

Clouds

Global Warming Science, EPS101

Minmin Fu and Eli Tziperman

<https://cloudappreciationsociety.org>



<https://courses.seas.harvard.edu/climate/eli/Courses/EPS101/>

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cumulus



altocumulus



stratocumulus



cumulonimbus



cirrus



cirrocumulus



stratus

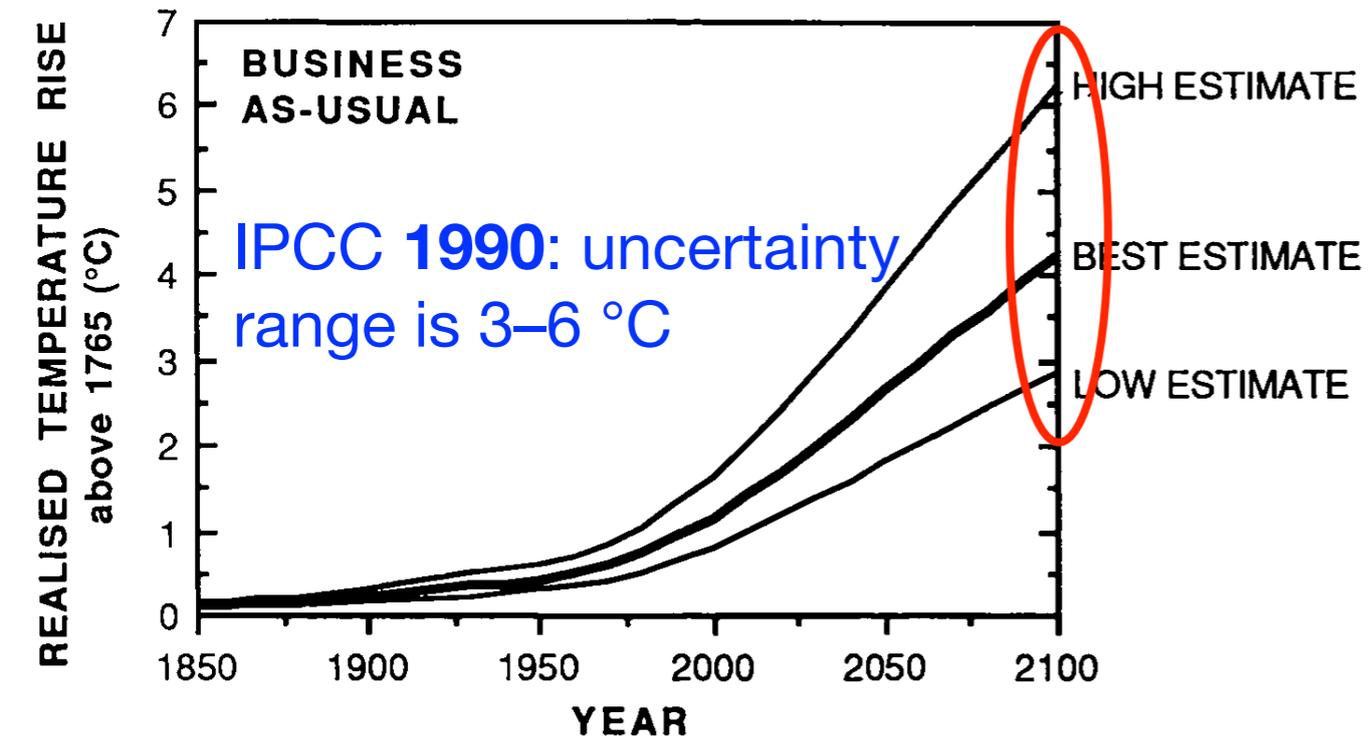


altostratus

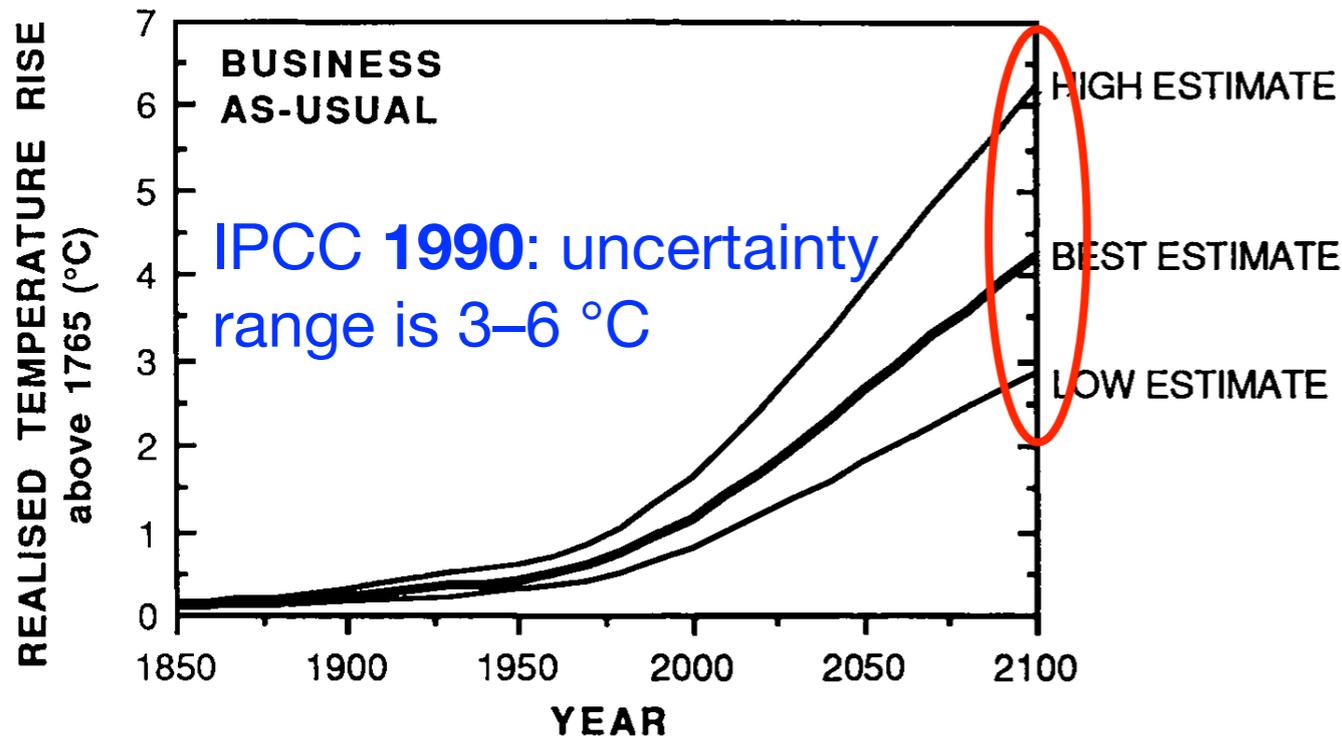


cirrostratus

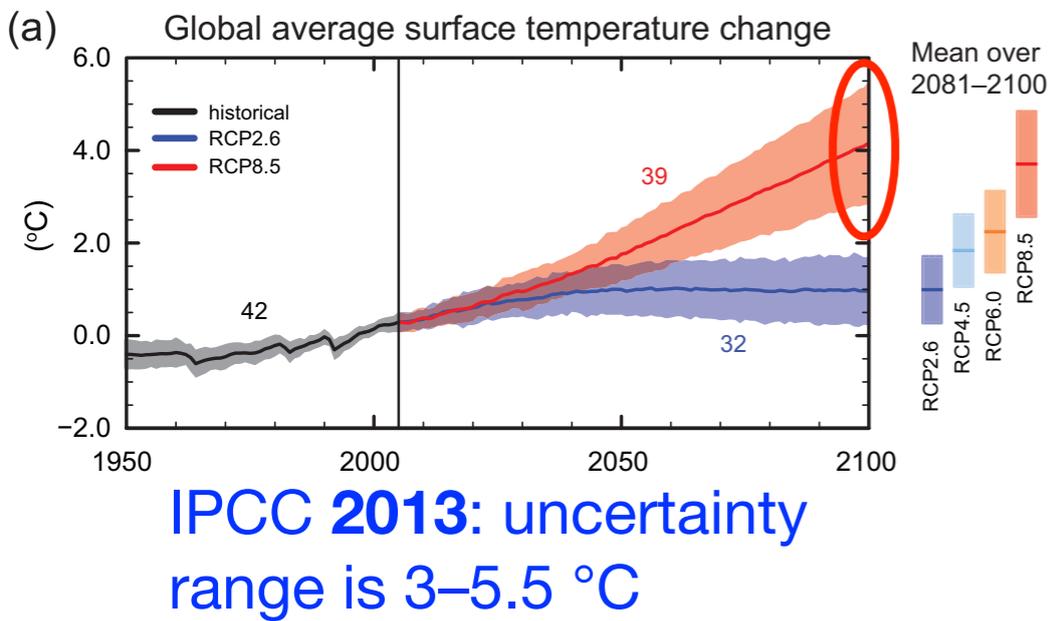
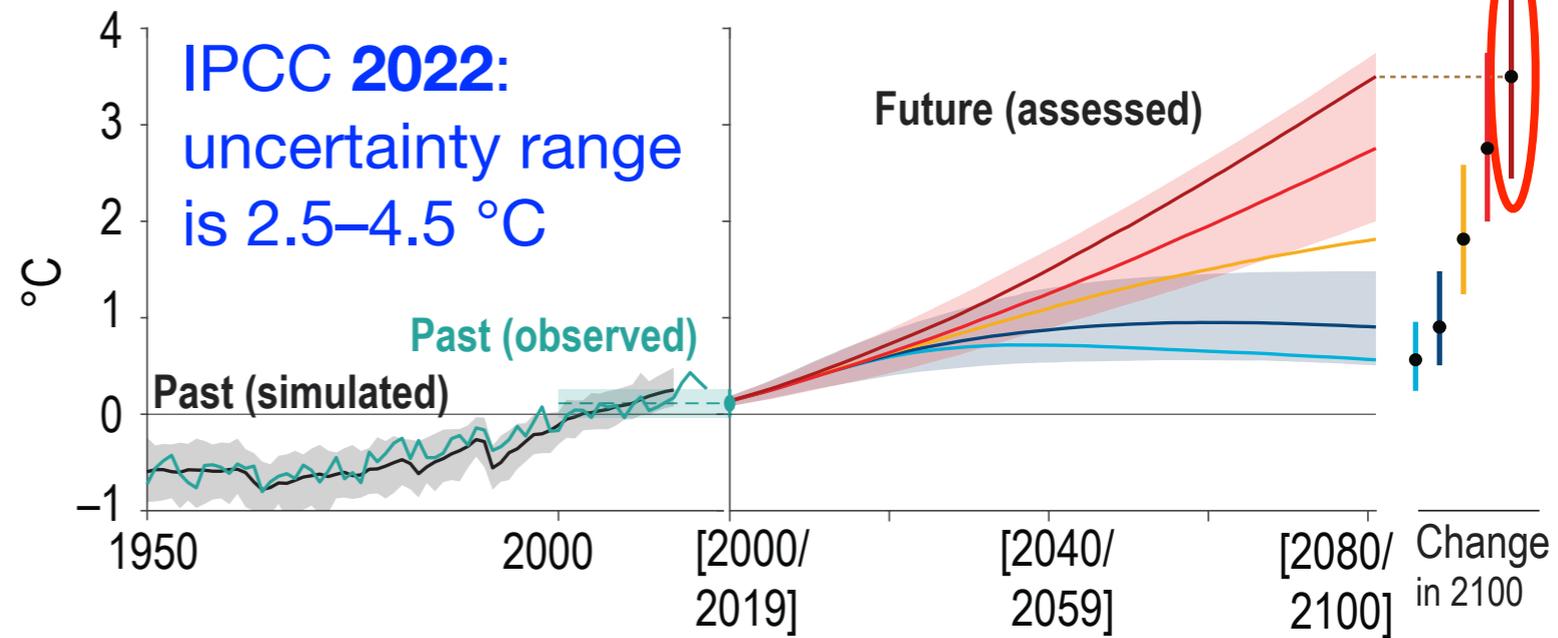
Climate prediction uncertainty



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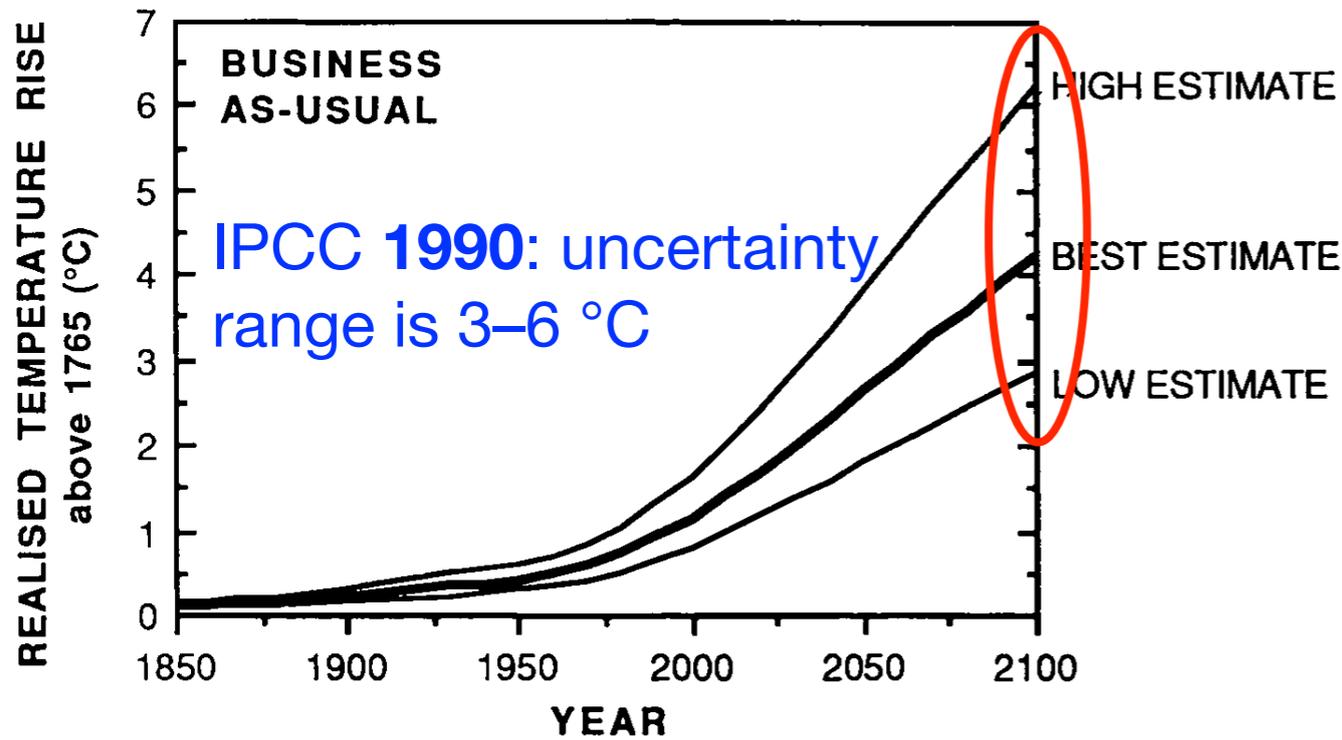


(a) Global surface air temperature Figure TS.8

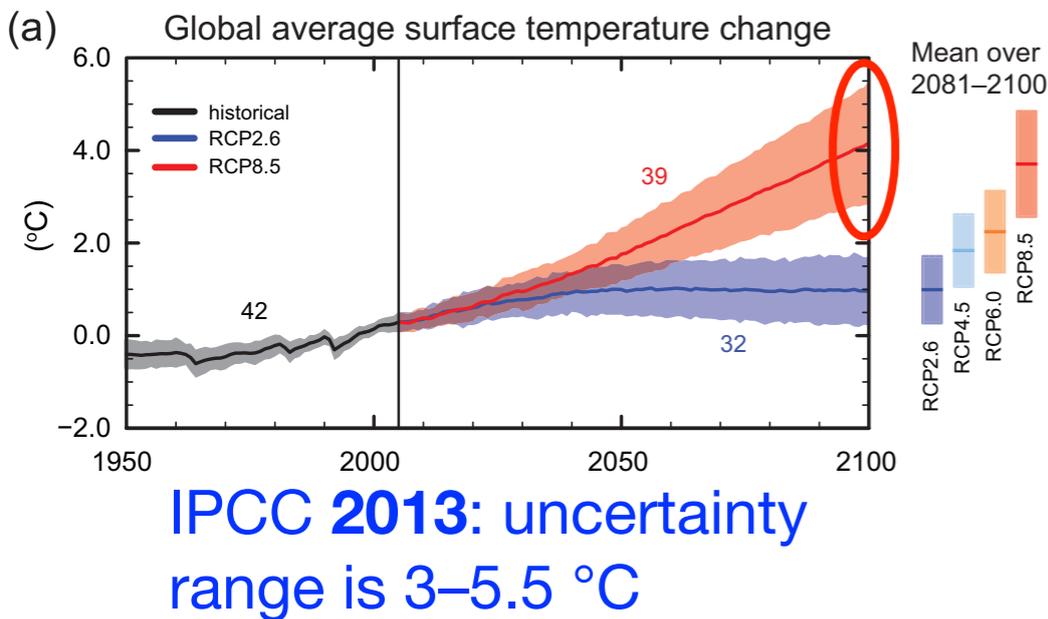
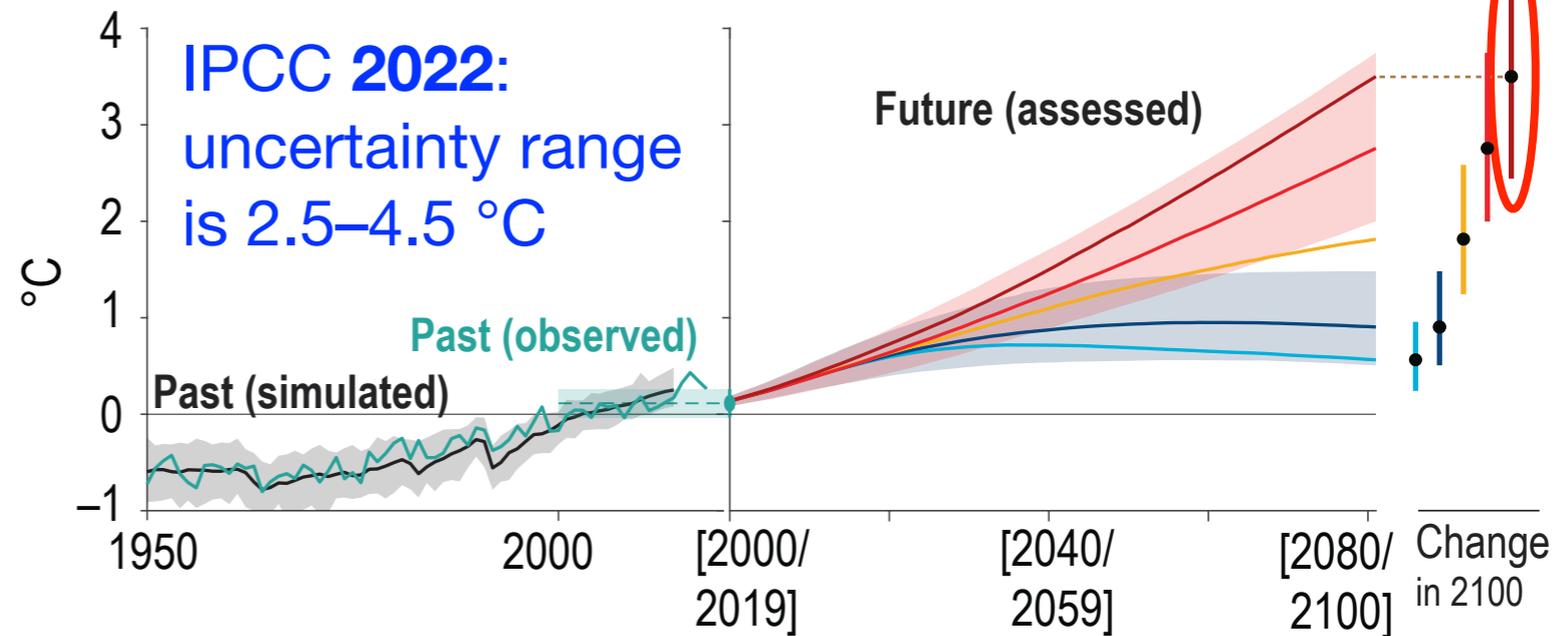


