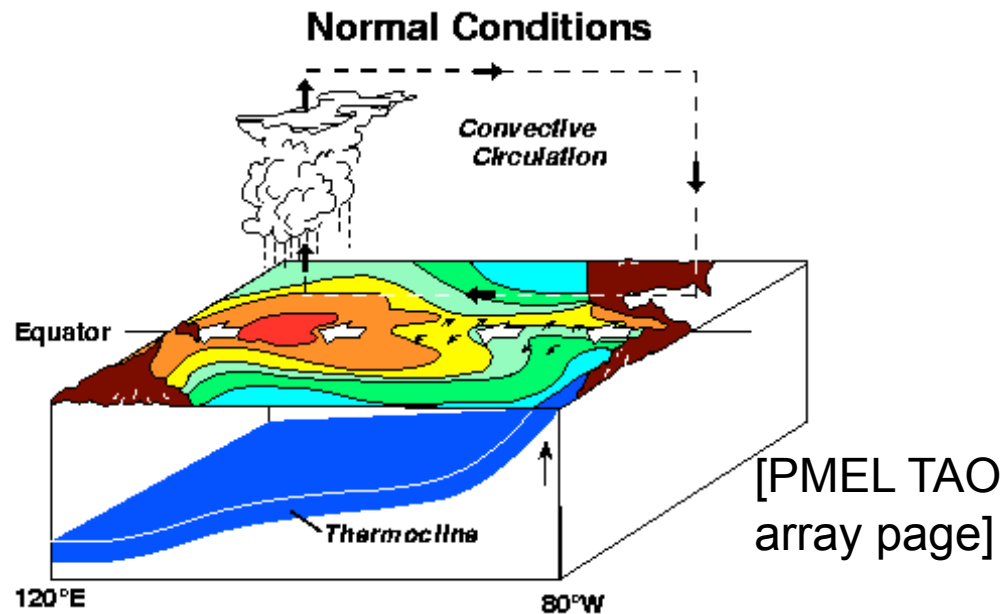
with: **Nathan Arnold, Brian Farrell**

Outline

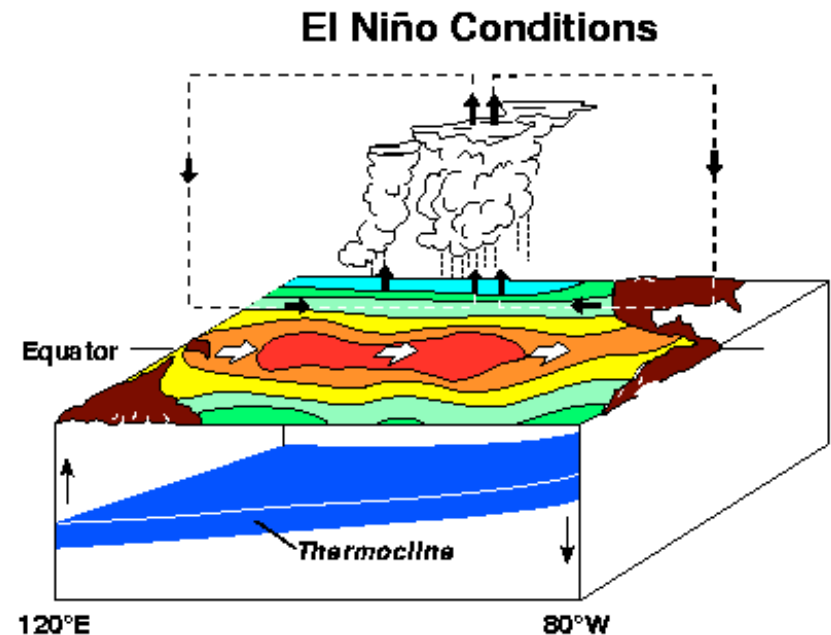
- Pliocene, “permanent El Nino”
- Some previous attempts
- Superrotation! weakening those easterlies
 - Rossby wave resonance mechanism
 - Increased convective activity with SST in SPCAM

LESSON 2

Reminder: El Nino, the equatorial Pacific easterlies & thermocline



Normal conditions: easterly winds push warm water to west, East Pacific cools bec. of shallow thermocline there.



El Niño: easterlies weaken, East Pacific thermocline deepens, EP surface ocean warms

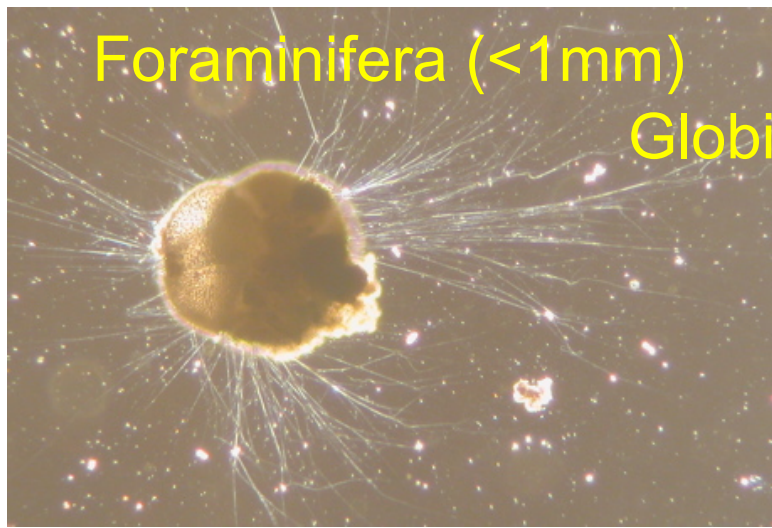
➔ Two possible ways of 'making' a permanent El Niño:
(1) weaken easterlies or (2) deepen thermocline.

The Pliocene (2-5 Myrs)