Climate prediction uncertainty

Uncertainty in the projected response to business-as-usual emission scenarios remains almost unchanged for 35 years

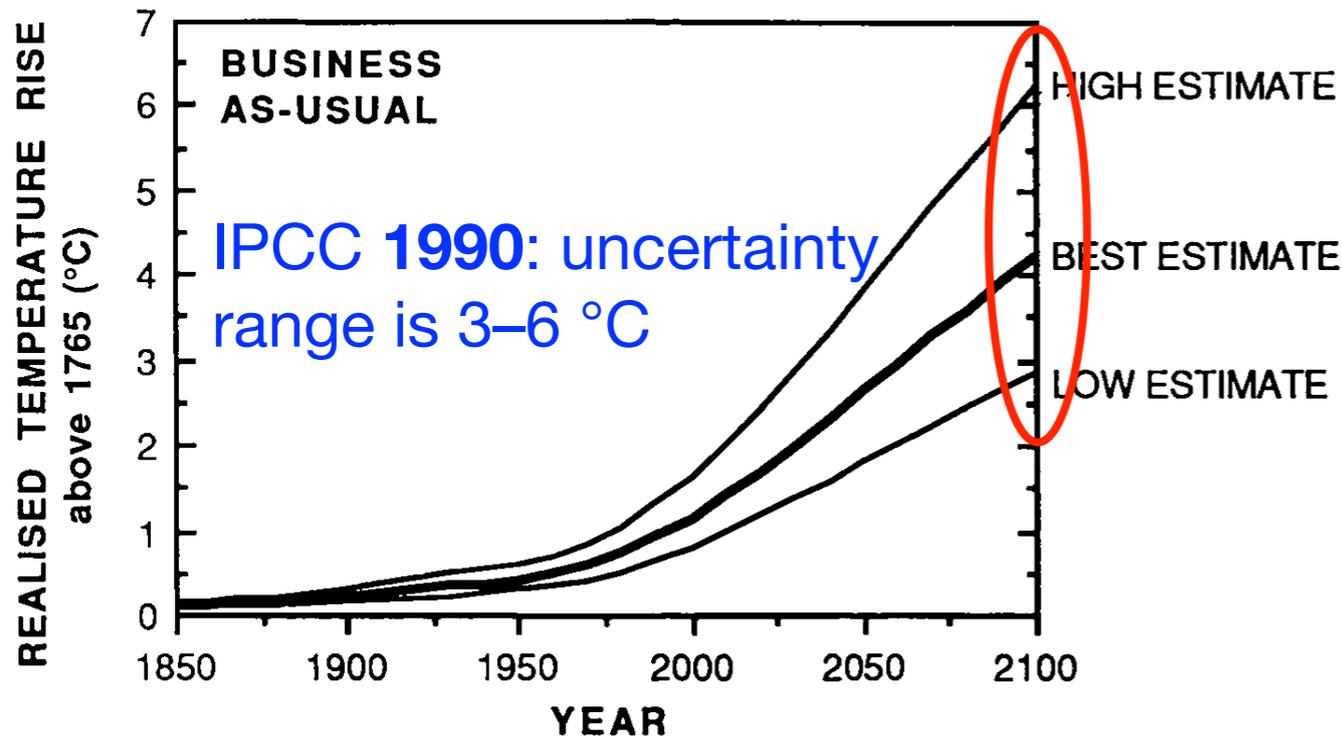


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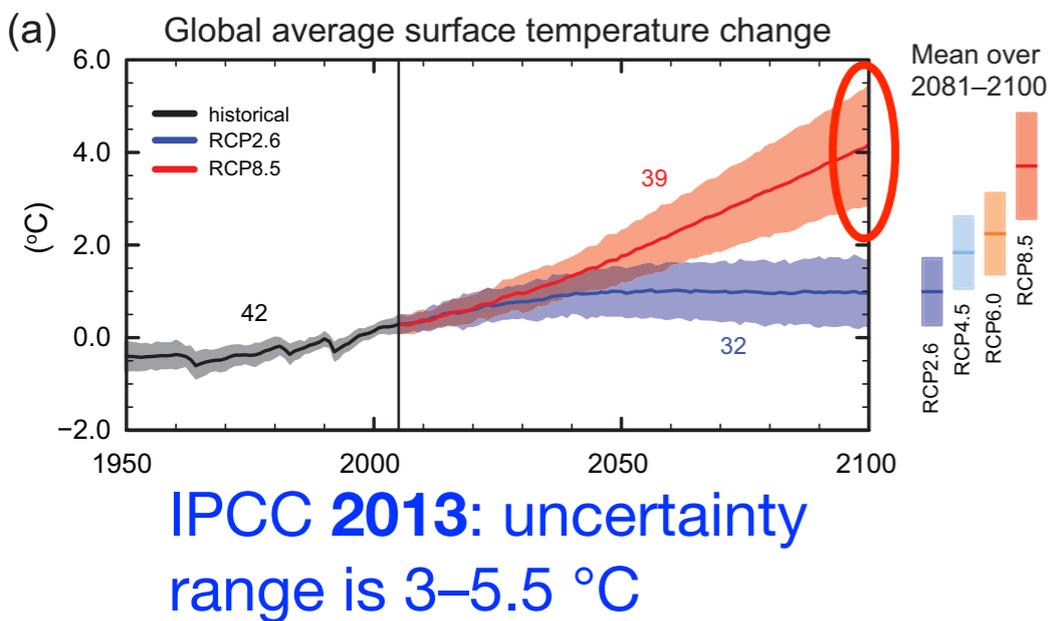
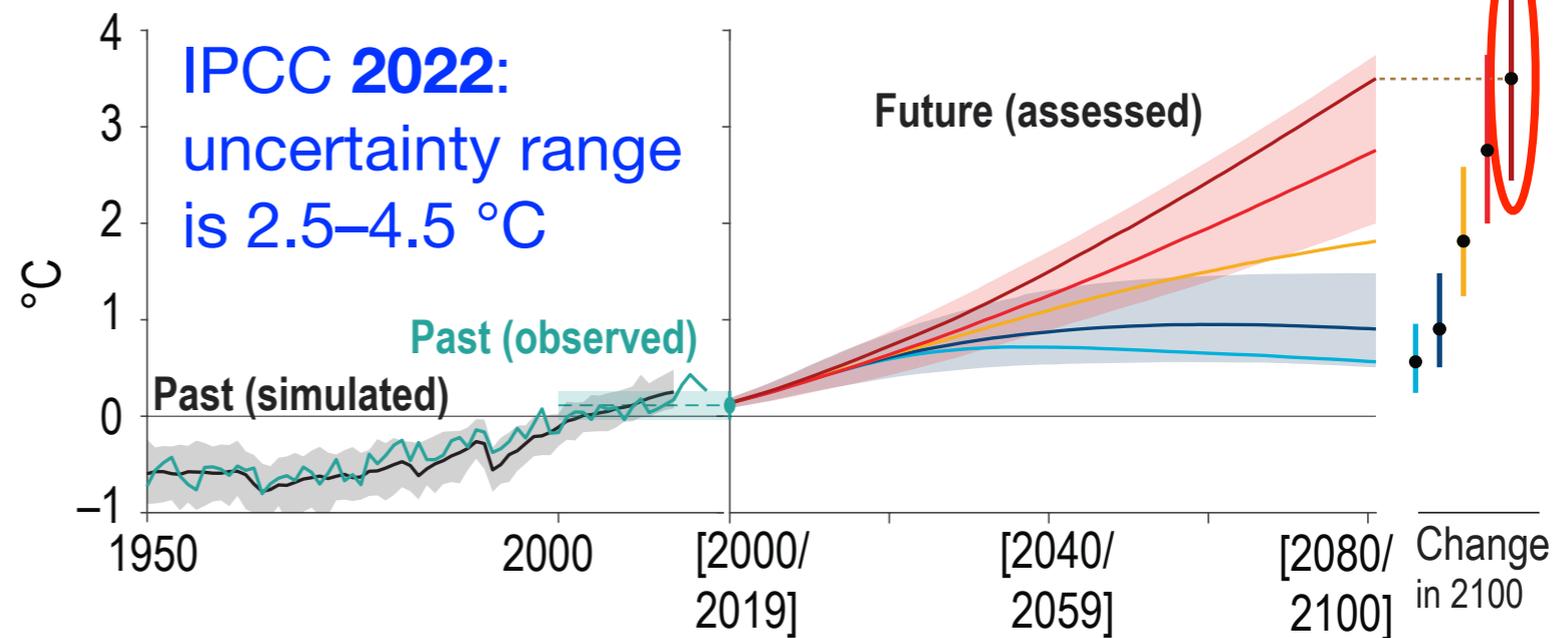


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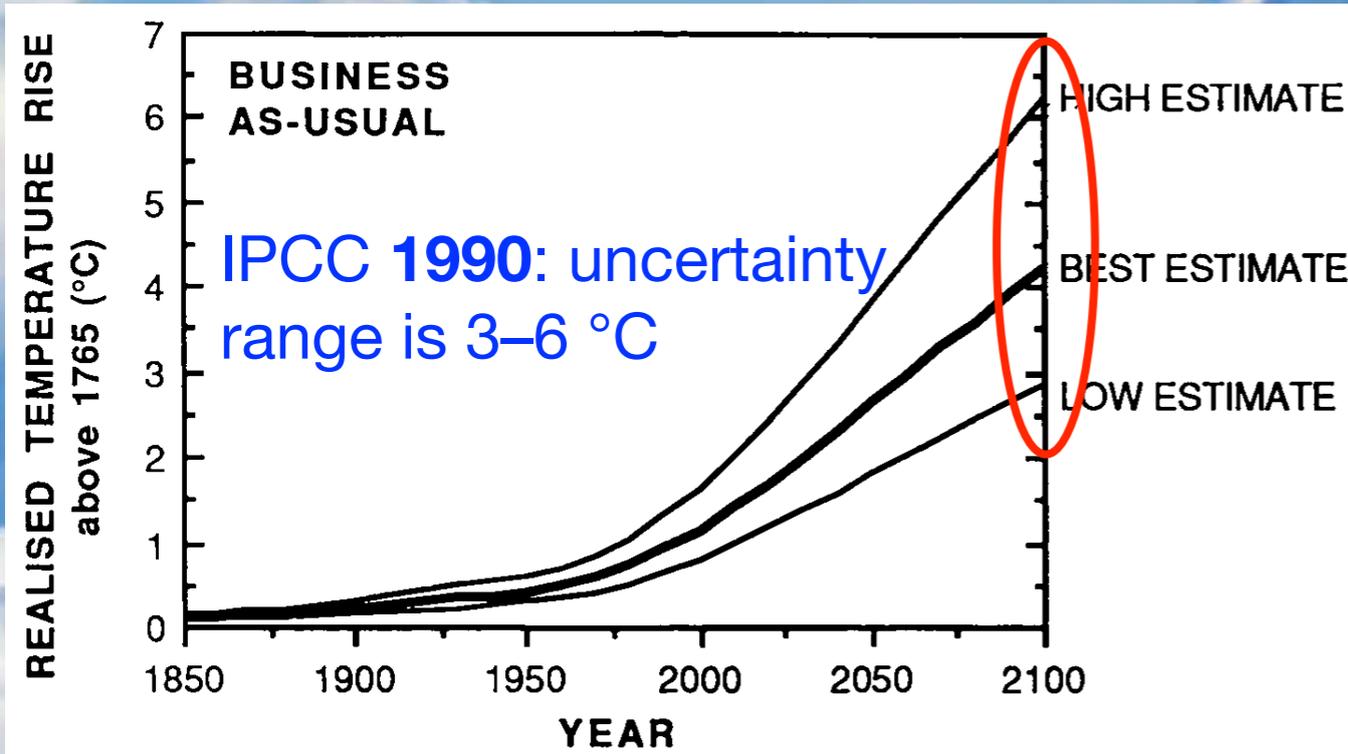
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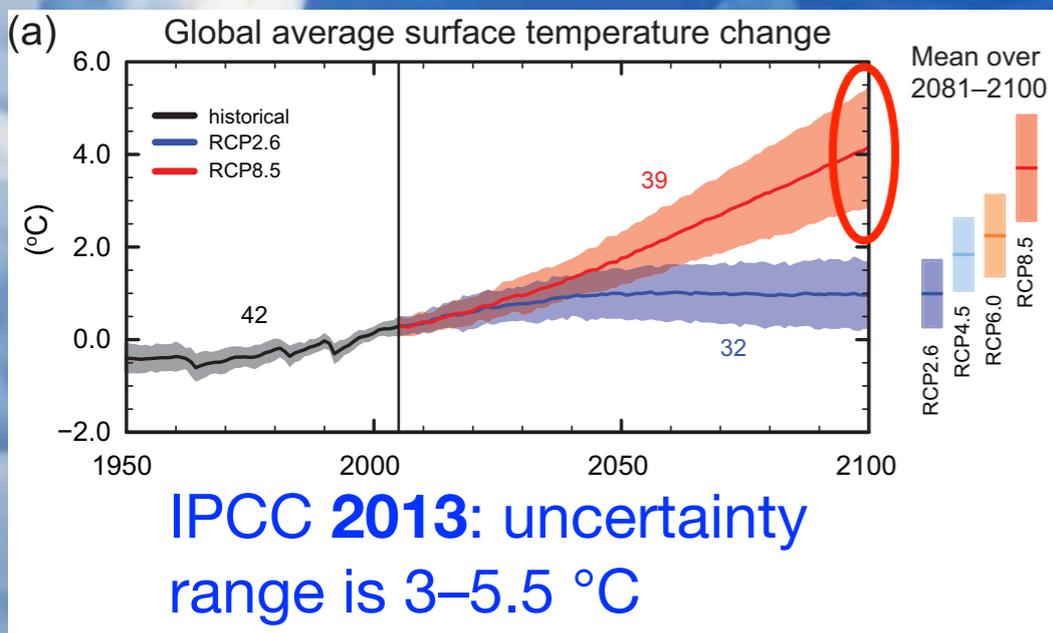
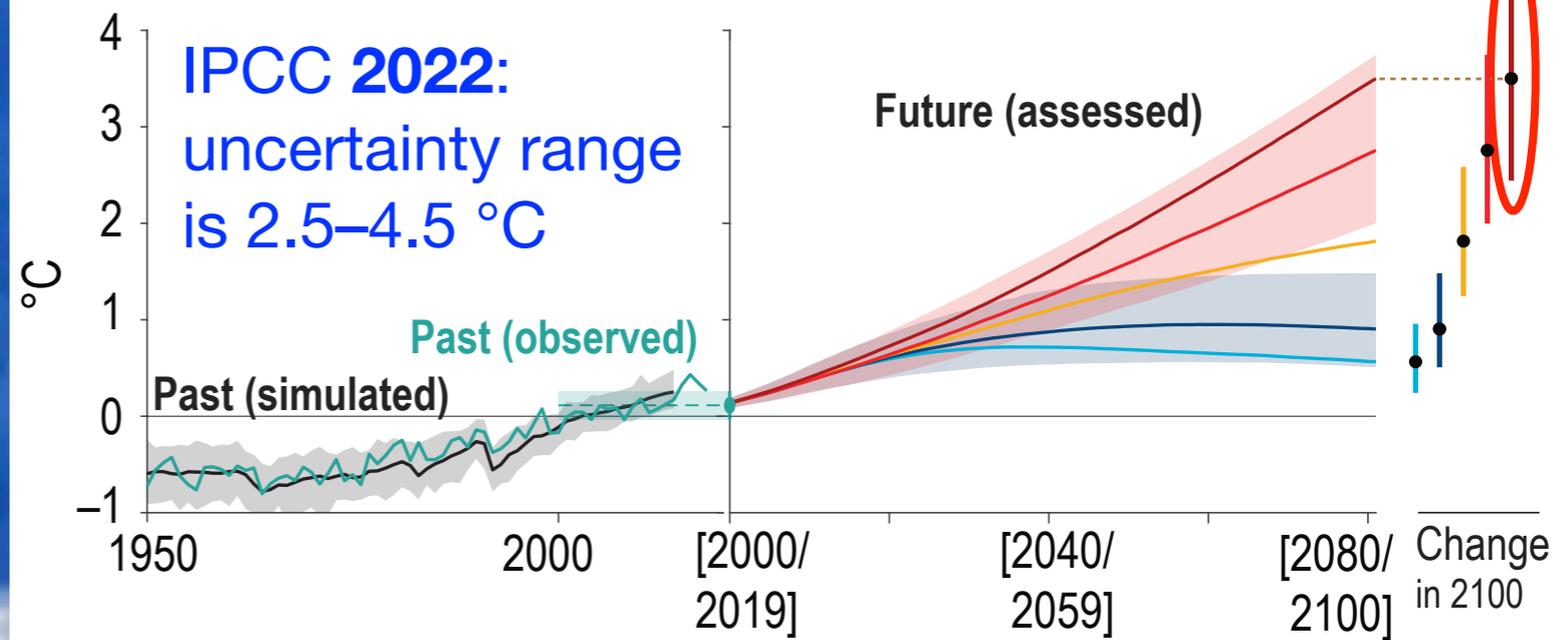
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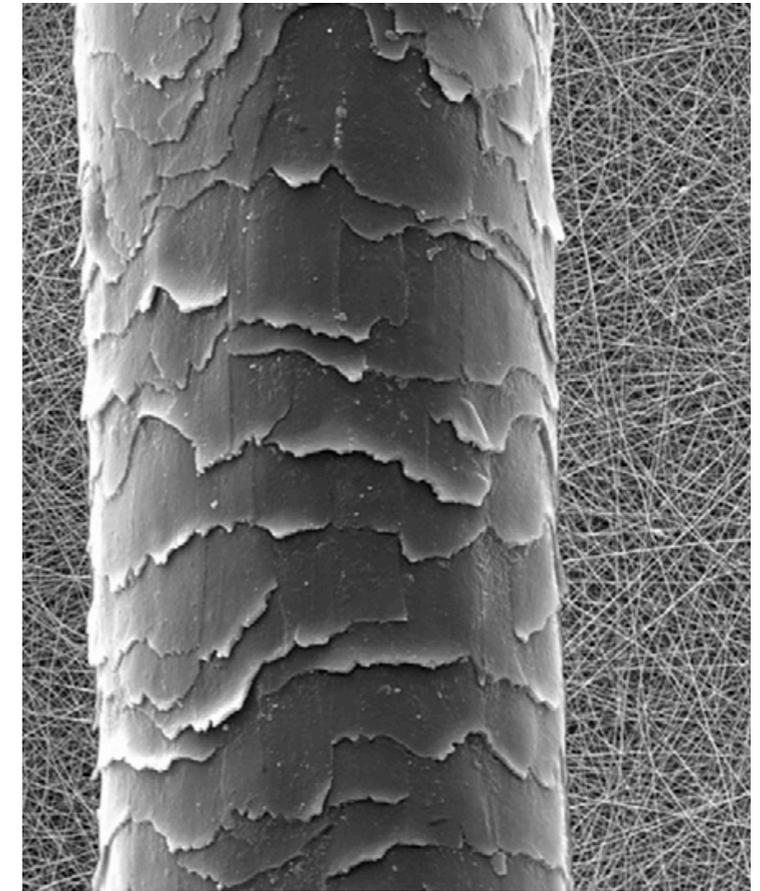