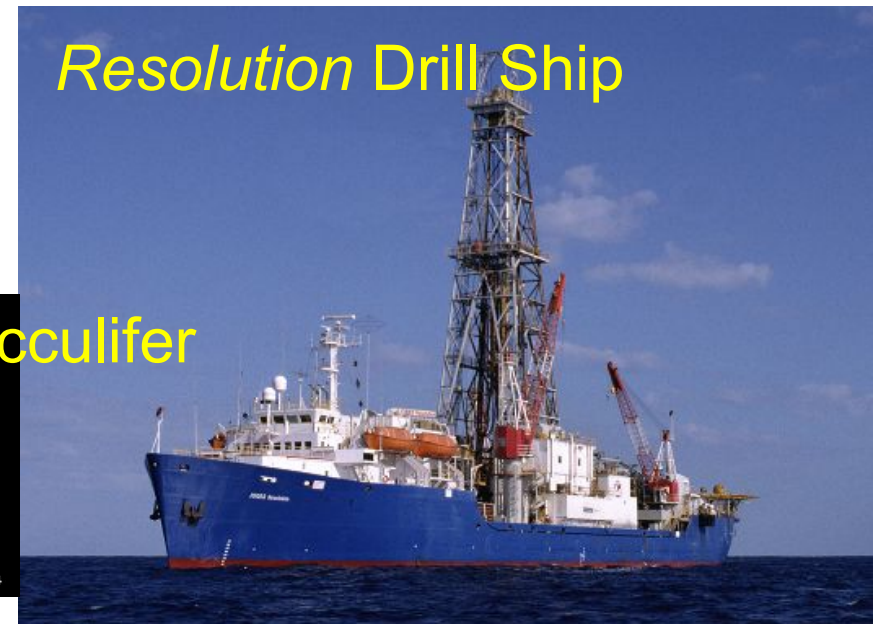
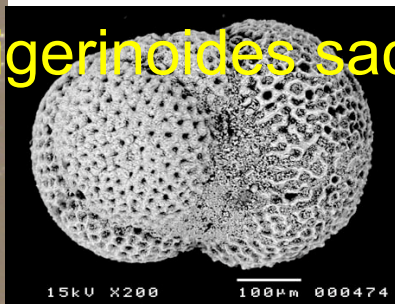
- CO₂ 350-500ppm? (Today: 380; preindustrial: 280; in 50 yrs ...)
- Global average surface temperature: $\approx 3^\circ$ warmer than today
- Ice: covers Antarctica, but not much in northern hemisphere (ice ages started ≈ 2.7 Myrs ago)

How do we know:

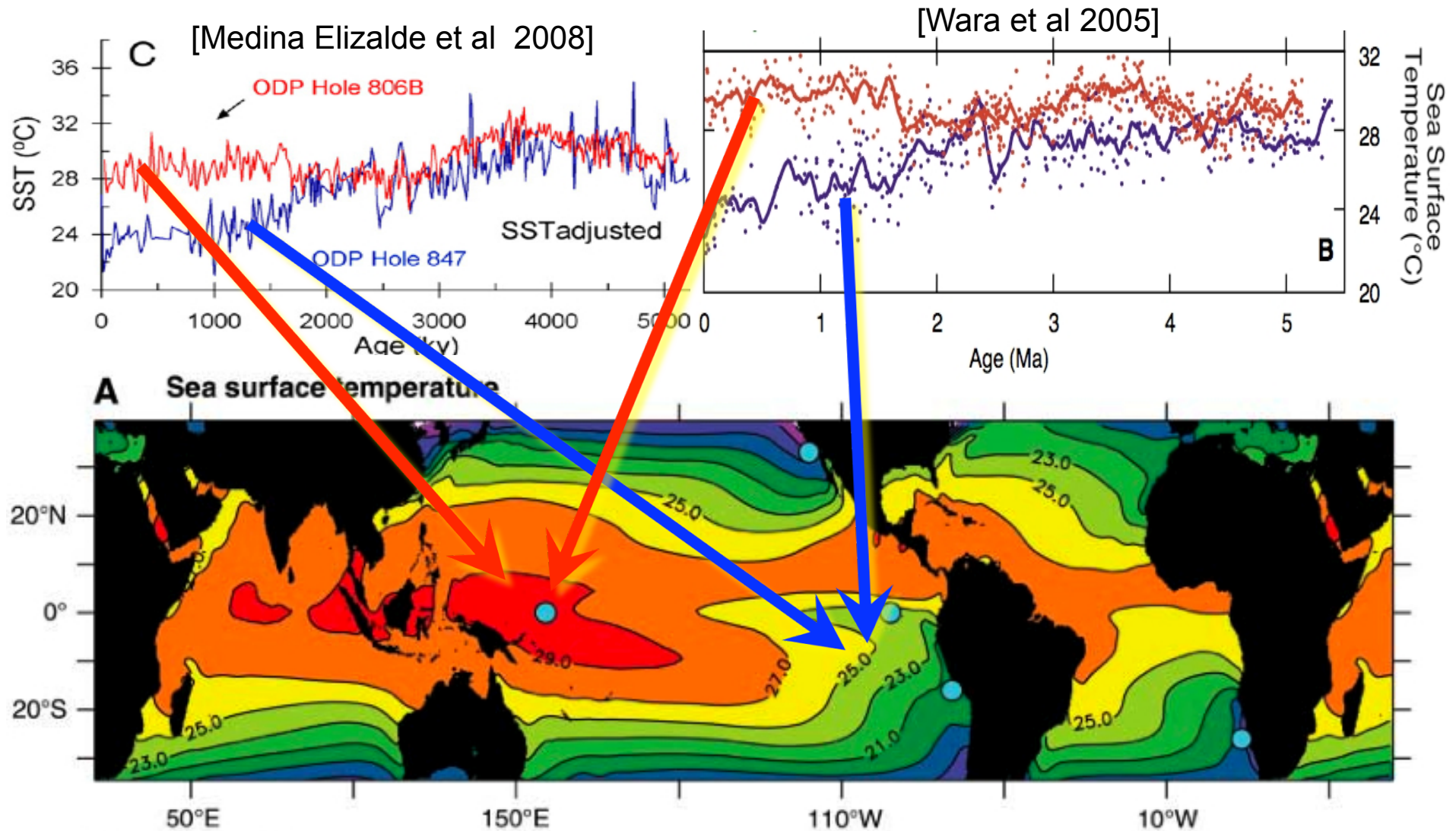
- Isotopic/ other proxy records from deep sea drilling.



Globigerinoides sacculifer



The equatorial Pacific during the Pliocene



➔ The east-west temperature gradient in the equatorial Pacific did not exist during the Pliocene (2-5 Myr ago)

Previous explanations of the Pliocene “*permanent El Nino*”

- Deeper global thermocline? Evidence: strong warming in upwelling sites off Africa, California, South America
- Mechanism [Fedorov et al]: a collapse of the thermocline by a very strong fresh water forcing [in the north Pacific?]

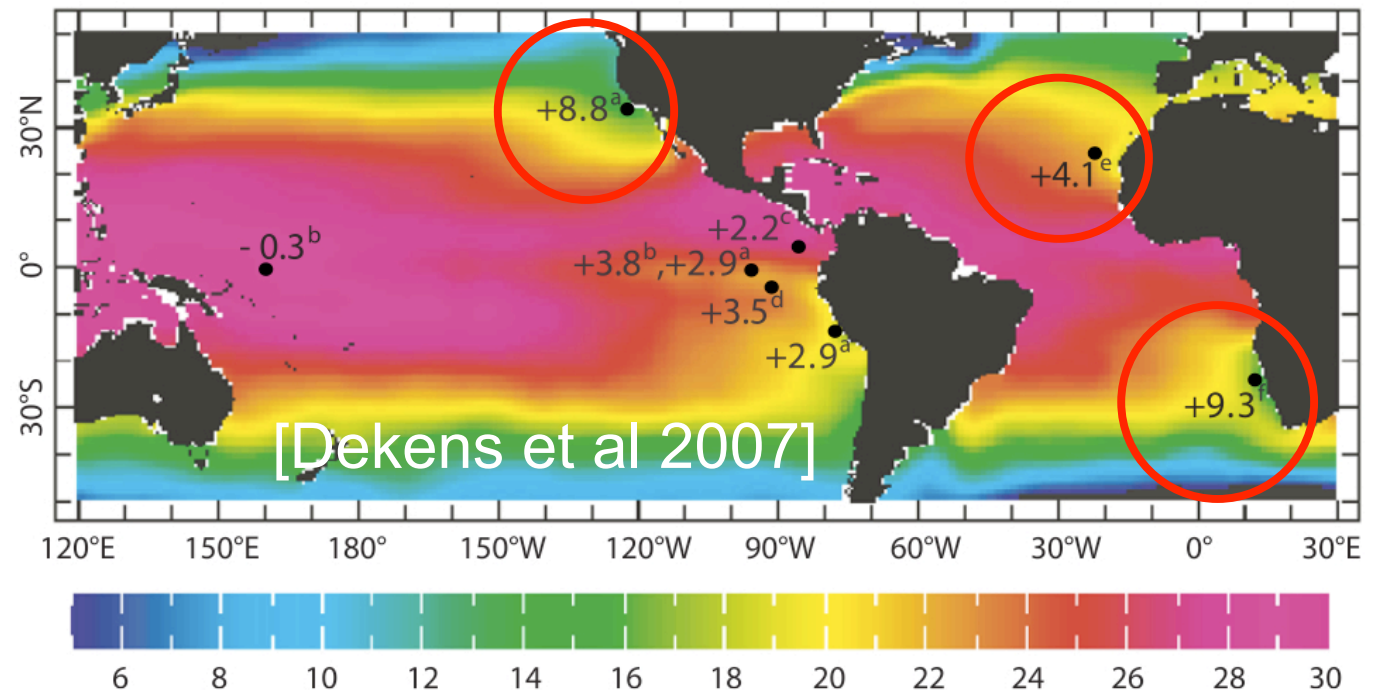
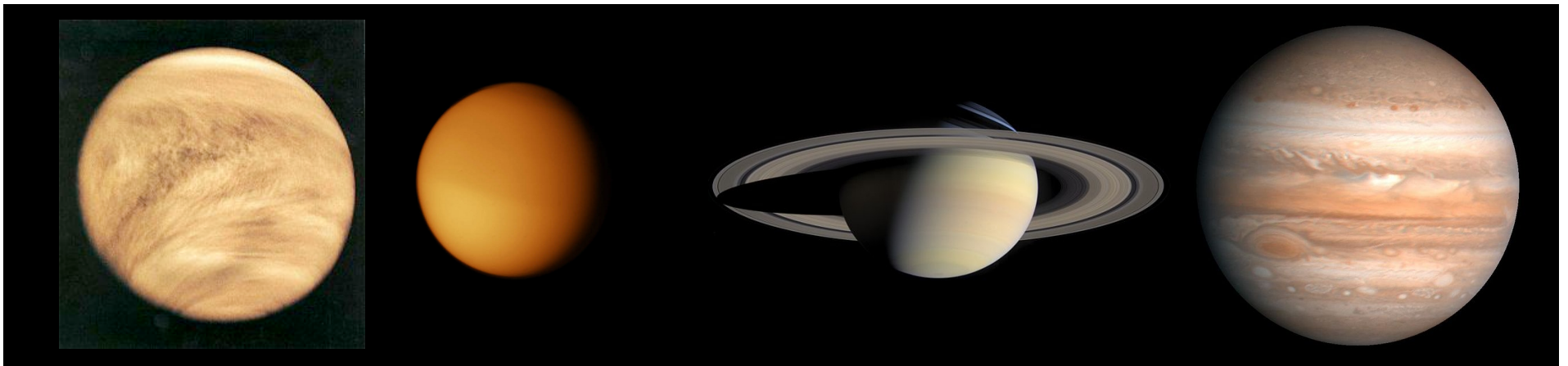


Figure 1. Difference in sea surface temperature (SST) between Pliocene and modern SST. The colored map shows modern mean annual SST [Levitus and Boyer, 1994]. Superimposed is the difference between

Superrotation

- Superrotation = Zonally-averaged westerly wind at the equator, basically the atmosphere rotating faster than earth itself
- Seen in the atmospheres of Venus, Titan, Saturn, and Jupiter:



- Also seen in the upper atmosphere during MJO
- **Forbidden** by angular momentum conservation in the absence of up-gradient angular momentum fluxes (Hide's theorem) → must involve some non-trivial eddy dynamics.

Superrotation dynamics: Rossby Wave reminder...

Consider a wave solution $\psi = A \cos(kx + ly - \sigma t)$

Rossby wave dispersion relation $\sigma = \frac{-\beta k}{k^2 + l^2 + L_R^{-2}}$

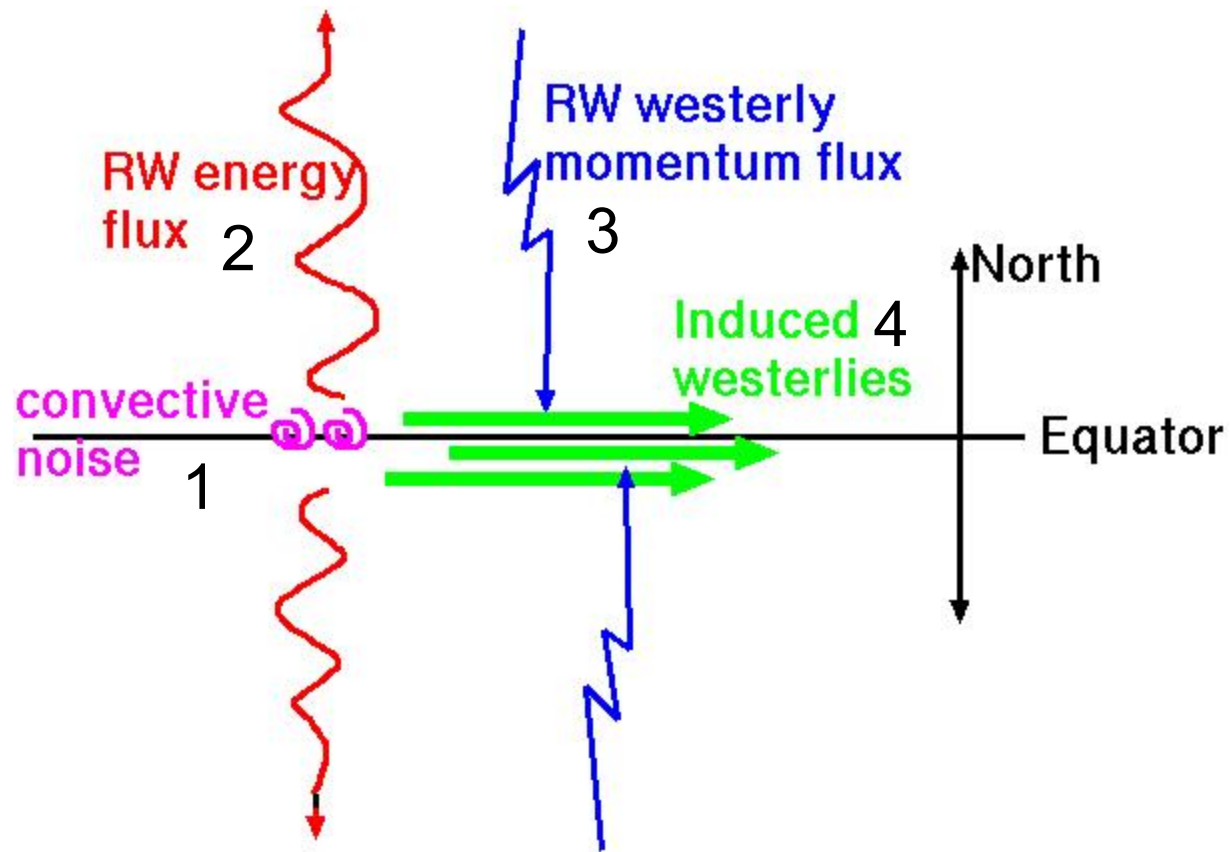
Meridional group velocity $c_g^{(y)} = \frac{2\beta kl}{(k^2 + l^2 + L_R^{-2})^2}$