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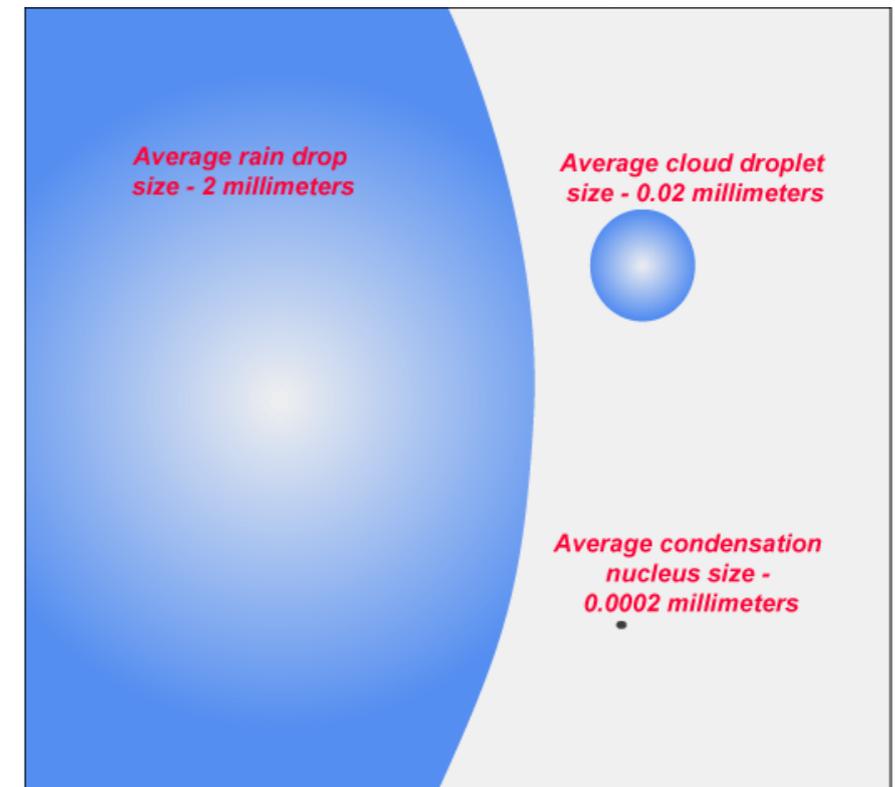
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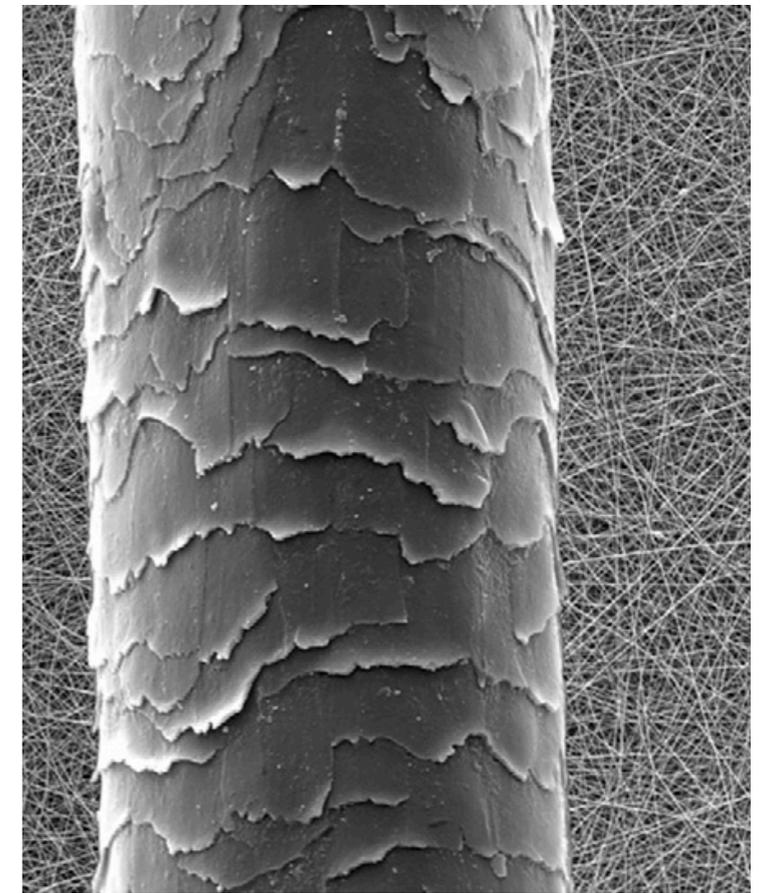


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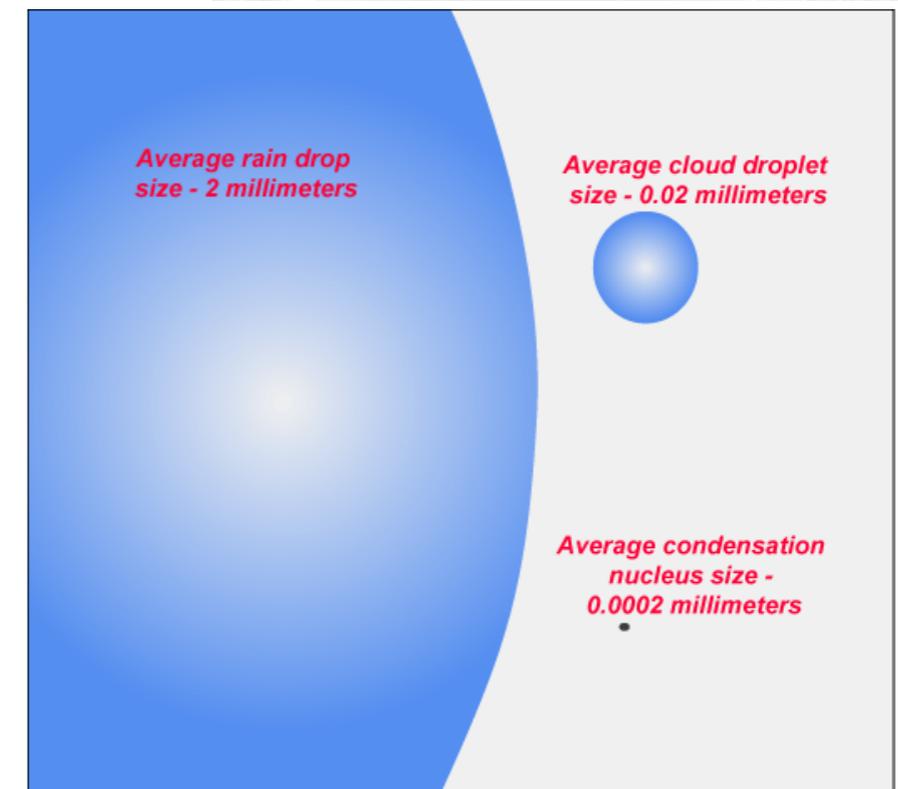
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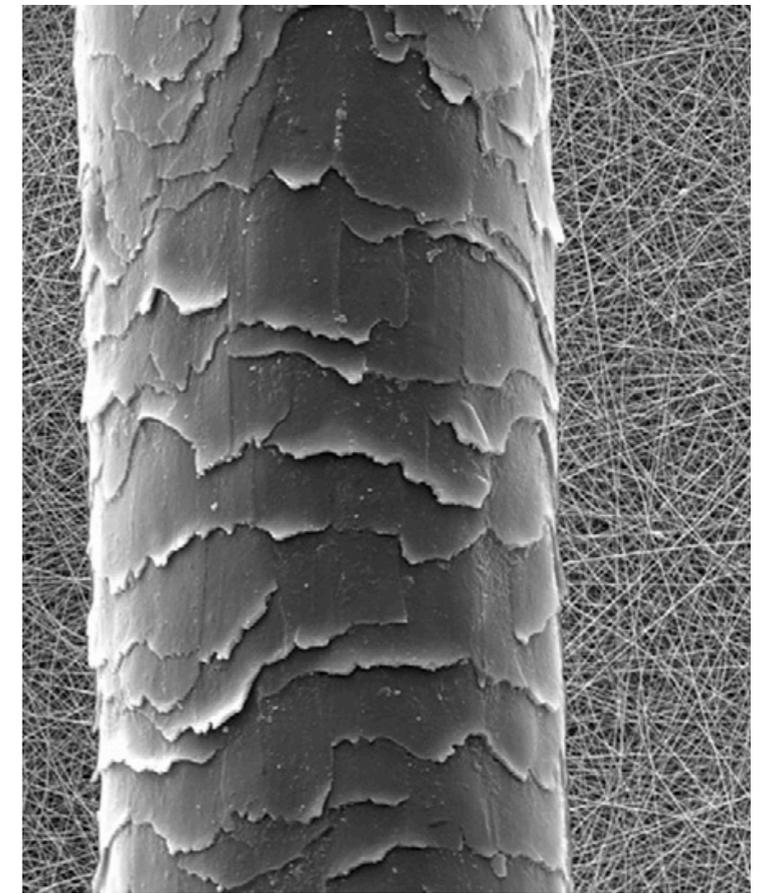


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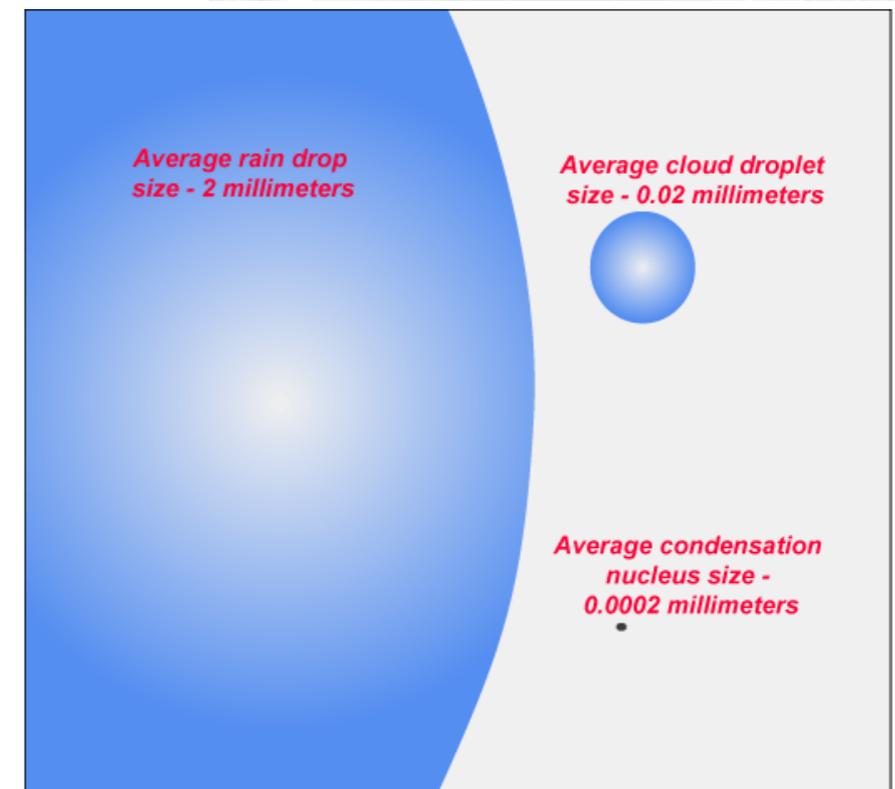
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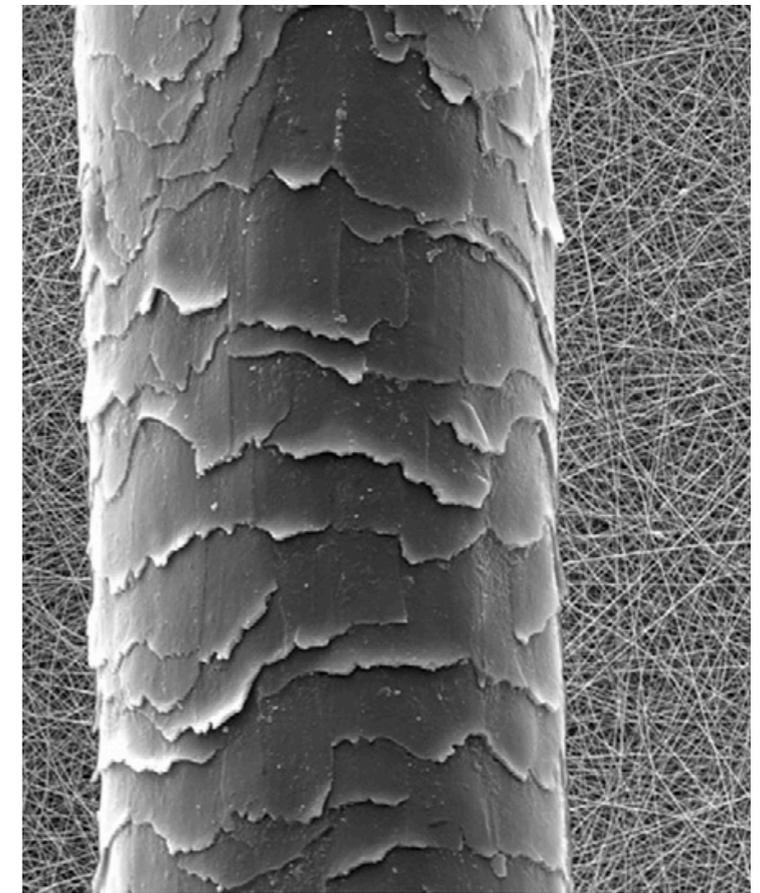


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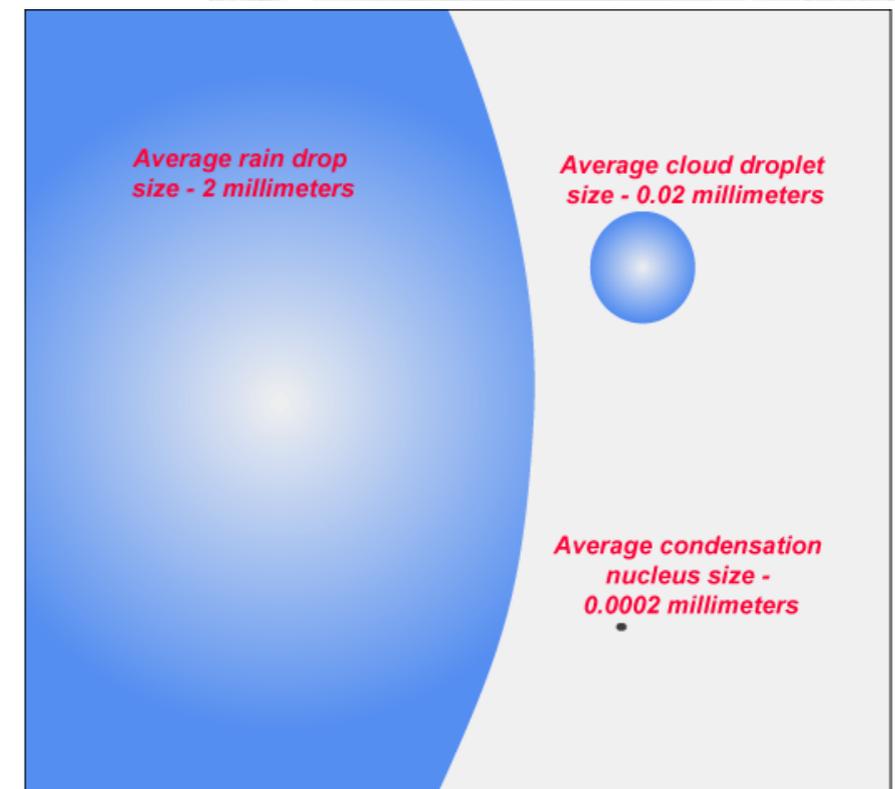
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- The typical distance between droplets in a cloud is ~ 1 mm.

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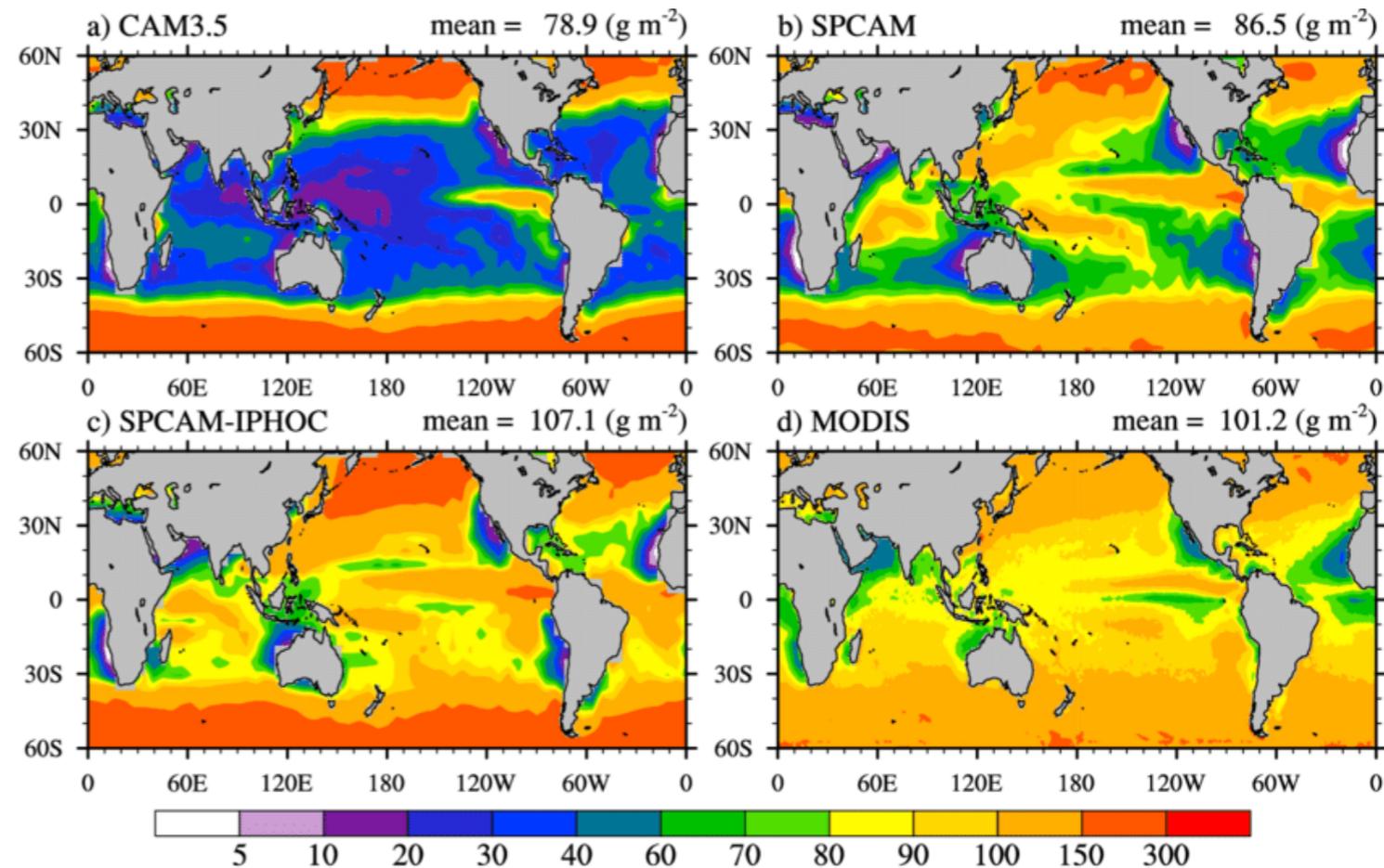
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- Considerable model disagreement on the future spatial distribution of clouds leads to uncertainties in these effects