Meridional momentum flux

$$\overline{u'v'} = \overline{(-\psi_y)(\psi_x)} = -kl A^2 \overline{\sin^2(kx + ly - \sigma t)}.$$

➔ Meridional momentum flux is in opposite direction to group velocity. Specifically, energy flux away from equator implies momentum flux toward equator westerly momentum induced at equator.

Superrotation dynamics: Rossby Wave reminder...



A partial superrotation literature review

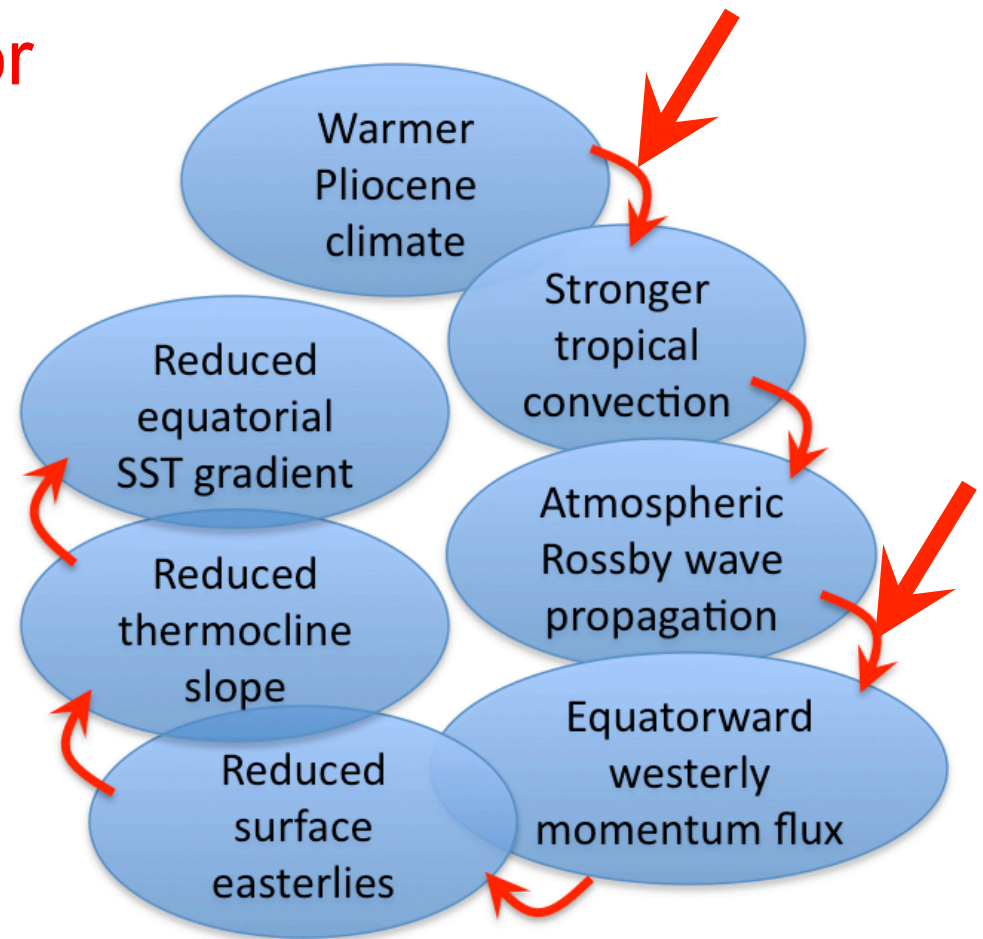
- 2-level PE models search for multiple equilibria due to eddy fluxes **from mid-latitudes**: [Suarez and Duffy, 1992; Saravanan, 1993].
- and later also 3d GCMs: [Williams , 2006, 2003]
- Theoretical considerations of wave propagation [Panetta et al., 1987]
- Superrotation multi-equilibria due to a feedback of mean circulation not involving momentum wave flux [Shell & Held , 2004]
- 18 level AGCM: Steady longitudinal variations in diabatic heating → horiz eddy momentum fluxes stationary planetary waves → superrotation [Kraucunas & Hartmann 05]
- Moving flame effect (Lindzen's book, Venus)
- **Possible superrotation & the collapse of the walker circulation in a future global warming scenario** [Held, 1999; Pierrehumbert 2002]
- Pierrehumbert [2002] writes:
“There is no evidence that a westerly superrotating state has ever occurred in any climate of the Earth's past...”
And this is where it gets interesting...