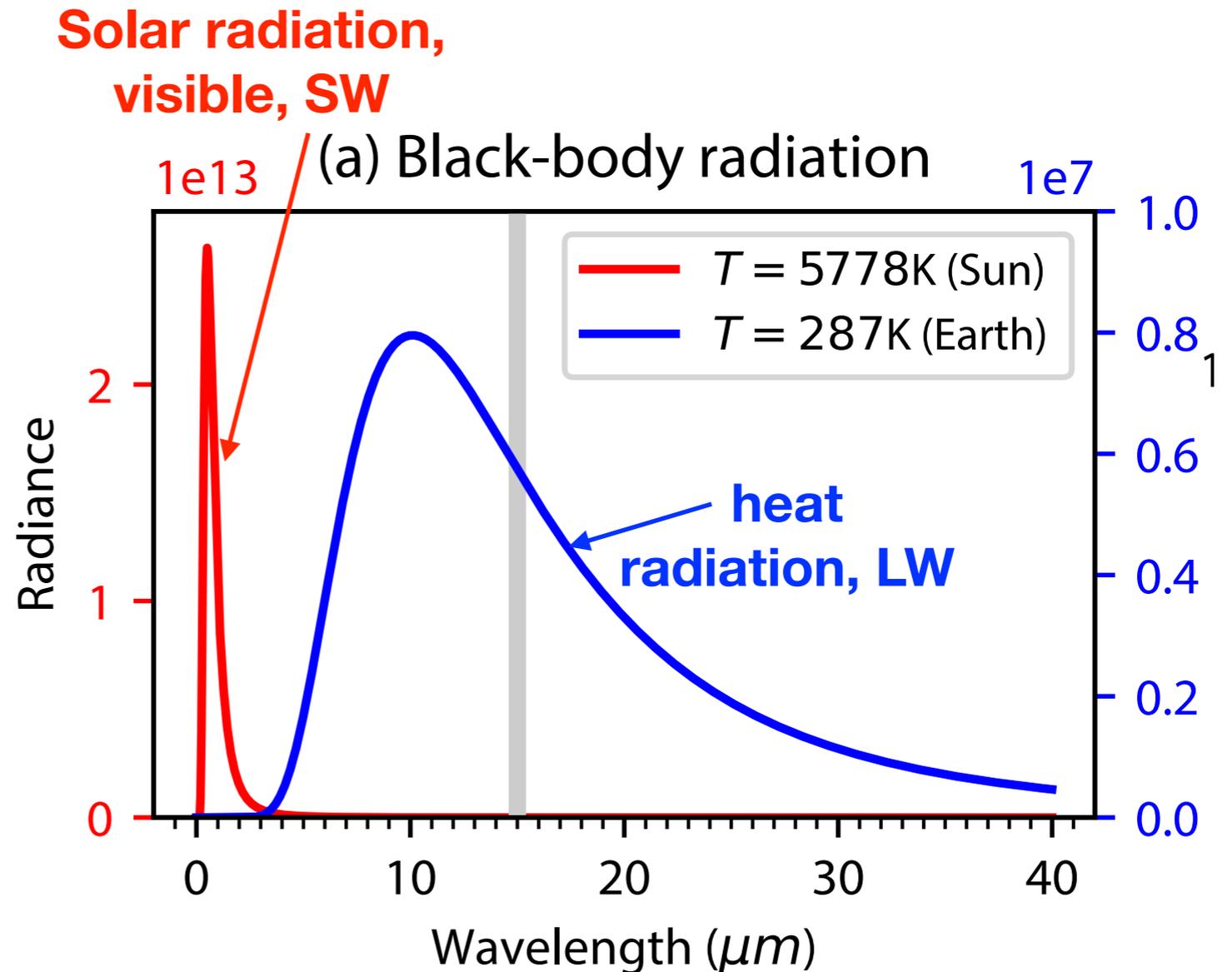


Mean “liquid water path” (total weight (gram) of liquid water in a 1 m^2 air column) of several climate models vs MODIS satellite observations. Although the mean values are comparable, there are large differences in spatial distribution. (Cheng and Xu 2011)

Clouds, radiation and climate

Reminder: Longwave (LW) vs Shortwave (SW) Radiation

- The wavelength of black body radiation emitted from an object depends on its temperature.
- Radiation emitted from the sun has short wavelengths (including visible light, **0.4–0.7 μm**), while radiation emitted from Earth to space has long wavelengths (**5–30 μm**), including Infra-Red)

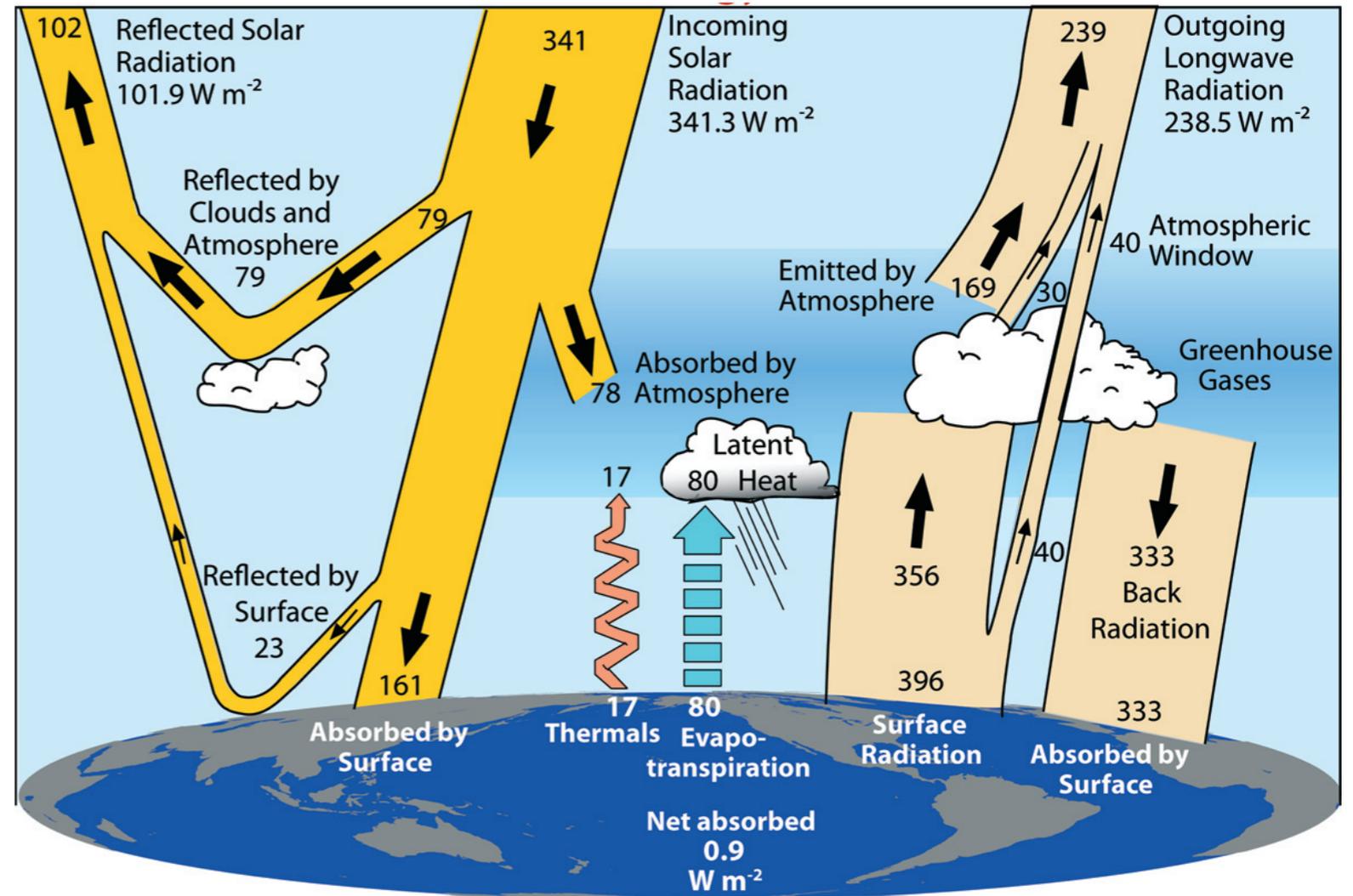


Black body radiation spectra
for the Earth and the sun

(Camille Hankel)

Clouds, radiation and energy balance

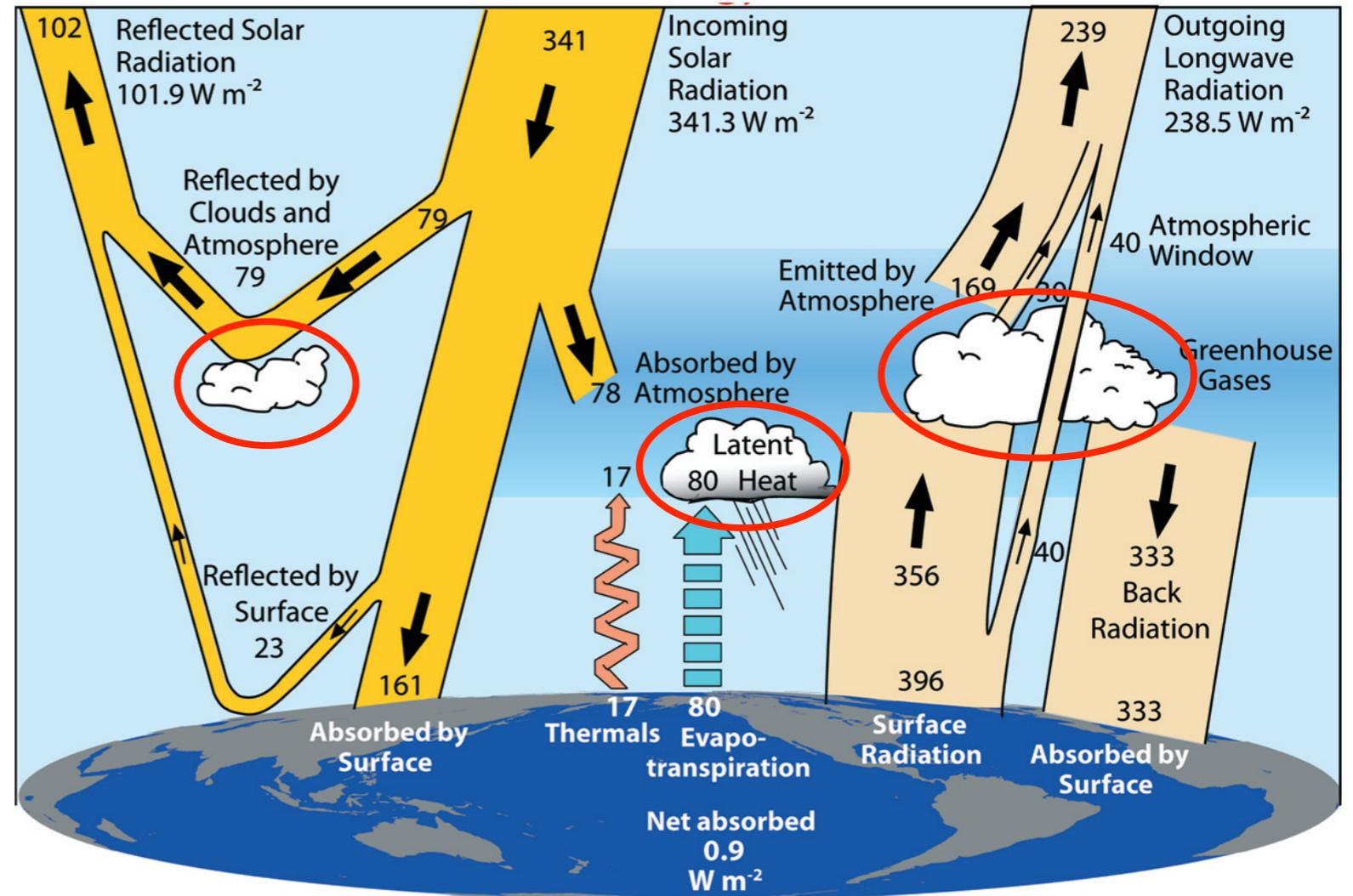
- Clouds are characterized by two competing radiative effects.
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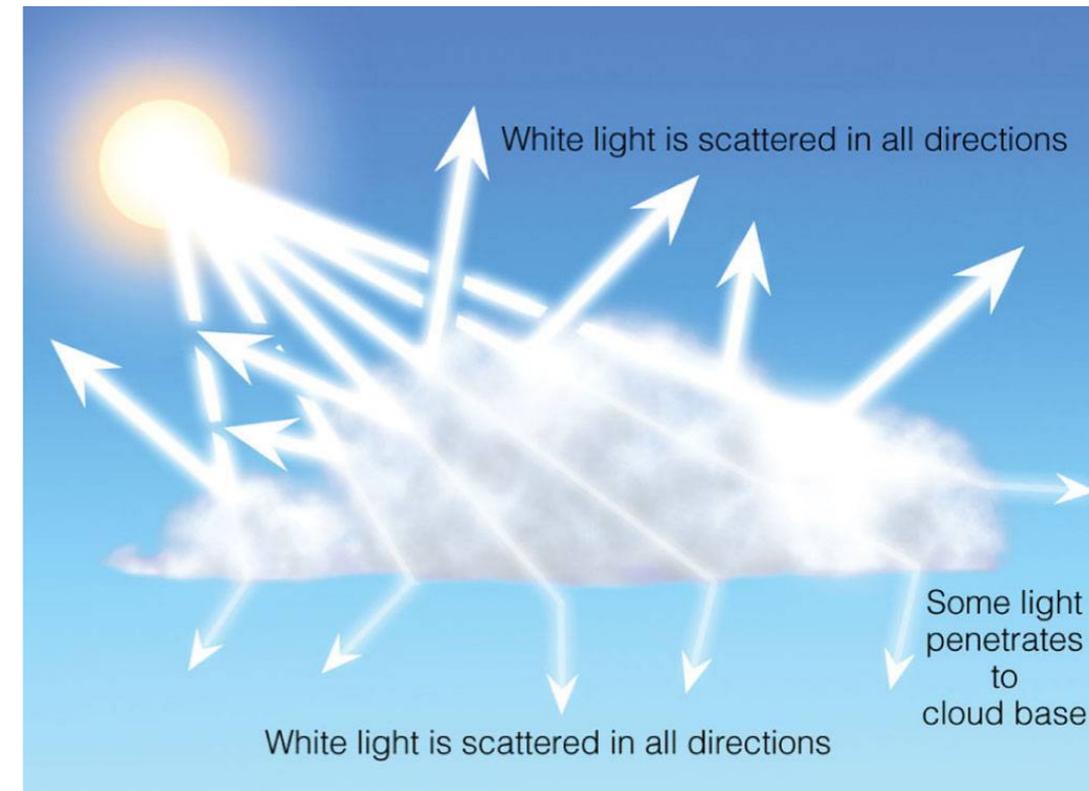
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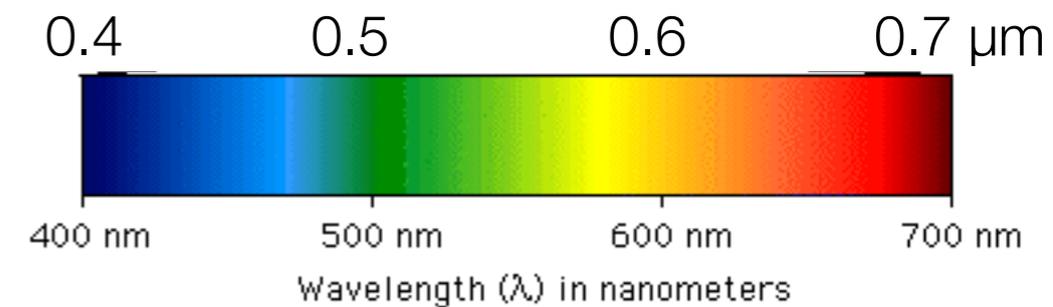
Schematic of SW Cloud Scattering



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“Mie-scattering” Gustav Mie (1868–1957)

http://apollo.lsc.vsc.edu/classes/met130/notes/chapter19/mie_scatt.html



https://web.mst.edu/~gbert/Color_Lg/spec/Aspec.html

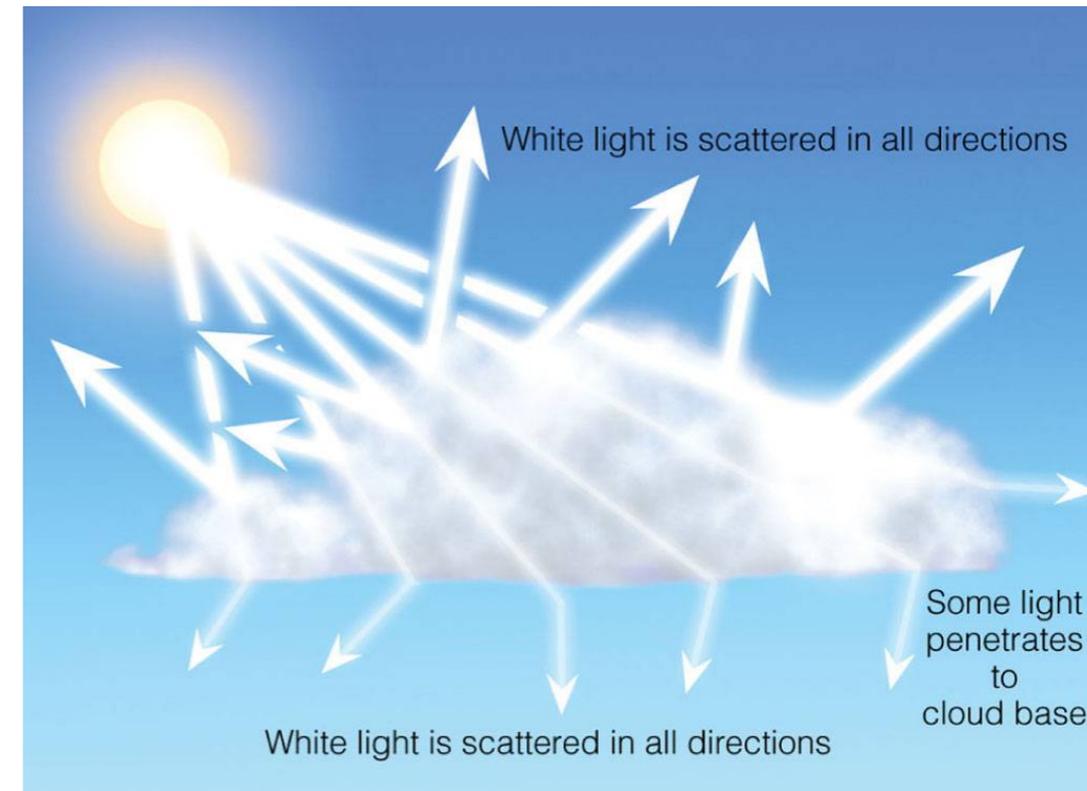
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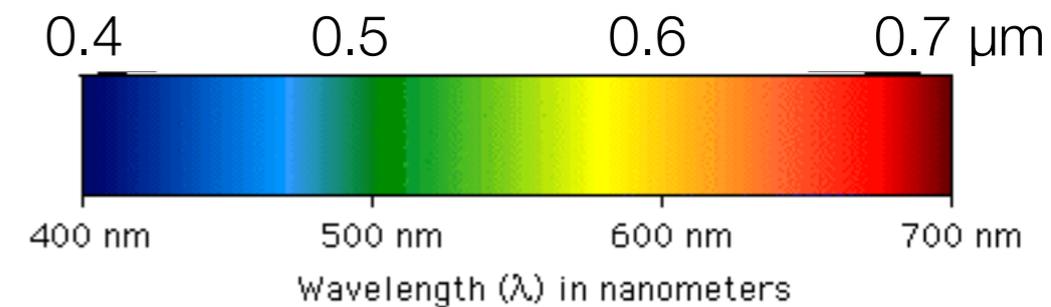
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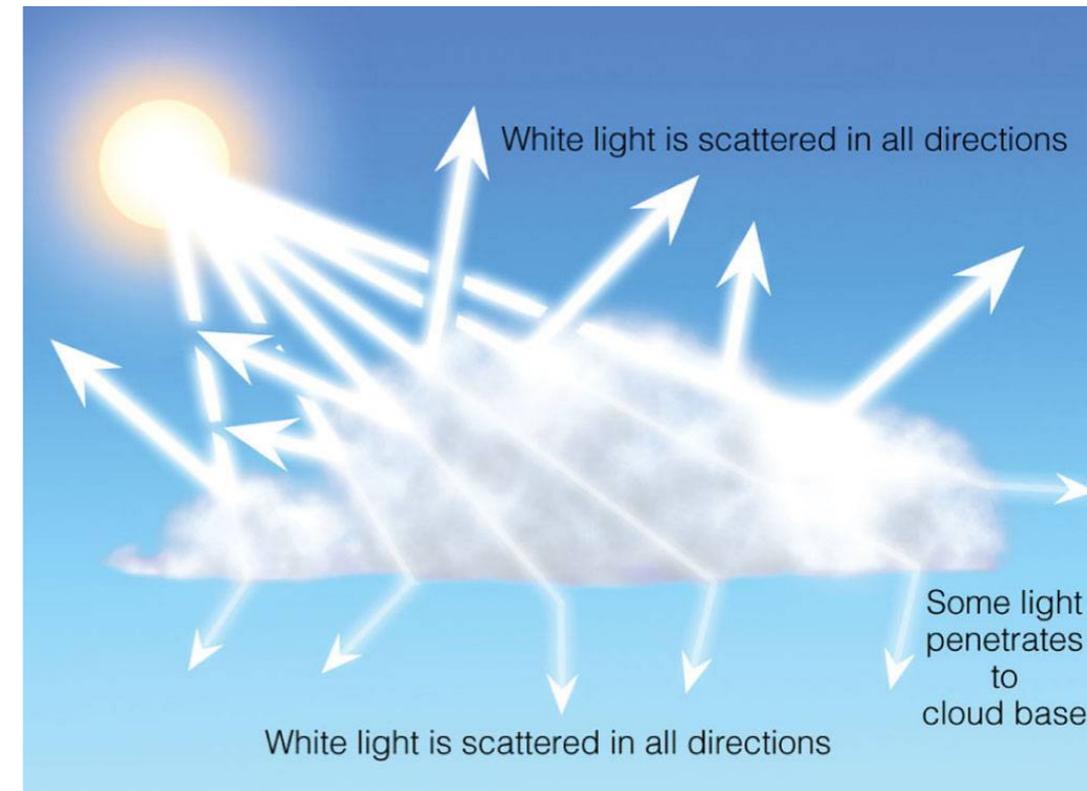
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- ☉ Rayleigh Scattering: by molecules, smaller than wavelengths. Wavelength dependence is $1/\lambda^4$: blue scattered more efficiently than red.