Proposed mechanism for permanent El Nino

1. Warmer Pliocene → stronger MJO-like tropical convection =stochastic forcing at equator.

2. Rossby wave energy flux away from equator → equatorward westerly momentum flux → weaken equatorial easterlies.

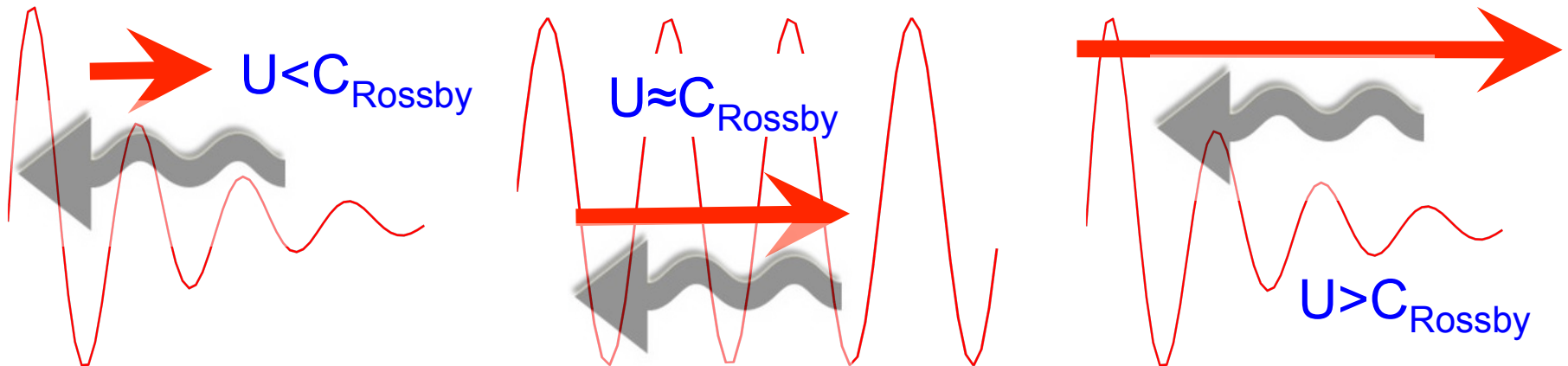
3. Weaker easterlies → decreased E-W thermocline slope → eliminate East Pacific cold tongue & E-W SST gradient → Permanent El Nino!



Exploring 2 critical elements of superrotation mechanism

- 1) From convective noise to Rossby wave propagation via a wave-mean flow resonance mechanism
- 2) Warm climate \rightarrow stronger convective activity

Forced Rossby waves are evanescent unless mean flow speed is equal and opposite to free Rossby wave phase speed.

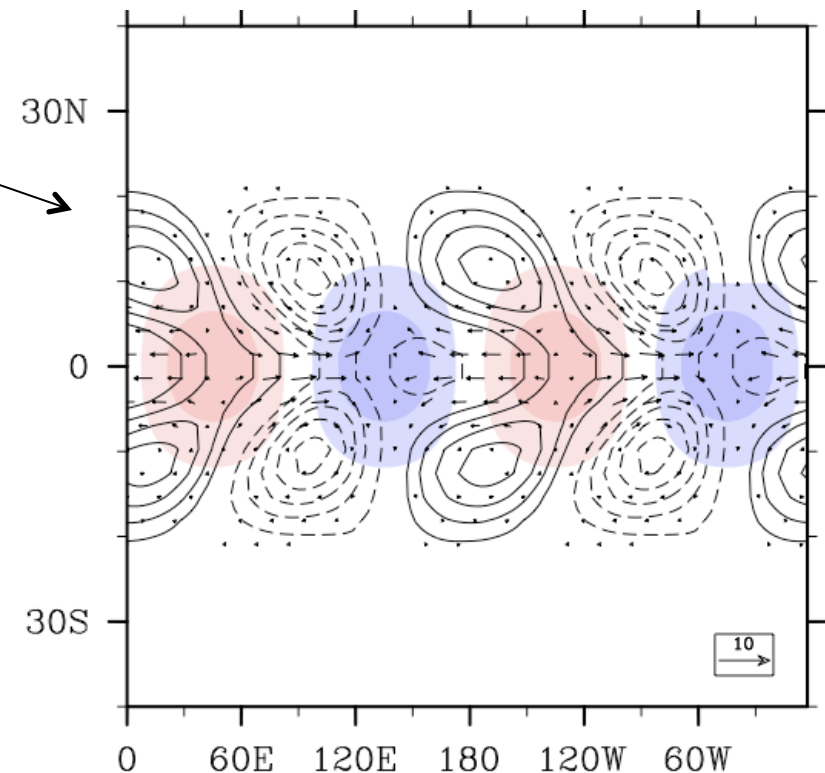
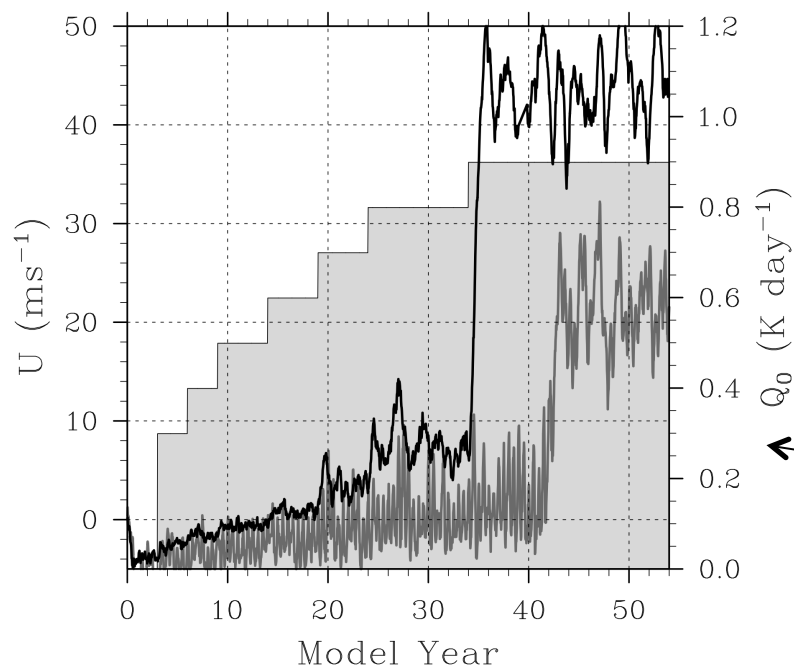


positive feedback: westerly wind strengthens \rightarrow approaches phase speed of free Rossby wave \rightarrow waves amplify \rightarrow stronger equatorward momentum flux \rightarrow enhanced westerlies
 \rightarrow A resonance! maximum wave amplification & westerly acceleration occur when westerly speed = Rossby wave speed.