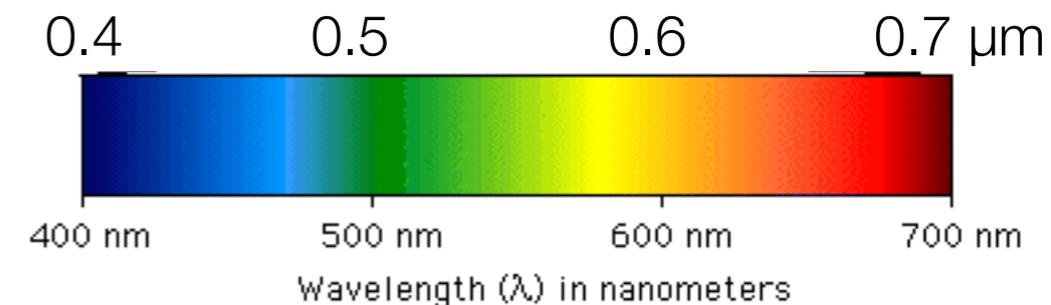
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The sum of longwave and shortwave radiative effect is the “net cloud radiative effect.”

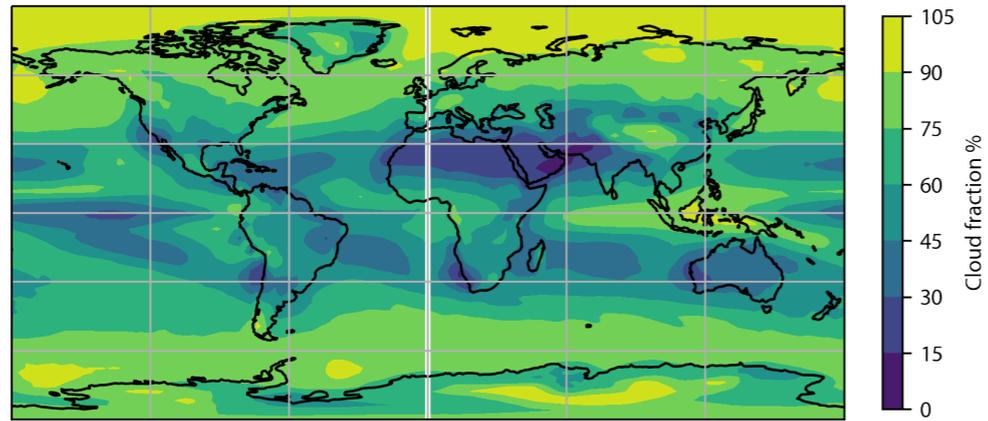
workshop #1

Clouds radiative effect and climate sensitivity in climate models

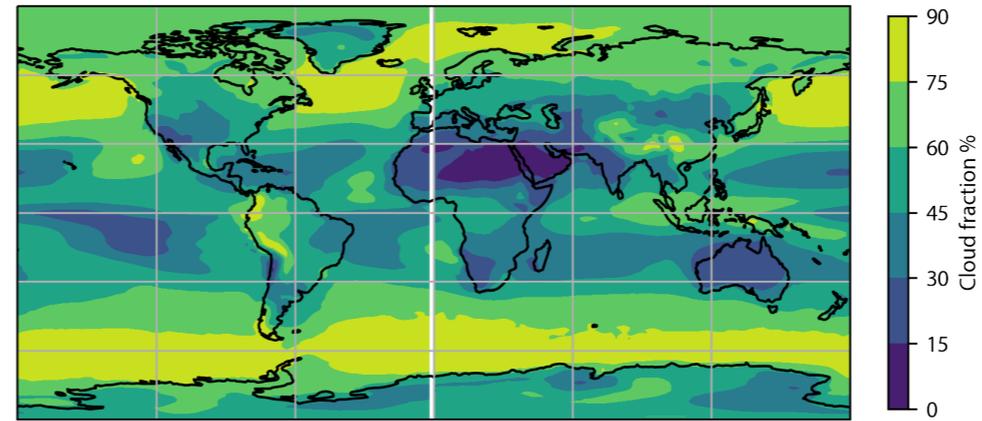
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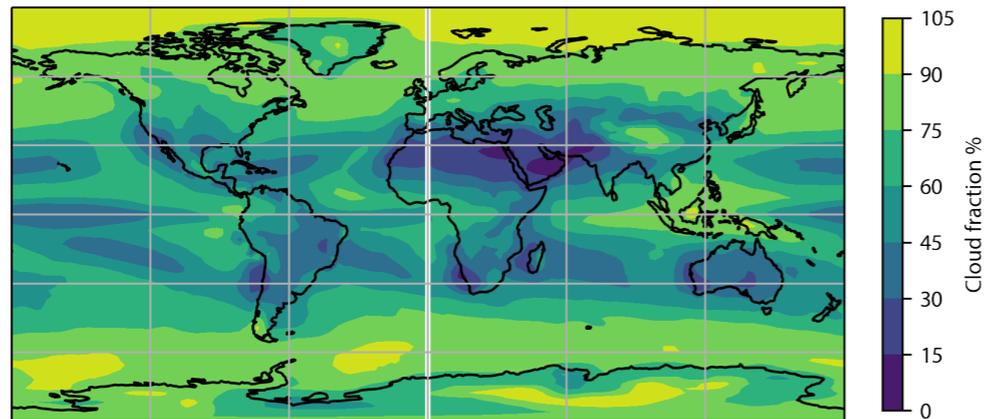
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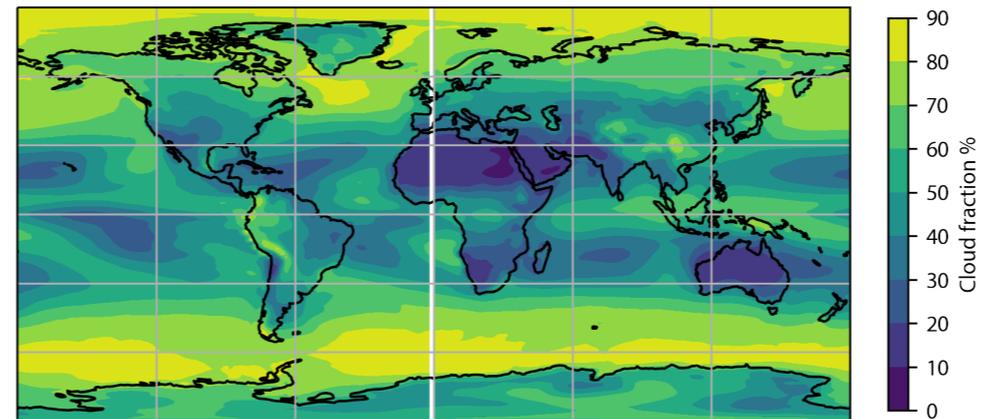
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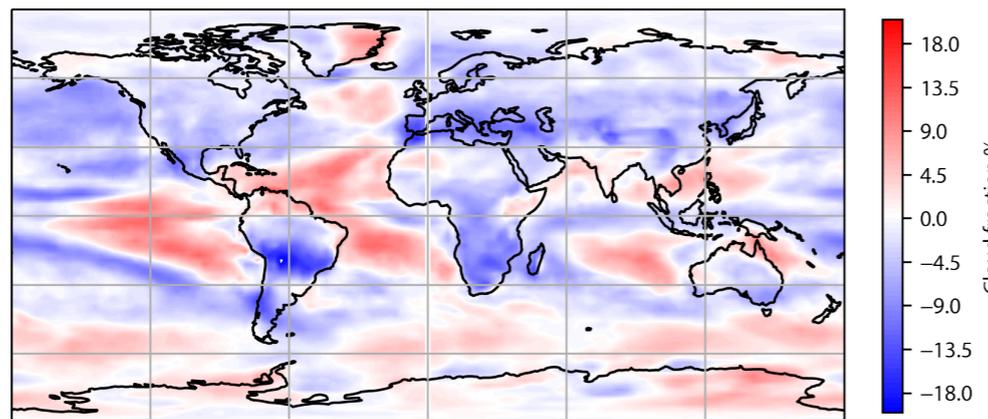
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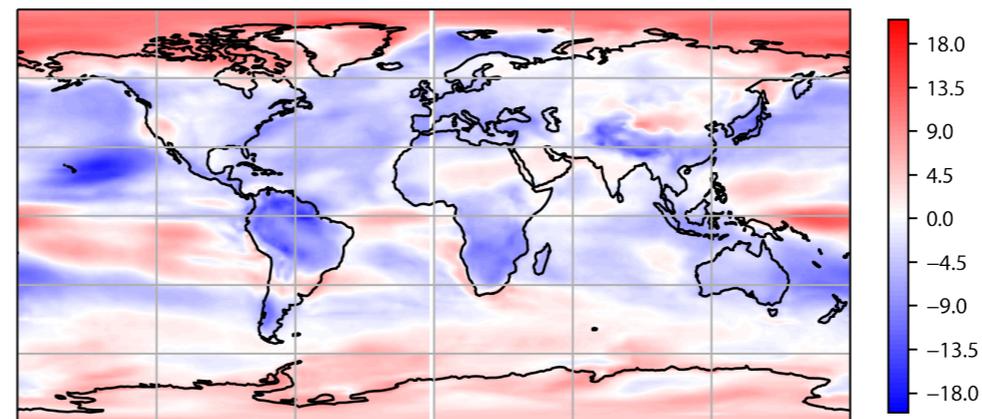
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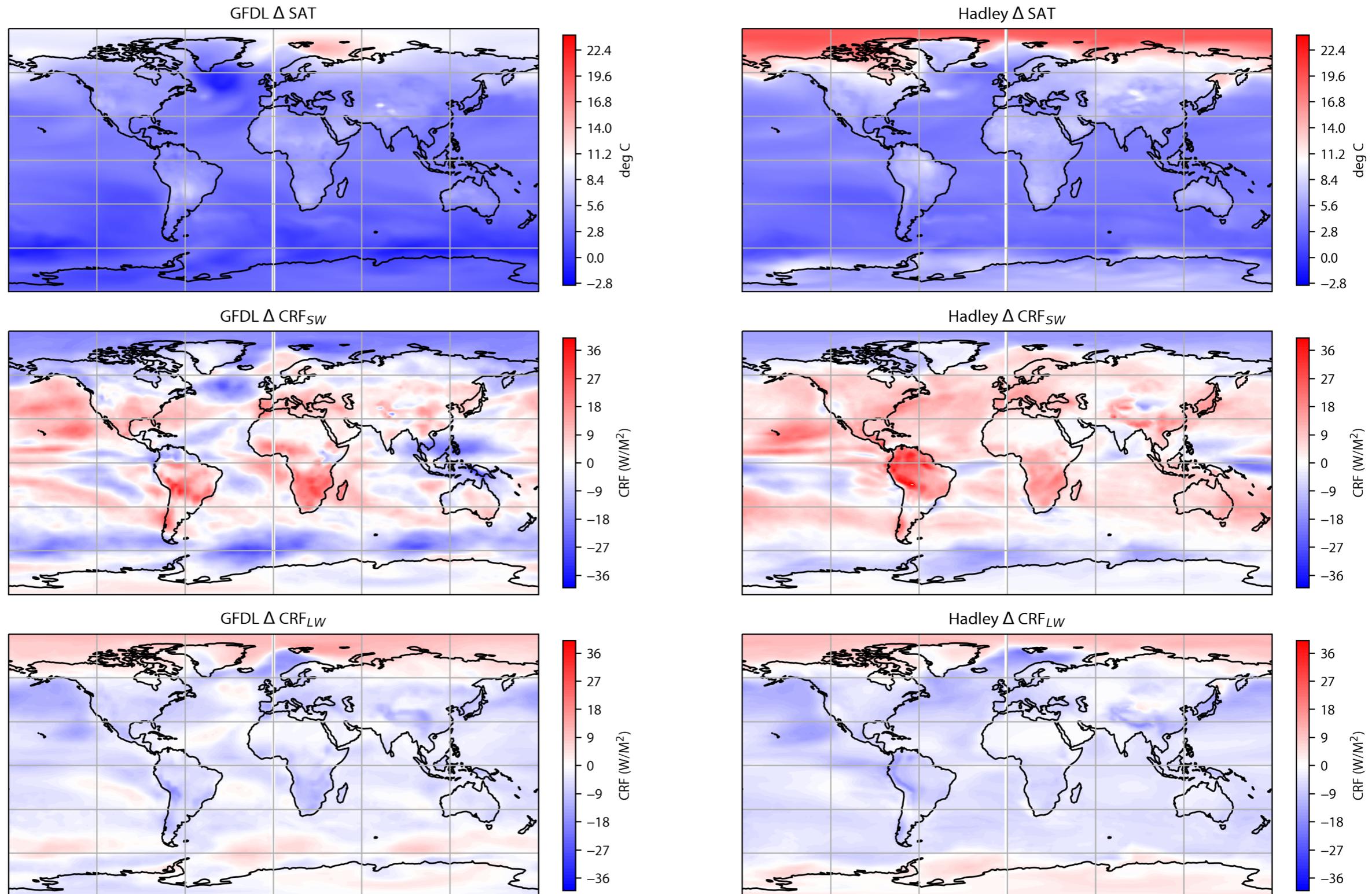


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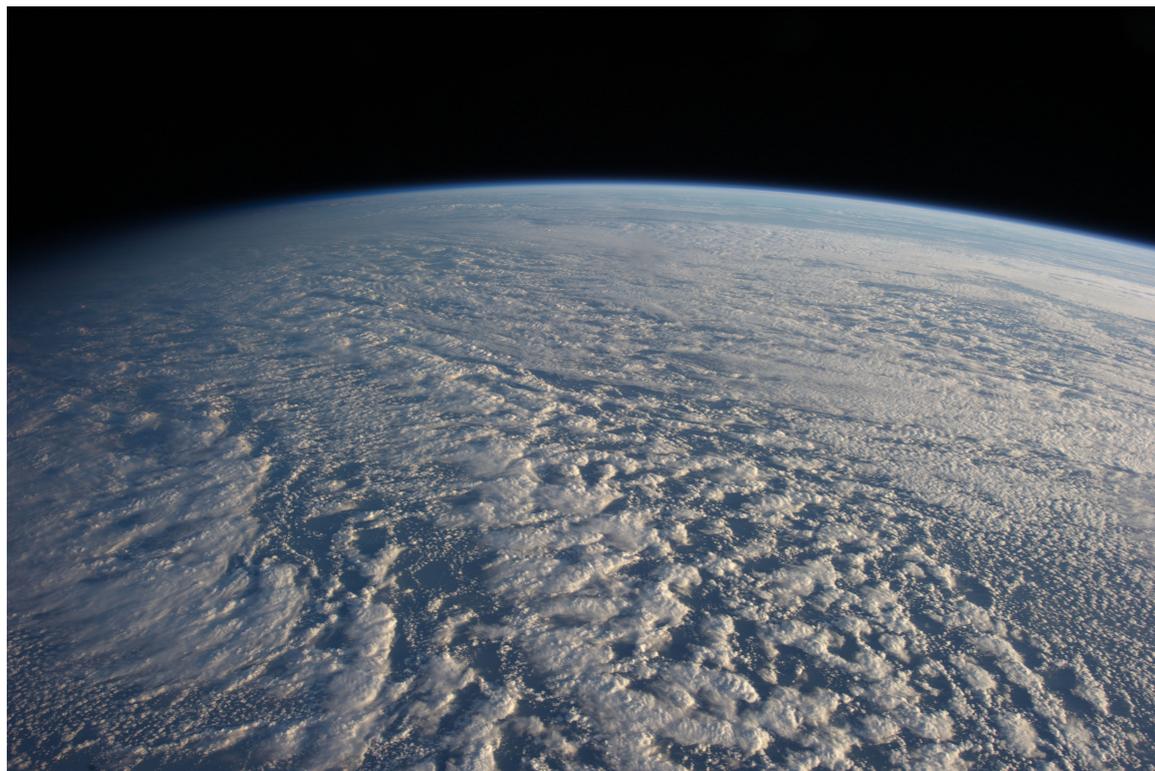
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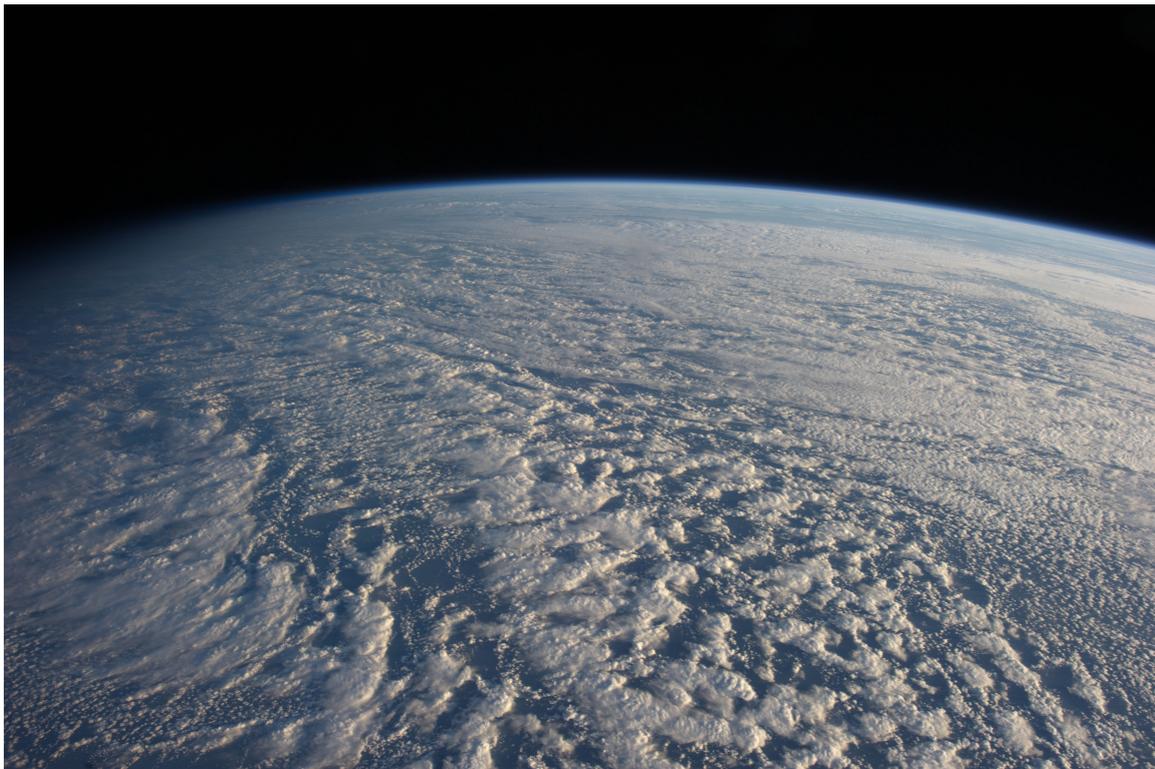
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- ☉ Globally, clouds reduce the absorbed SW radiation by about 47 W/m²



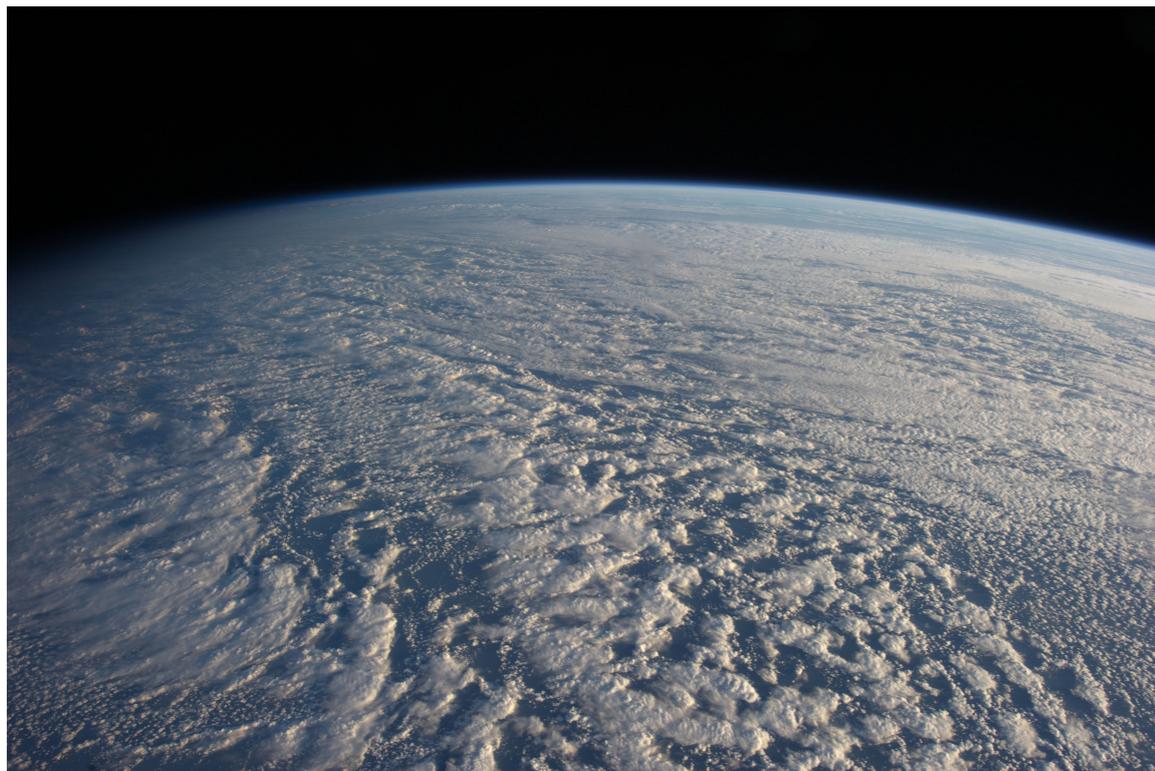
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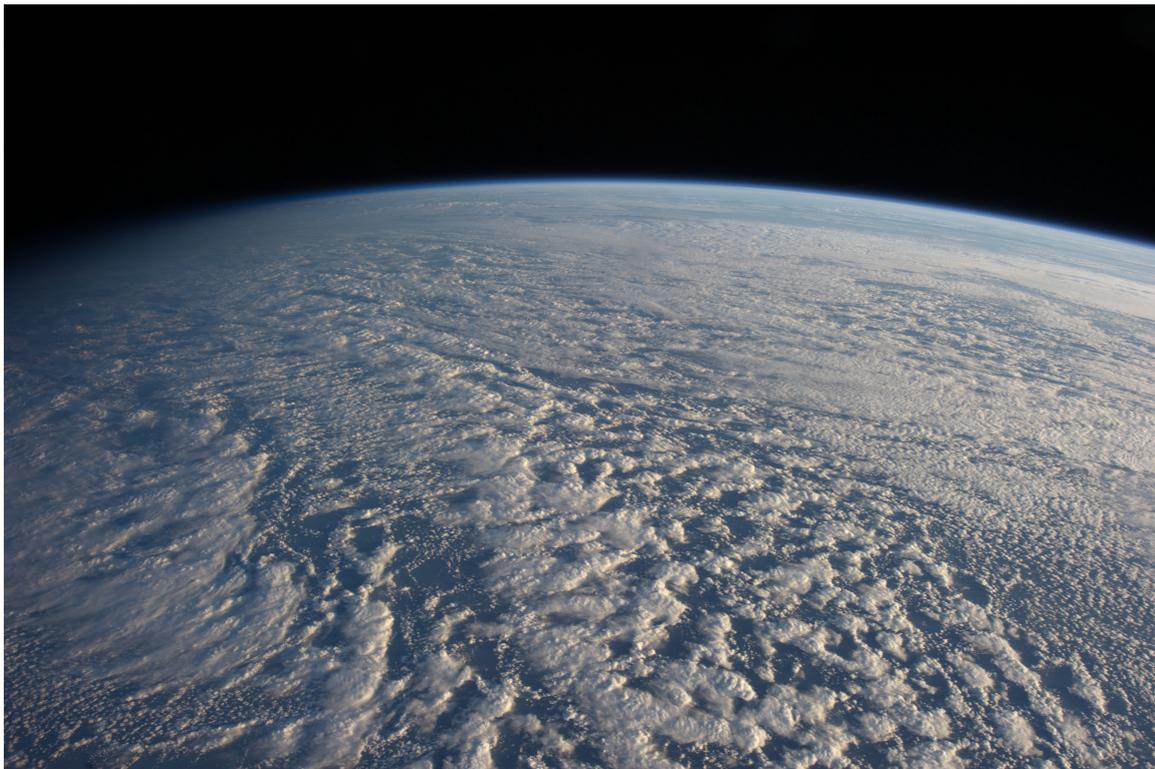
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Stratocumulus: low (lowest 2 km) clouds; high SW albedo, small LW CRE composed of water; **have a net cooling effect.**



cirrus: high (>6 km) clouds; low SW albedo, large LW emissivity, composed of ice; **have a net warming effect.**

Longwave Cloud Radiative Effect

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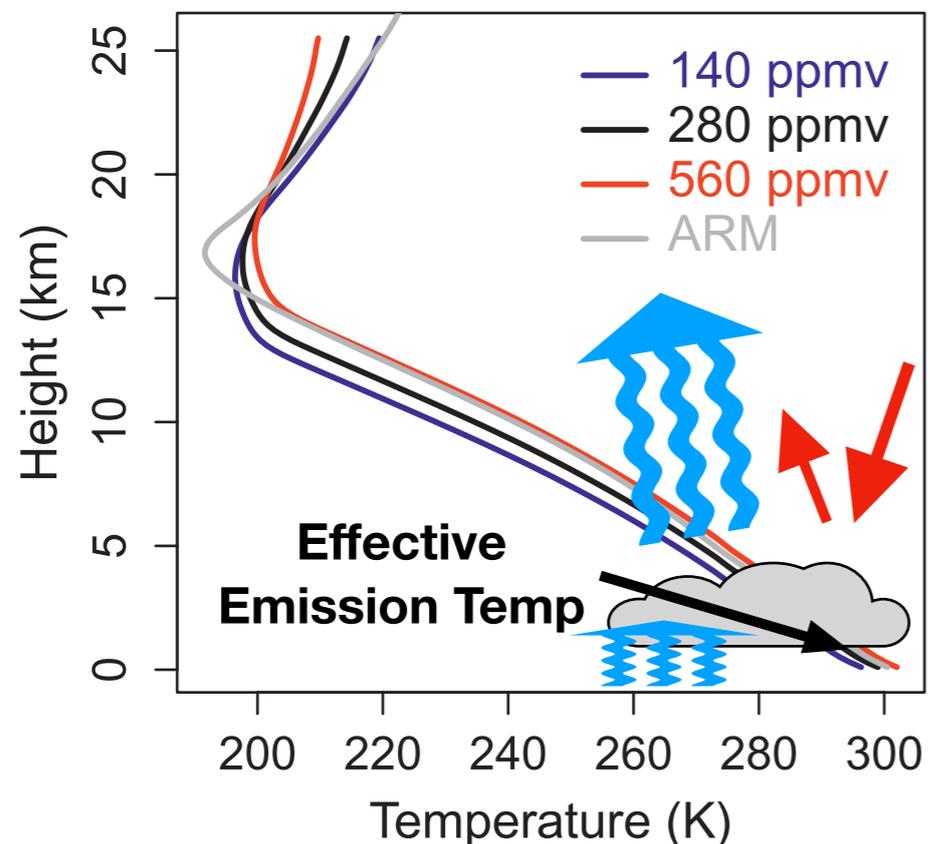
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CRE summary: high cloud warming vs. low cloud cooling

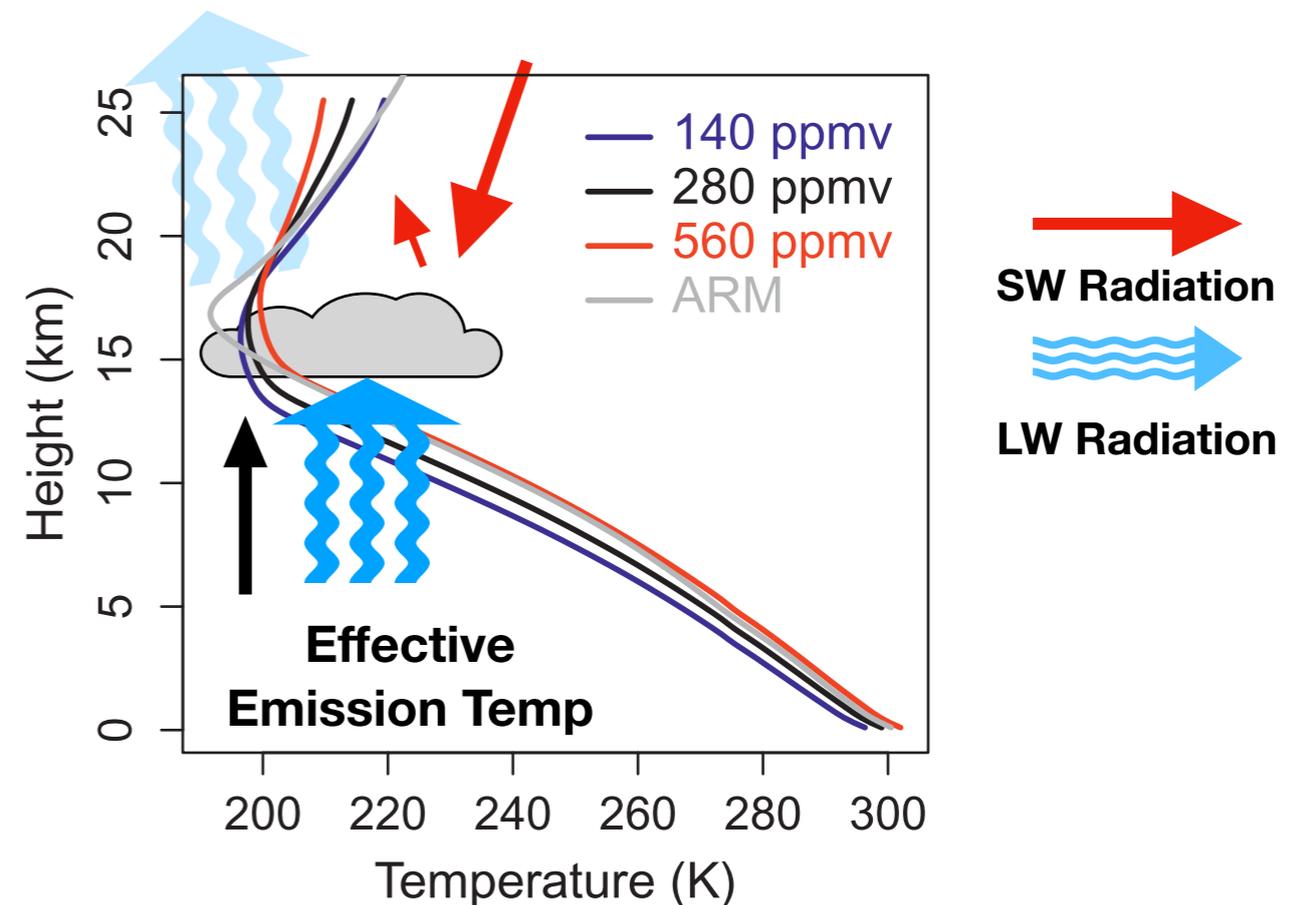
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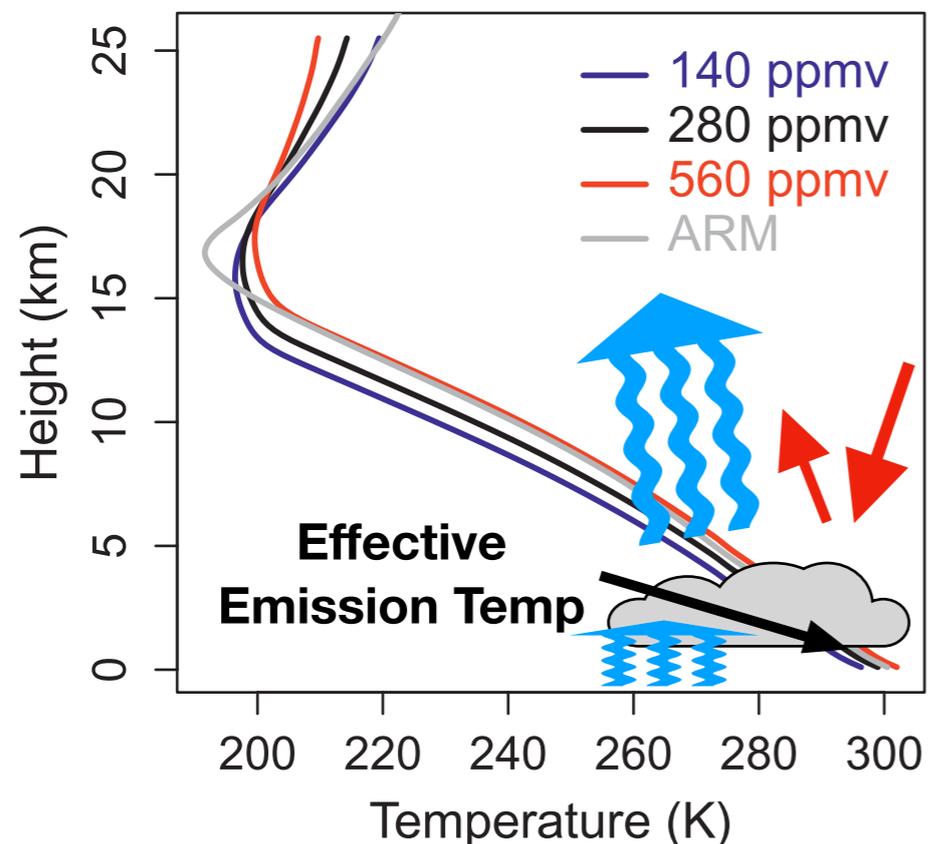
Low Clouds: Small LW CRE, large albedo, Net cooling effect



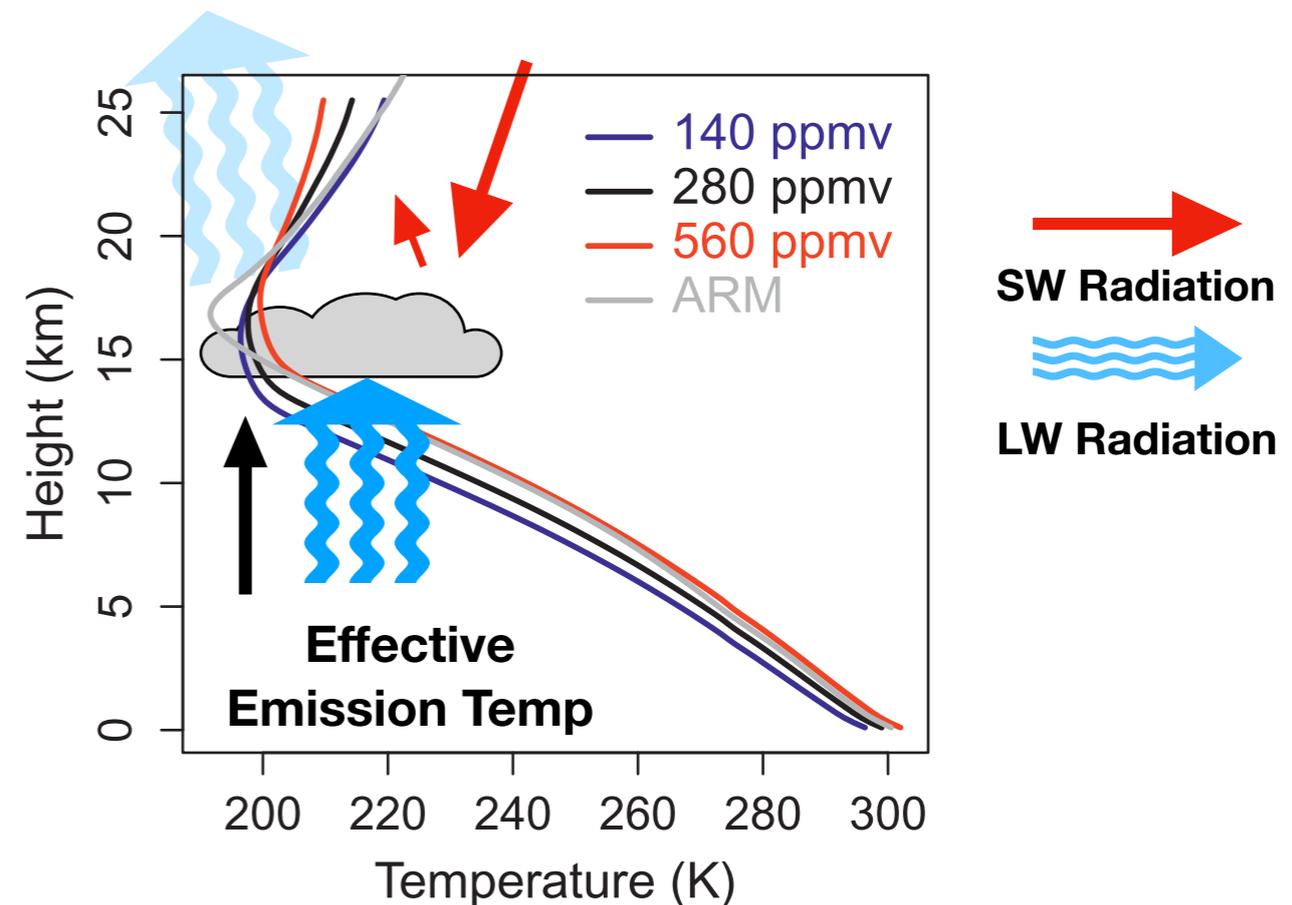
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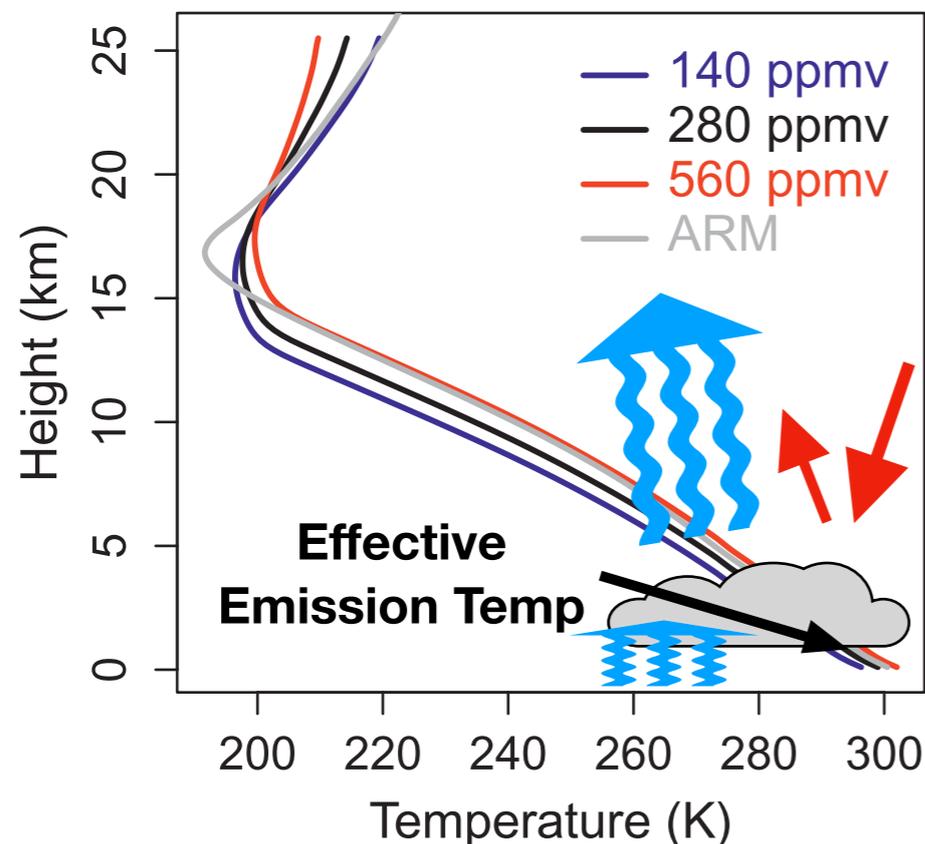
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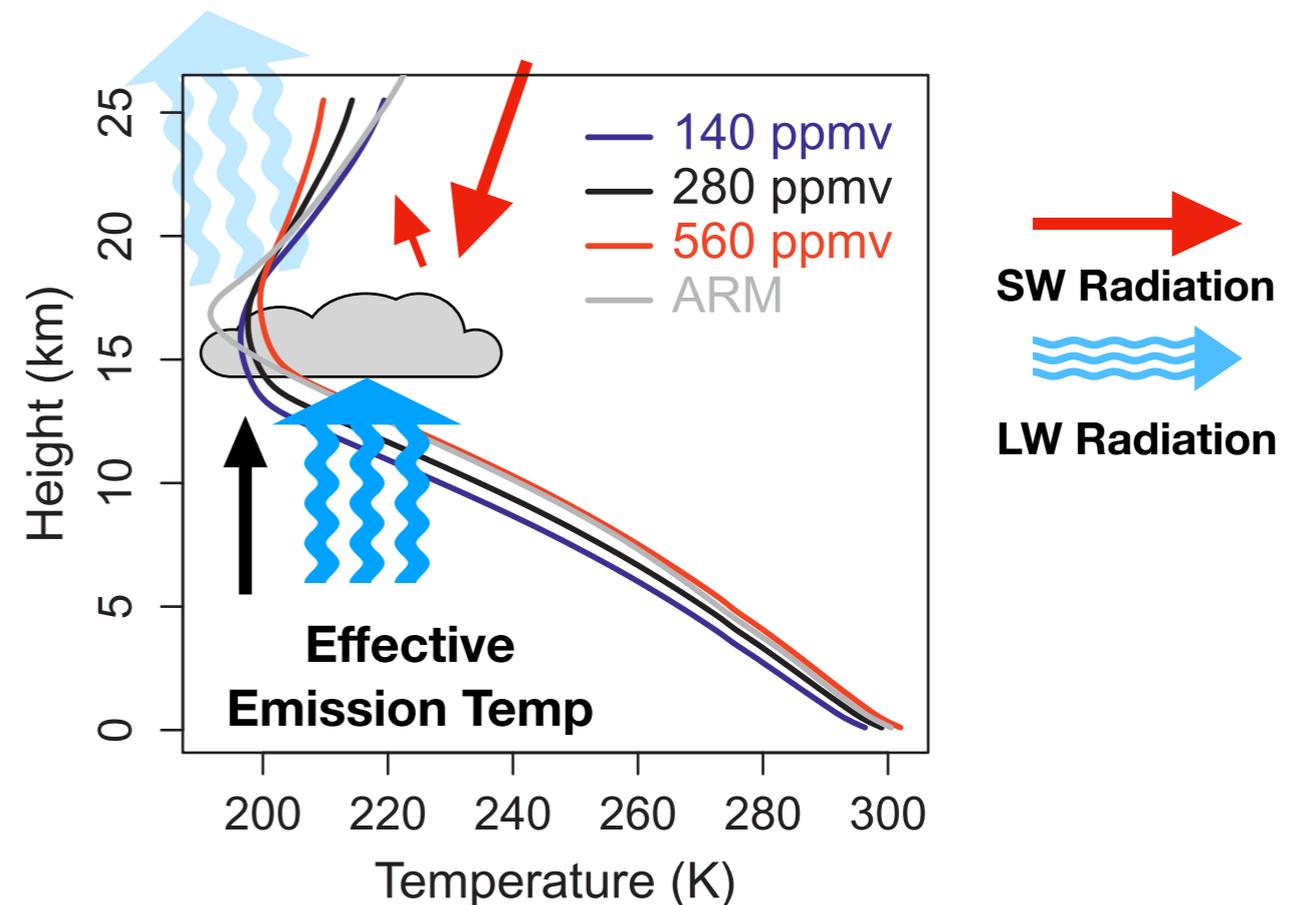
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- **Low clouds** radiate at a temperature close to the surface temperature, radiate upward most of the heat emitted by the surface, and thus **have little LW CRE**.
- Because it takes little water for a cloud to behave as a black body, LW CRE is primarily a function of cloud height. **High clouds**, radiate at a very low temperature and **have a strong longwave warming CRE**.