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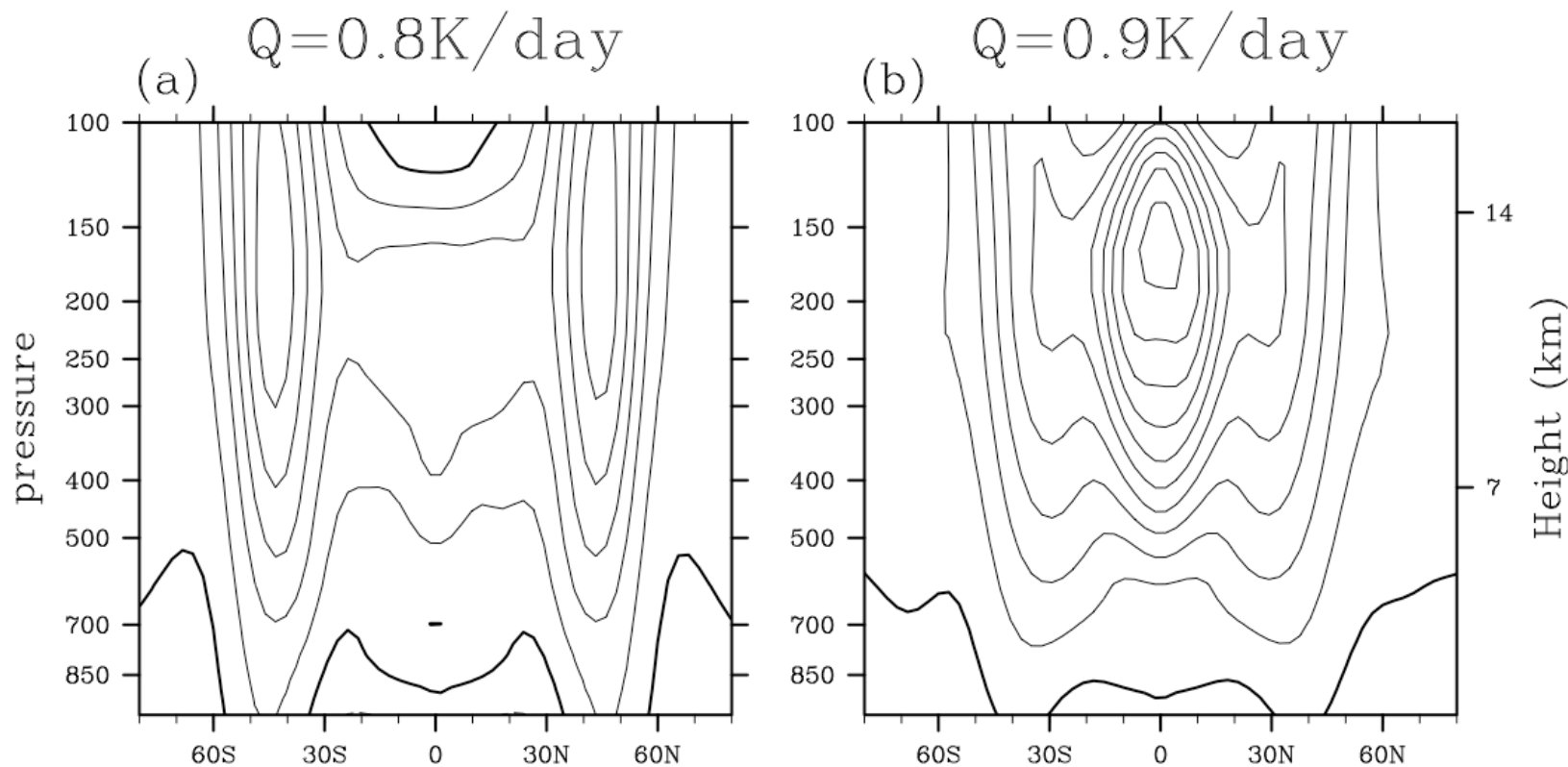
CAM experiment: (1) Impose eddy heating at equator (colors); (2) gradually increase heating rate...



Results: abrupt transition to superrotation

Exploring 2 critical elements of superrotation mechanism

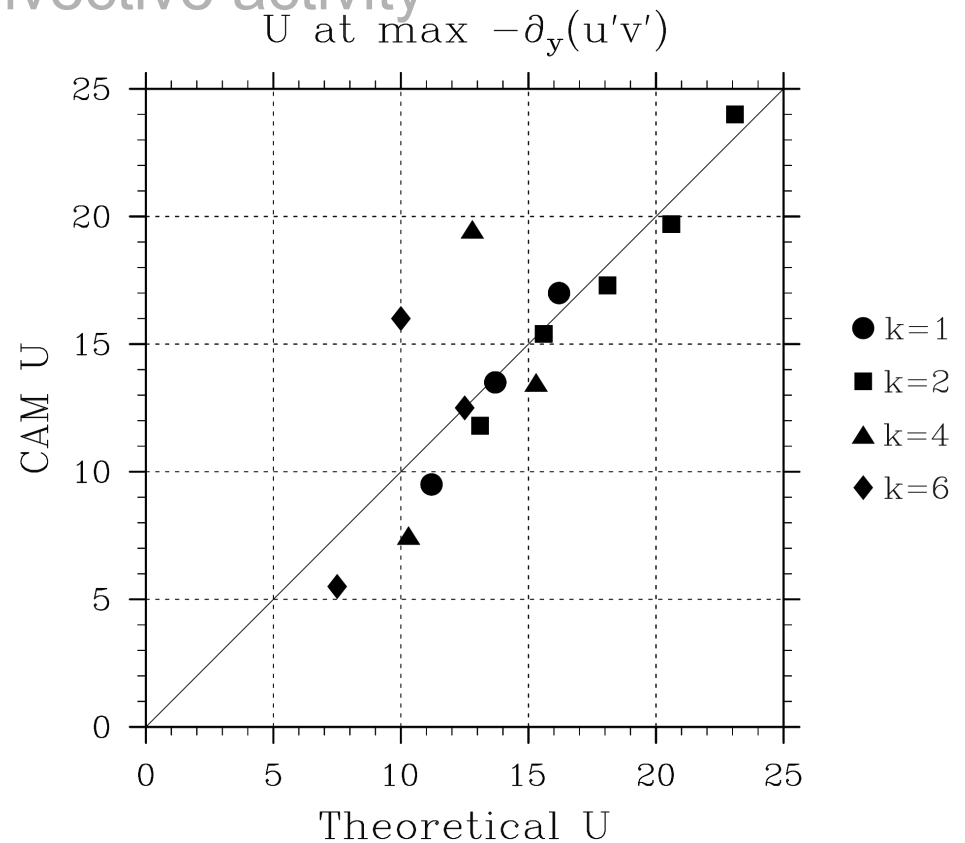
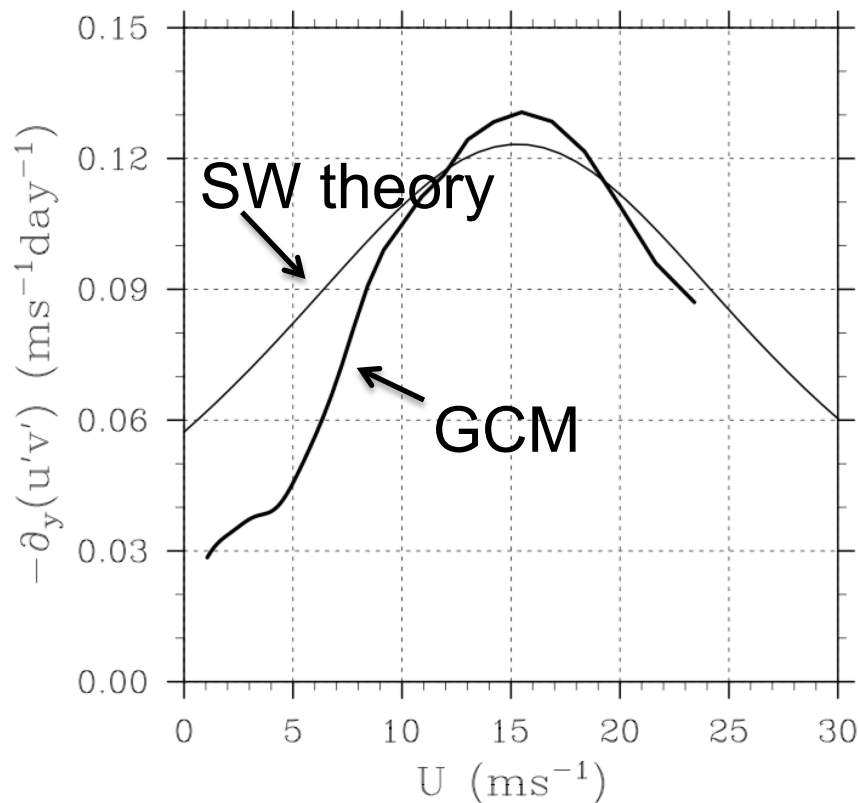
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Zonal Wind before and after bifurcation,
showing transition to a strong superrotation

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Comparison with shallow water analytical solution confirms resonance; experiments specify k & propagation speed of heating.

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Some previous evidence:

Observations: [Slingo et al. 1999]: increased MJO activity since 1970s? due to decadal tropical SSTs warming?

Idealized AGCM: [Lee, 1999] eddy flux convergence due to “MJO” twice as strong due to a uniform 3 degree warming

SPCAM results:

Compare aquaplanet with imposed zonally-averaged SST, with a uniformly increased SST run.

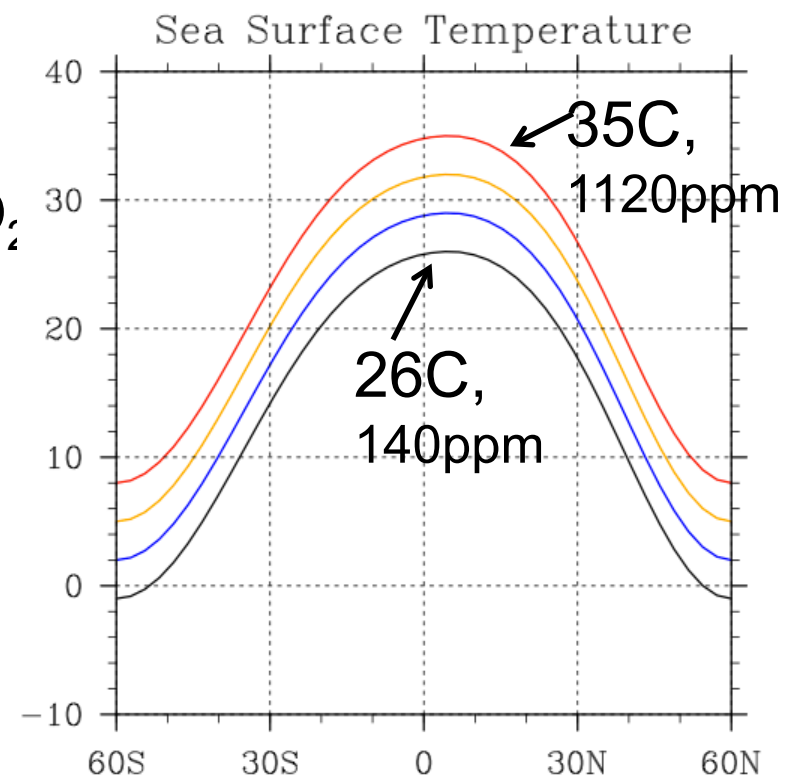
Exploring 2 critical elements of superrotation mechanism

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Effect of high SST in aquaplanet SPCAM3.5

Experimental Setup:

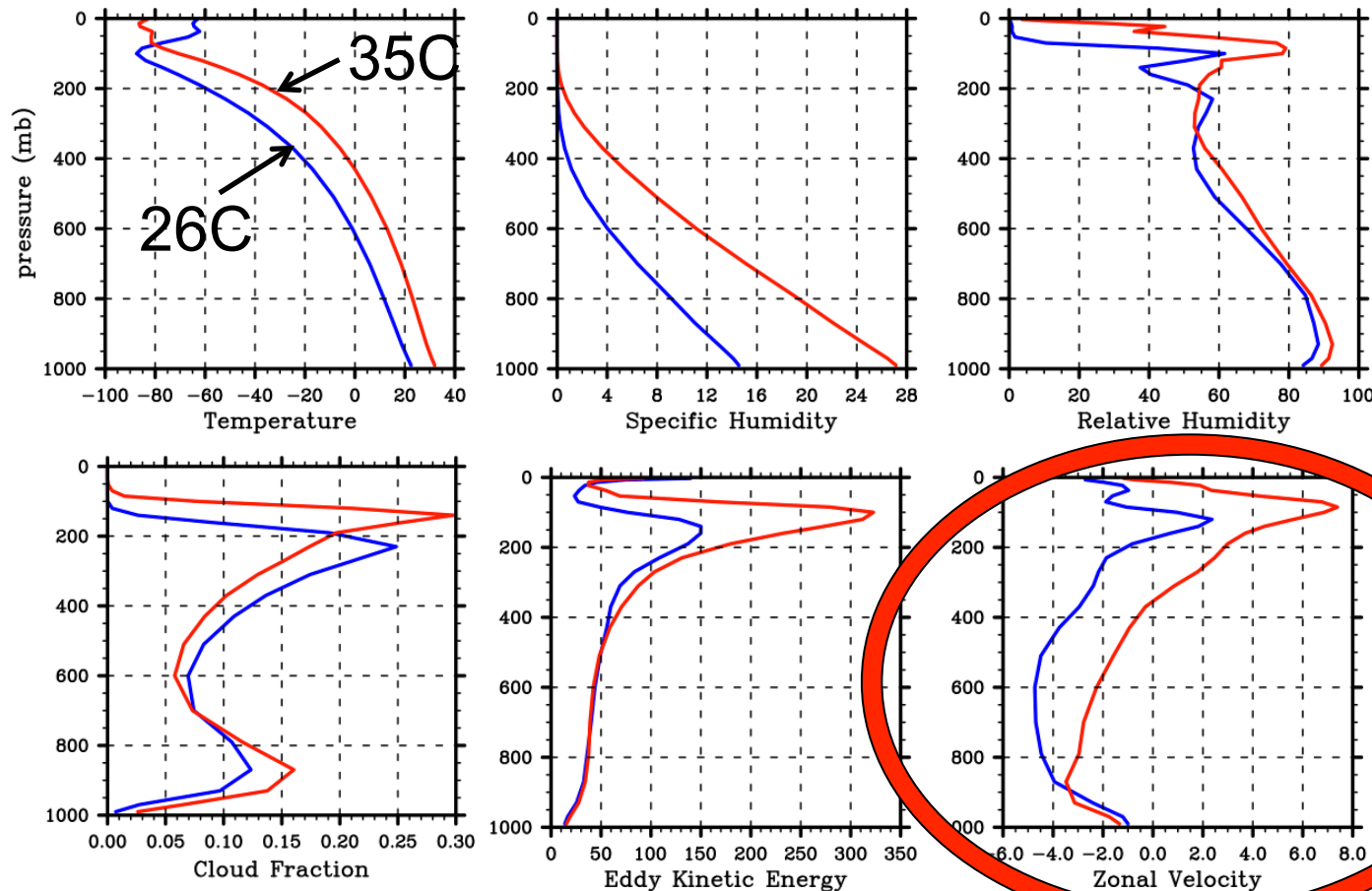
- Prescribed zonally symmetric SST, constant in time.
- Uniformly increased by 3C with CO₂ doublings
- SST peak offset to 5N,
 - creates “ITCZ”
 - cross-equator flow opposes superrotation
- No sea ice