Low Clouds: Small LW CRE, large albedo, Net cooling effect

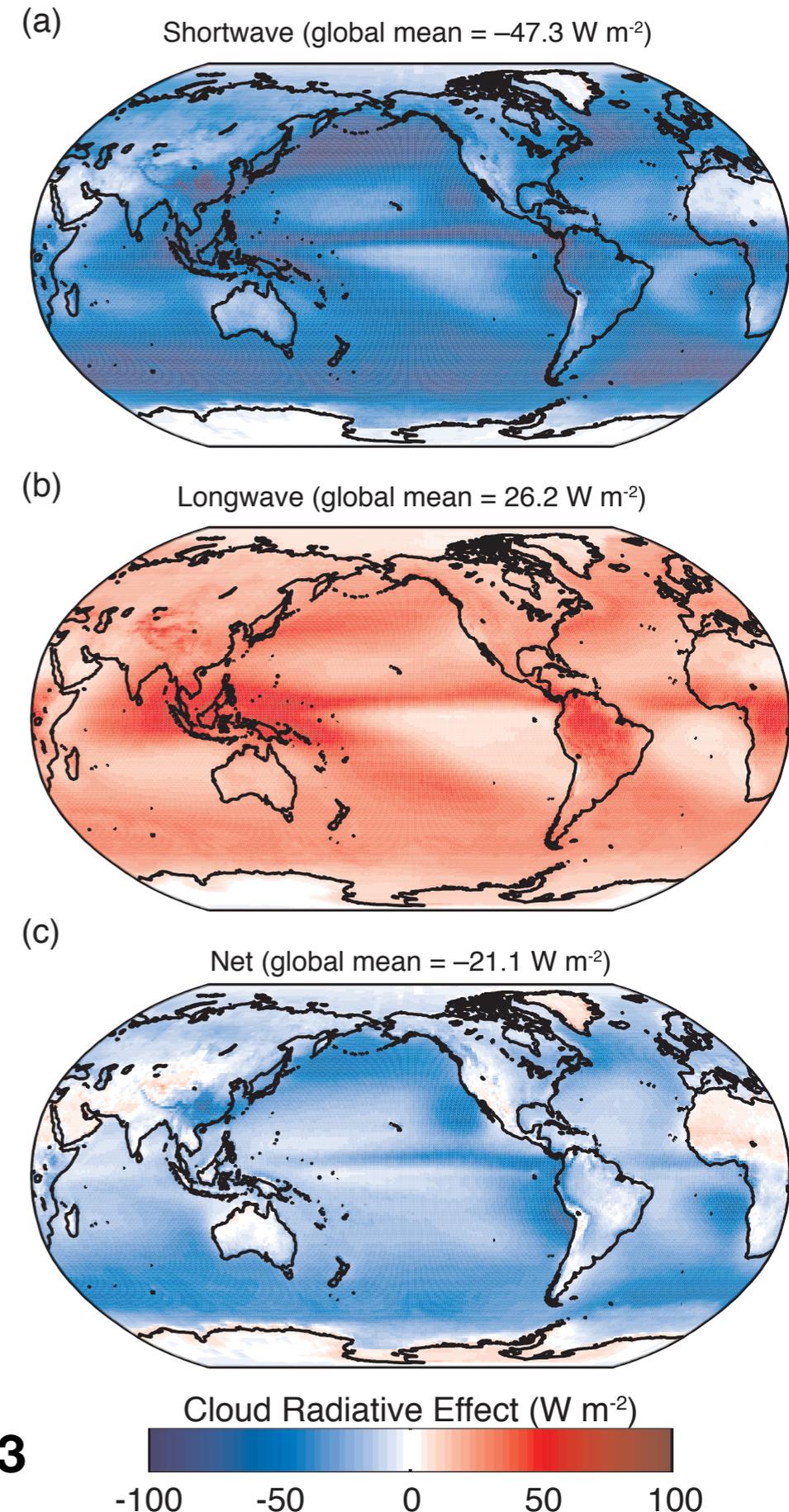


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Longwave vs Shortwave Cloud Radiative Effect

The sum of longwave and shortwave radiative effect is the “net cloud radiative effect” (CRE).

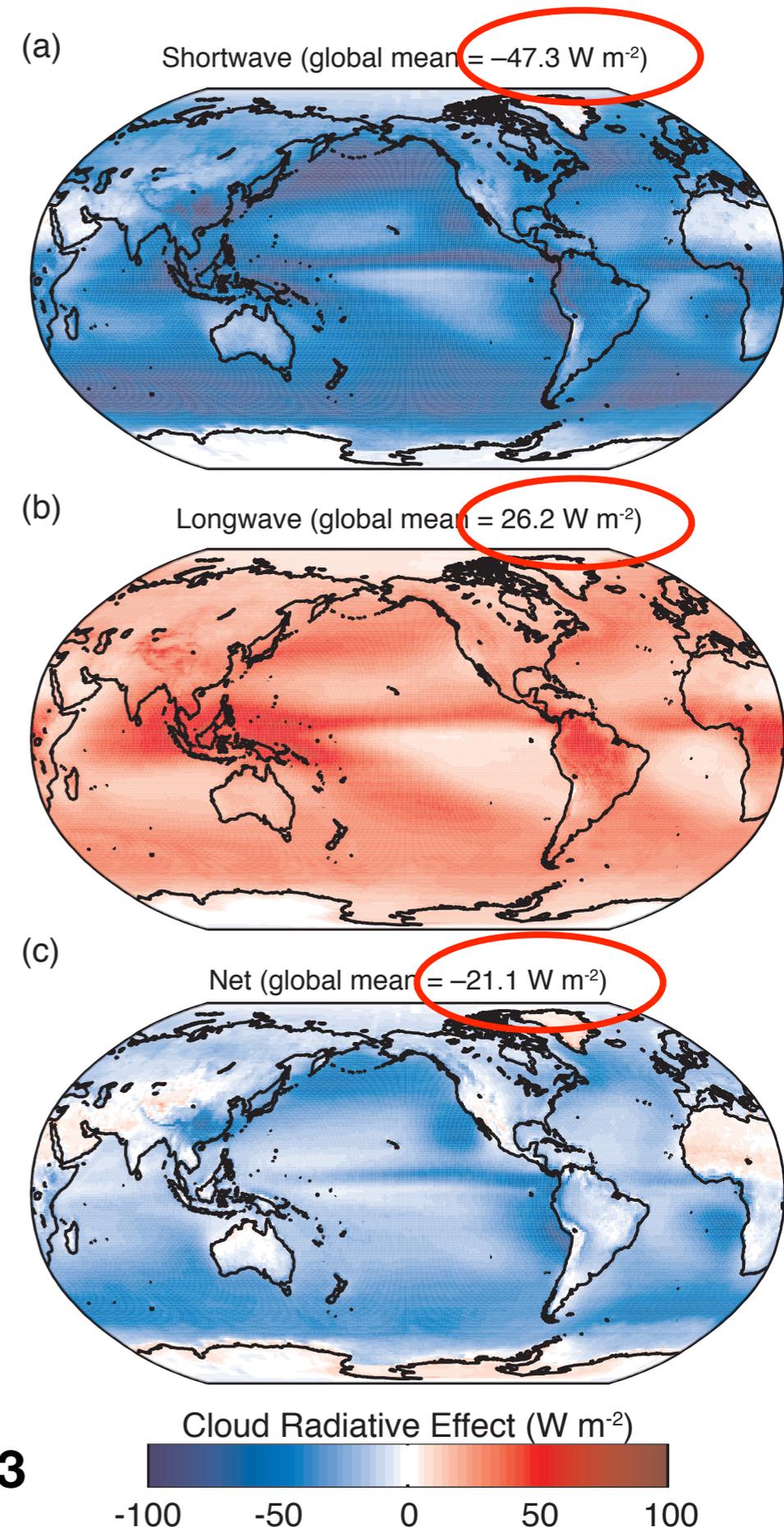
Figure 7.7 | Distribution of annual-mean top of the atmosphere (a) shortwave, (b) longwave, (c) net cloud radiative effects averaged over the period 2001–2011 from the Clouds and the Earth’s Radiant Energy System (CERES) Energy Balanced and Filled (EBAF) Ed2.6r data set (Loeb et al., 2009) and (d) precipitation rate (1981–2000 average from the GPCP version 2.2 data set; Adler et al., 2003).



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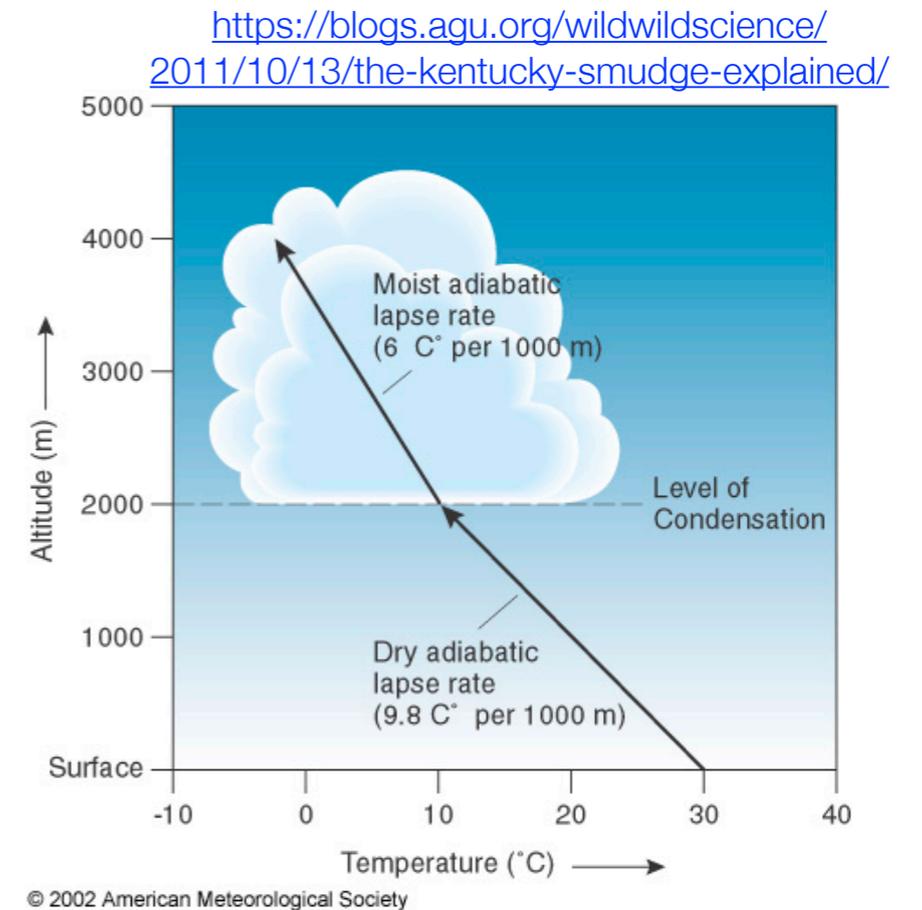
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Cloud formation

How do clouds form?

- Clouds generally form in rising air (updraft).

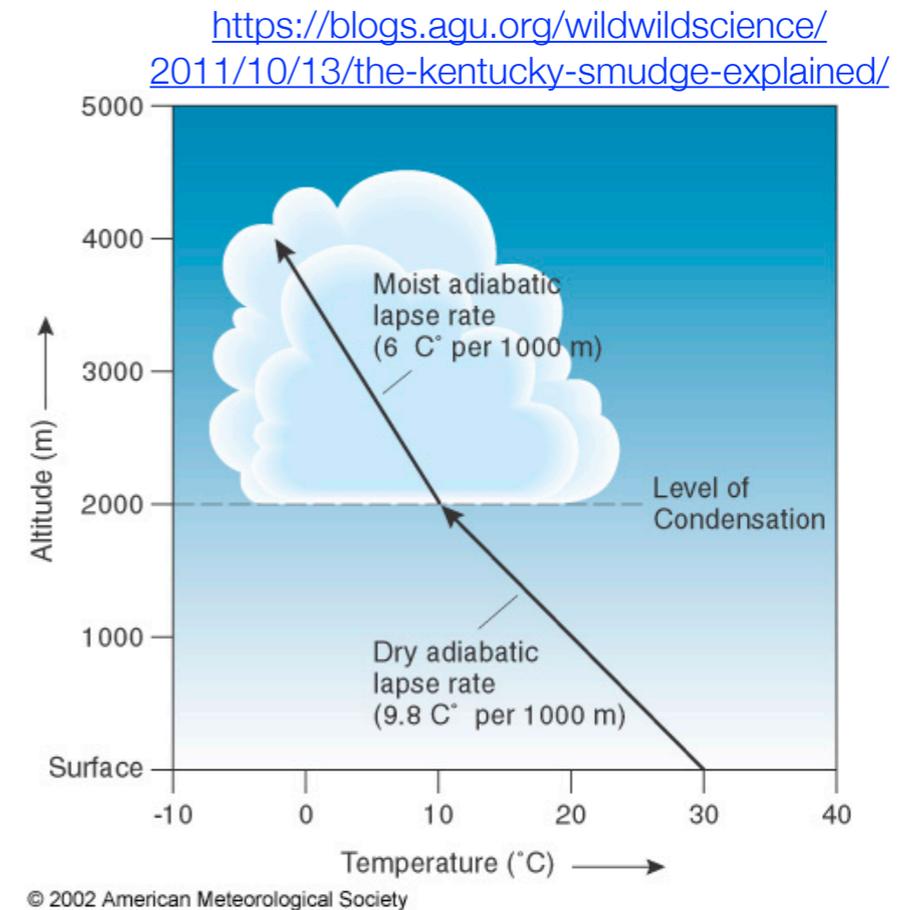


Schematic of Cloud Formation

https://www.youtube.com/watch?v=XH_M4jitiKw

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- ☉ Clouds generally form in rising air (updraft).
- ☉ As a moist parcel of air rises, it expands due to lower pressure & cools adiabatically.