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Effect of high SST in aquaplanet SPCAM3.5

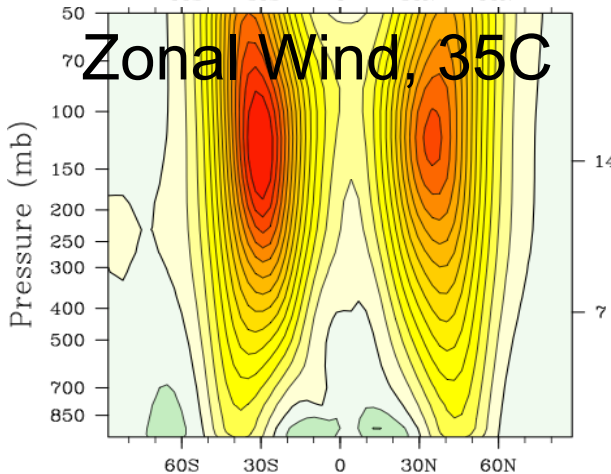
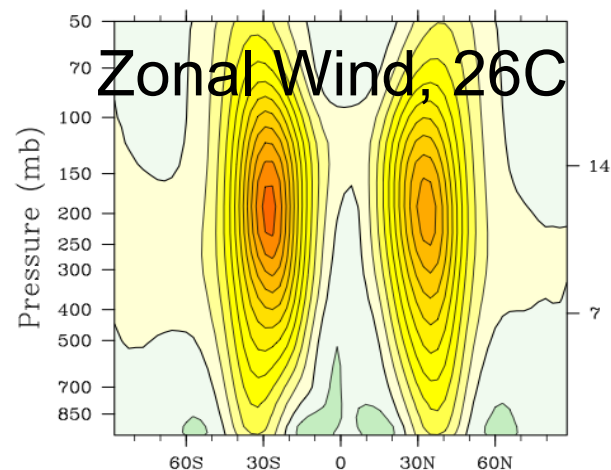


Encouraging,
but no *surface*
superrotation

Exploring 2 critical elements of superrotation mechanism

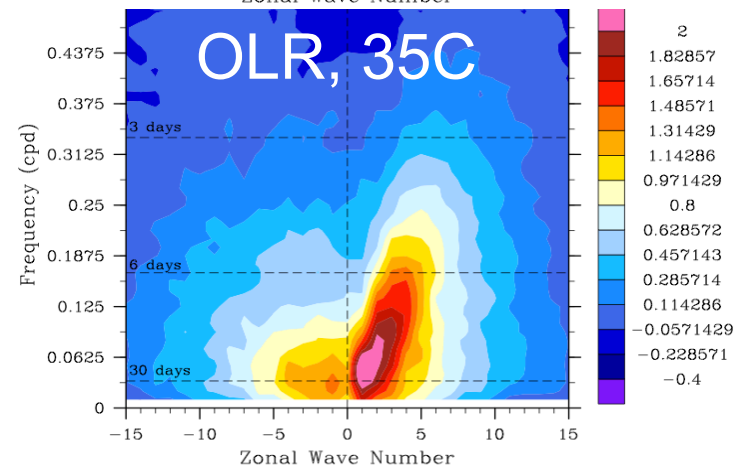
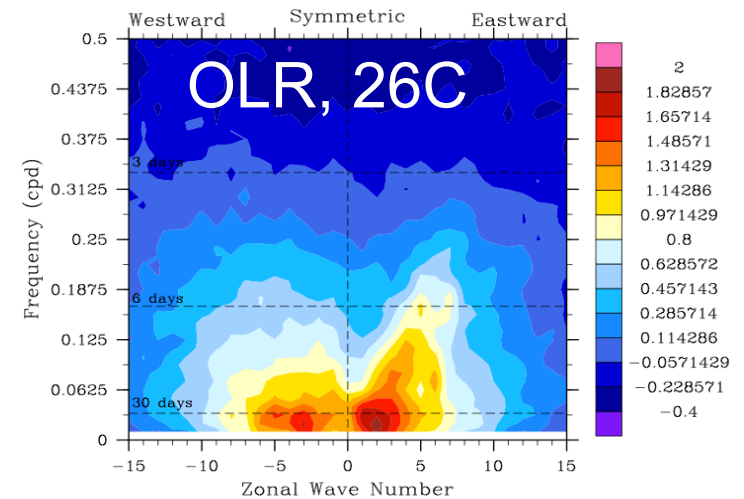
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Effect of high SST in aquaplanet SPCAM3.5



Increased MJO-like activity with SST!

Mechanism:
positive feedback due to steepening of mean MSE profile in warmer climate [Arnold, Kuang, Tziperman, in press]



Conclusions: permanent El Nino & superrotation

- We tried to make the case for superrotation as a mechanism for the vanishing equatorial Pacific SST gradient 3-5 Myr ago.
- Mechanism: enhanced convective “noise” at equator, radiating Rossby waves and inducing westerlies at equator.
- We proposed a Rossby wave resonance mechanism, which tends to lead to *abrupt transition* (bifurcation) to superrotation
- We find evidence in SPCAM, of enhanced convective activity & tendency to superrotation at high altitudes for warmer SST.
- Major challenge now: getting superrotation to surface... CMT?
- Could this mechanism lead to a permanent El Nino in the near future? **[as suggested by Held 1999 & Pierrehumbert 2002]**

[Tziperman & Farrell 2009; Arnold, Tziperman & Farrell 2011; Arnold, Kuang, Tziperman, 2012]