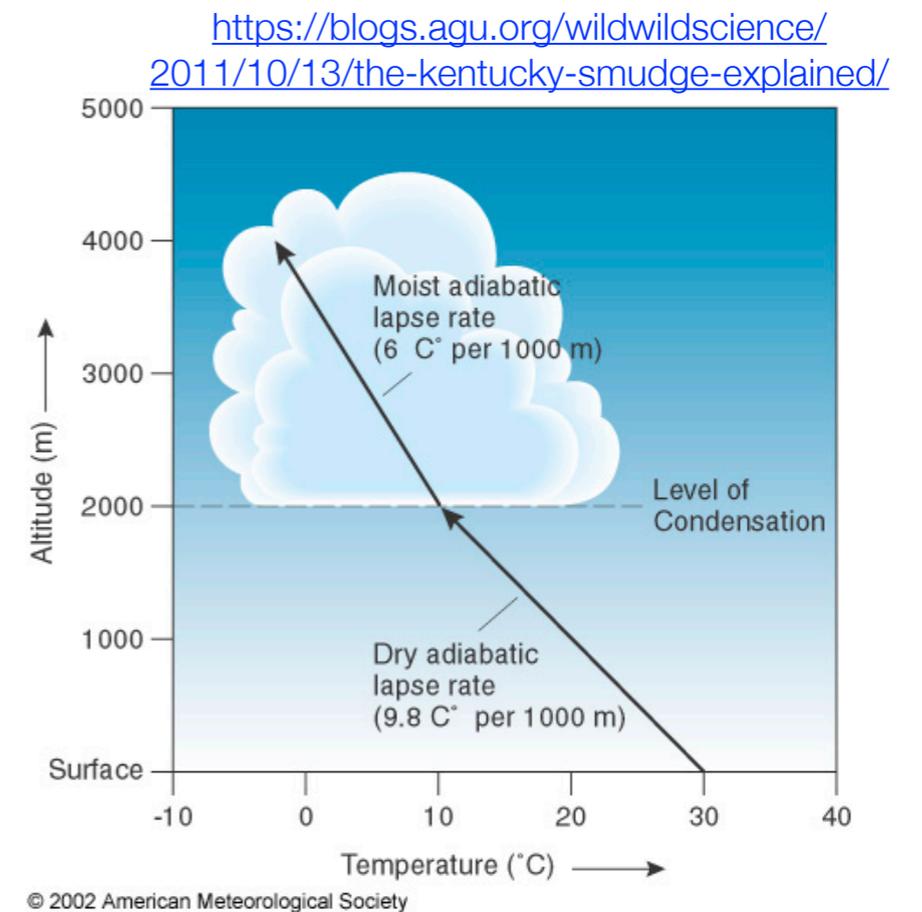


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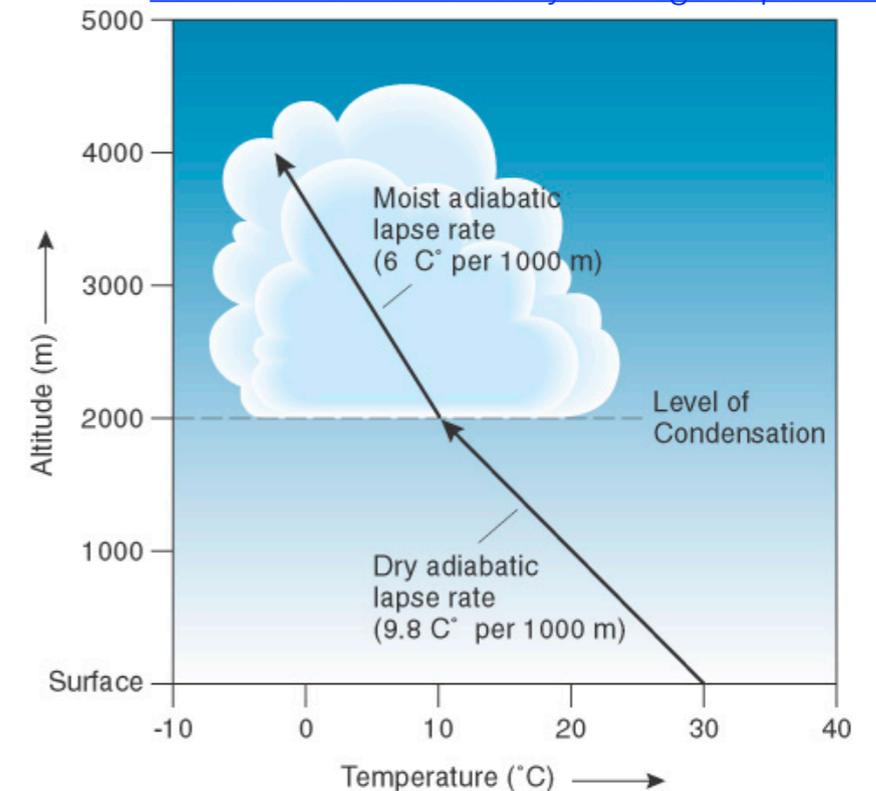
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<https://blogs.agu.org/wildwildscience/2011/10/13/the-kentucky-smudge-explained/>



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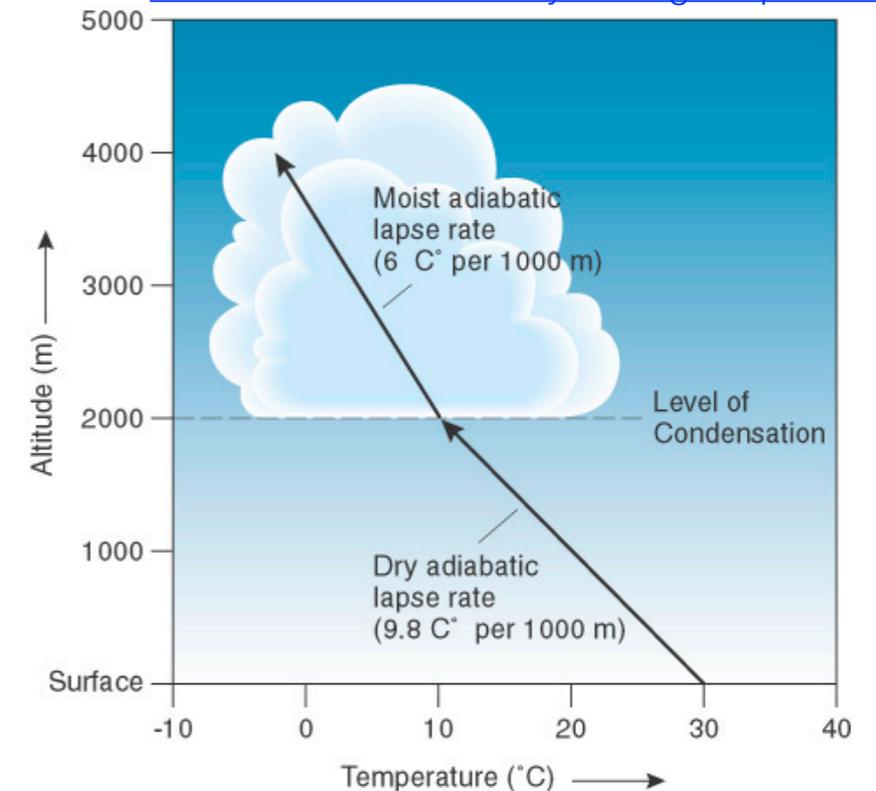
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- Further ascent & cooling causes droplets to condense as saturation humidity falls further, heat is released into the parcel, and it further rises. Positive feedback!

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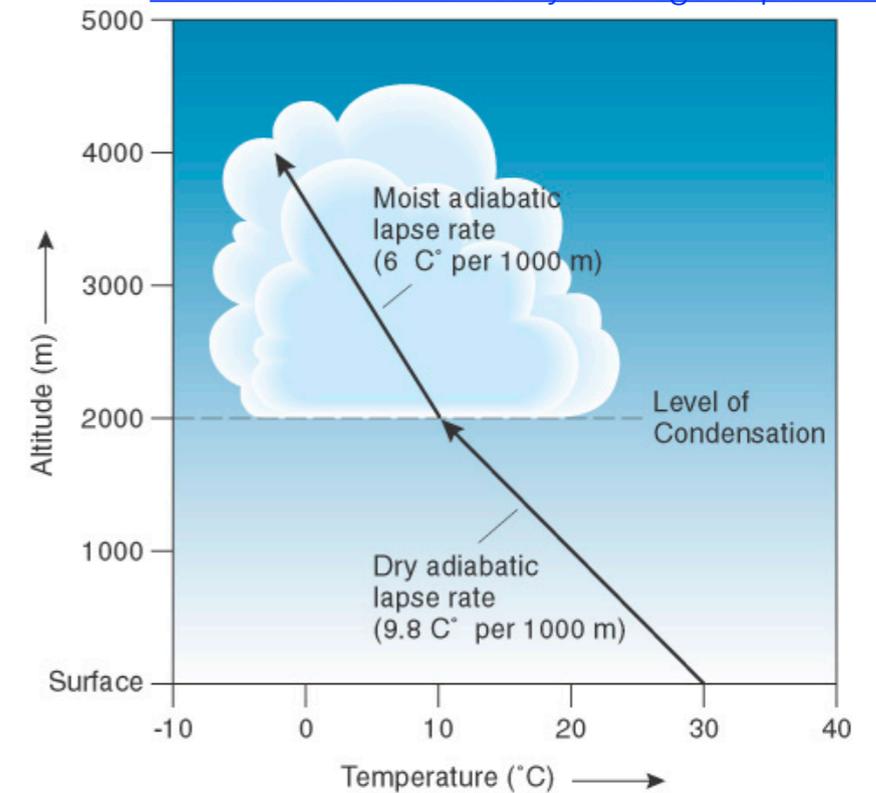
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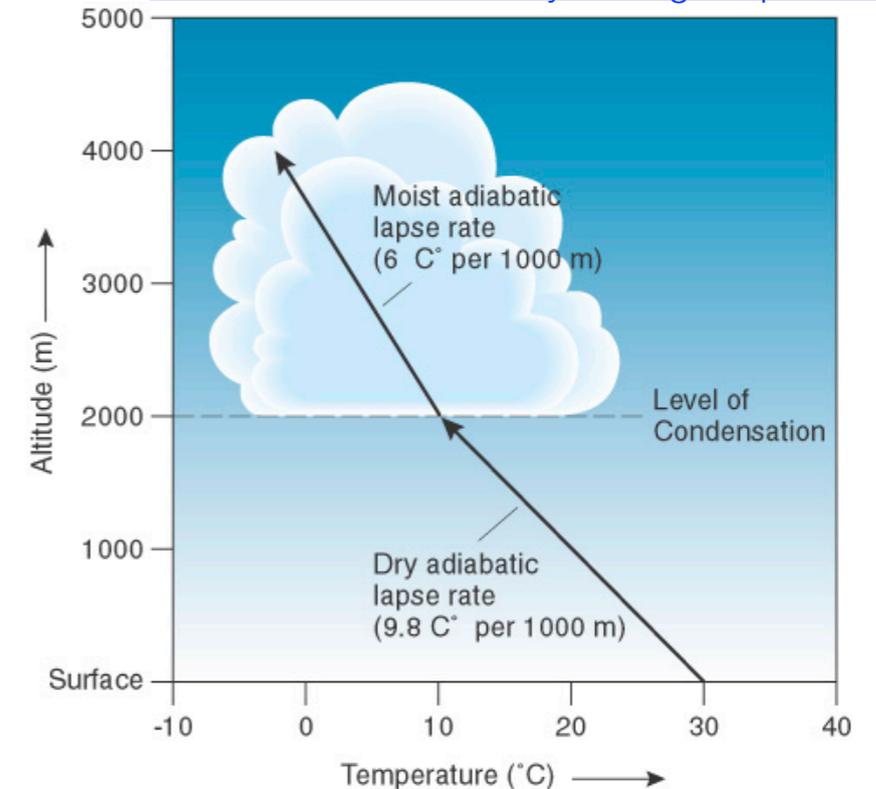
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Cloud formation by flow over topography.
(University of Wyoming)

Adiabatic heating demo: Fire Syringe

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notes section 7.2:
Moist convection and cloud formation

Moist convection and cloud formation

Moist static energy (MSE), the energy per unit mass of a moist air parcel, is conserved when the parcel is lifted adiabatically in the atmosphere,

$$MSE = c_p T(z) + Lq(z) + gz.$$

c_p : specific heat, J/(kg K); L : latent heat of condensation J/kg;

$q(z)$: specific humidity (kg water vapor/kg air)

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Parcel starts at surface, with $MSE_s = c_p T_s + Lq_s$. Initially, the rising air parcel is not saturated & there is no condensation, $q(z) = q_s$ so that the conservation may be written in terms of the *Dry Static Energy* as:

$$DSE(z) = c_p T(z) + gz = c_p T_s$$

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This leads to a solution for the temperature profile and lapse rate,

$$T(z) = T_s - gz/c_p \quad dT/dz = -g/c_p = -9.8 \text{ K/km}$$

Moist convection and cloud formation

The parcel keeps rising & cooling, until the saturation moisture is smaller than the parcel's moisture & condensation occurs, $q(z) = q^*(T(z), p(z))$. The conservation law is $c_p T(z) + Lq^*(T(z), p(z)) + gz = MSE_s$ which may be solved graphically for $T(z)$.

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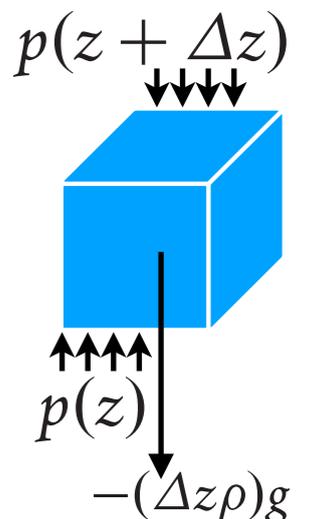
$$MSE_s = c_p T(z) + L \min(q_s, q^*(T(z), p(z))) + gz$$

To solve, need to find $p(z)$ from the vertical momentum (hydrostatic) balance for an air parcel: $dp/dz = -\rho g$. Using $\rho = p/(RT)$ this becomes $dp/dz = -pg/RT$, or $d \ln p = -(g/RT) dz$.

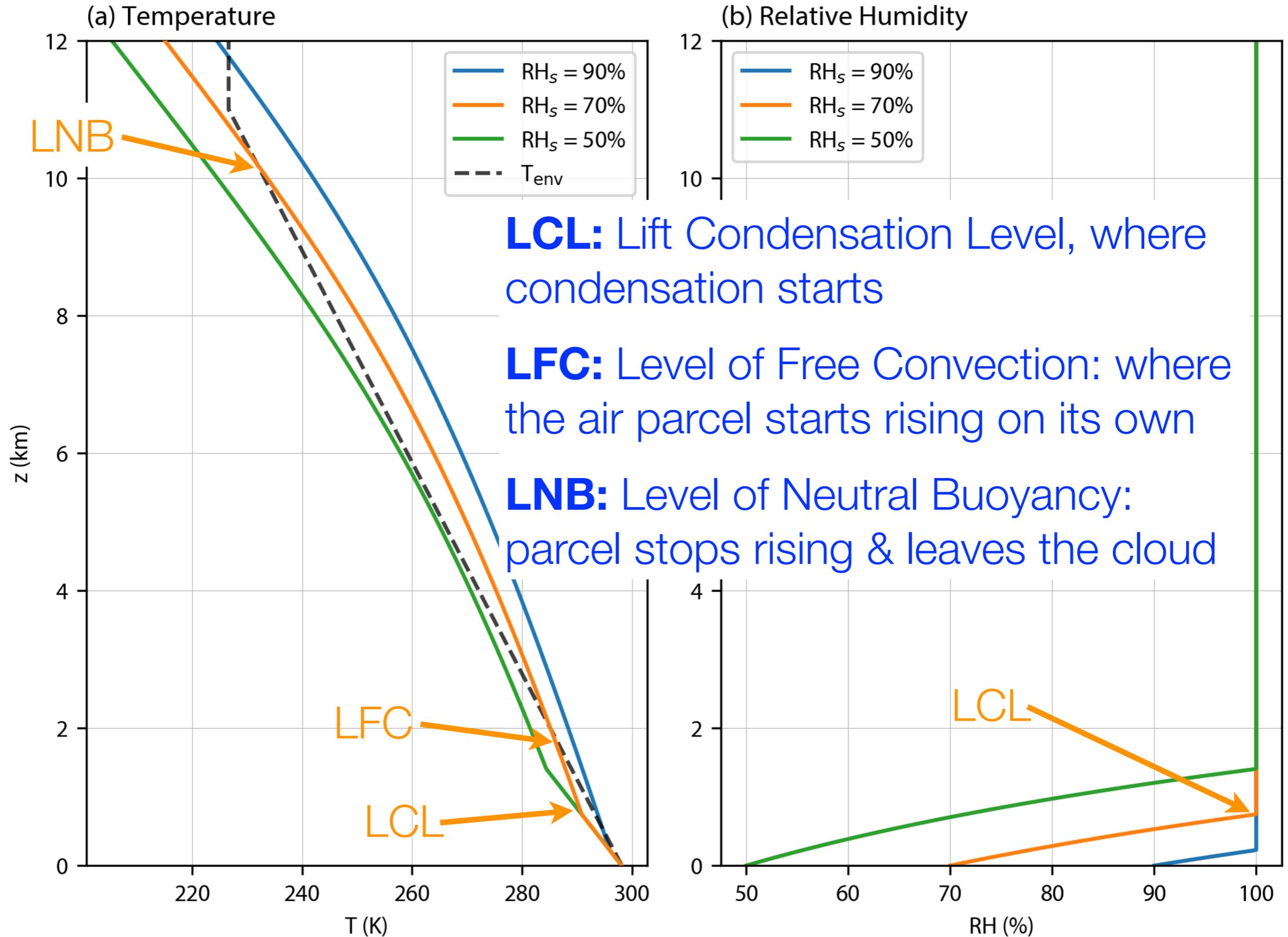
Integrating, we find $\ln p(z) - \ln p_s = -gz/R\bar{T}$,

so that $p(z) = p_s e^{-gz/(R\bar{T})}$

➔ pressure is exponential in height.



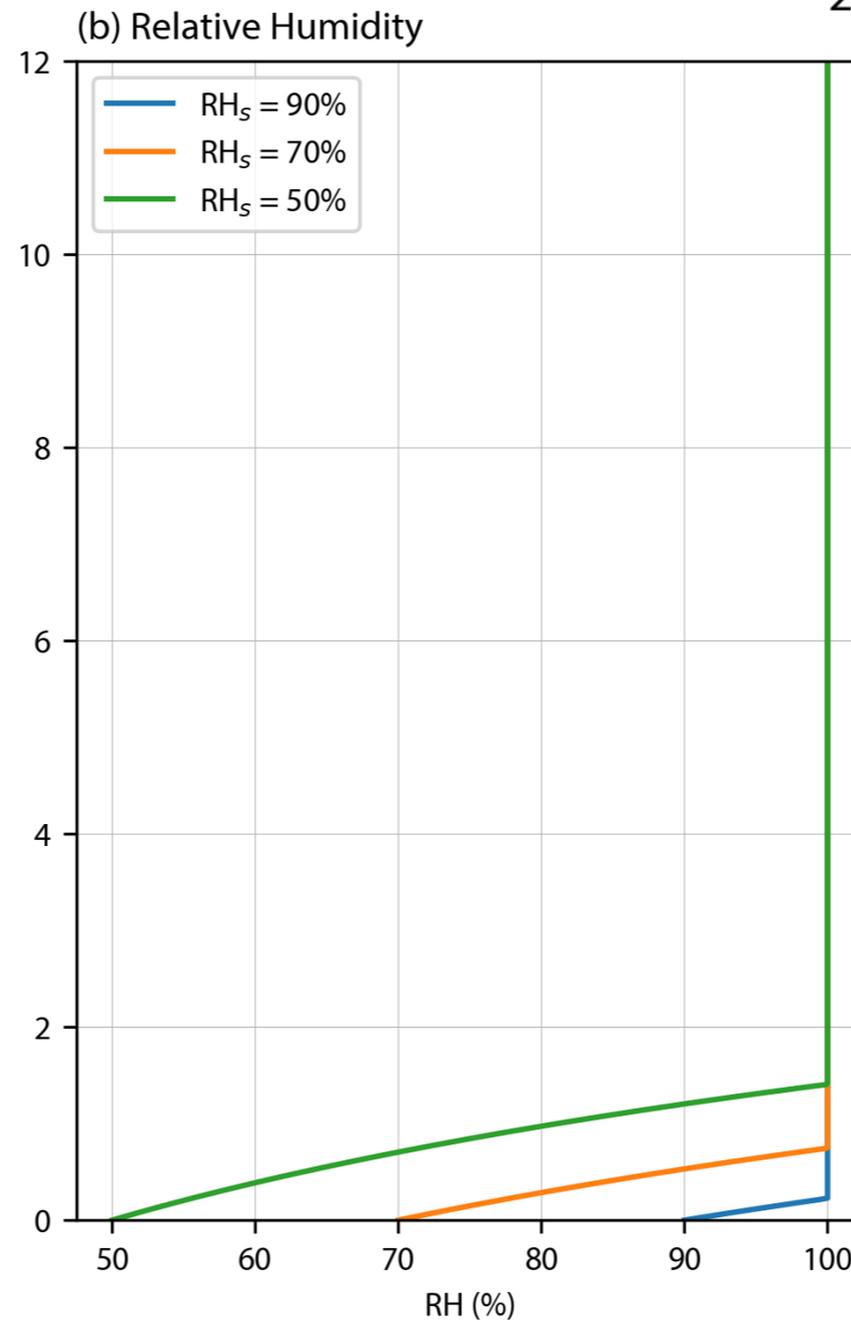
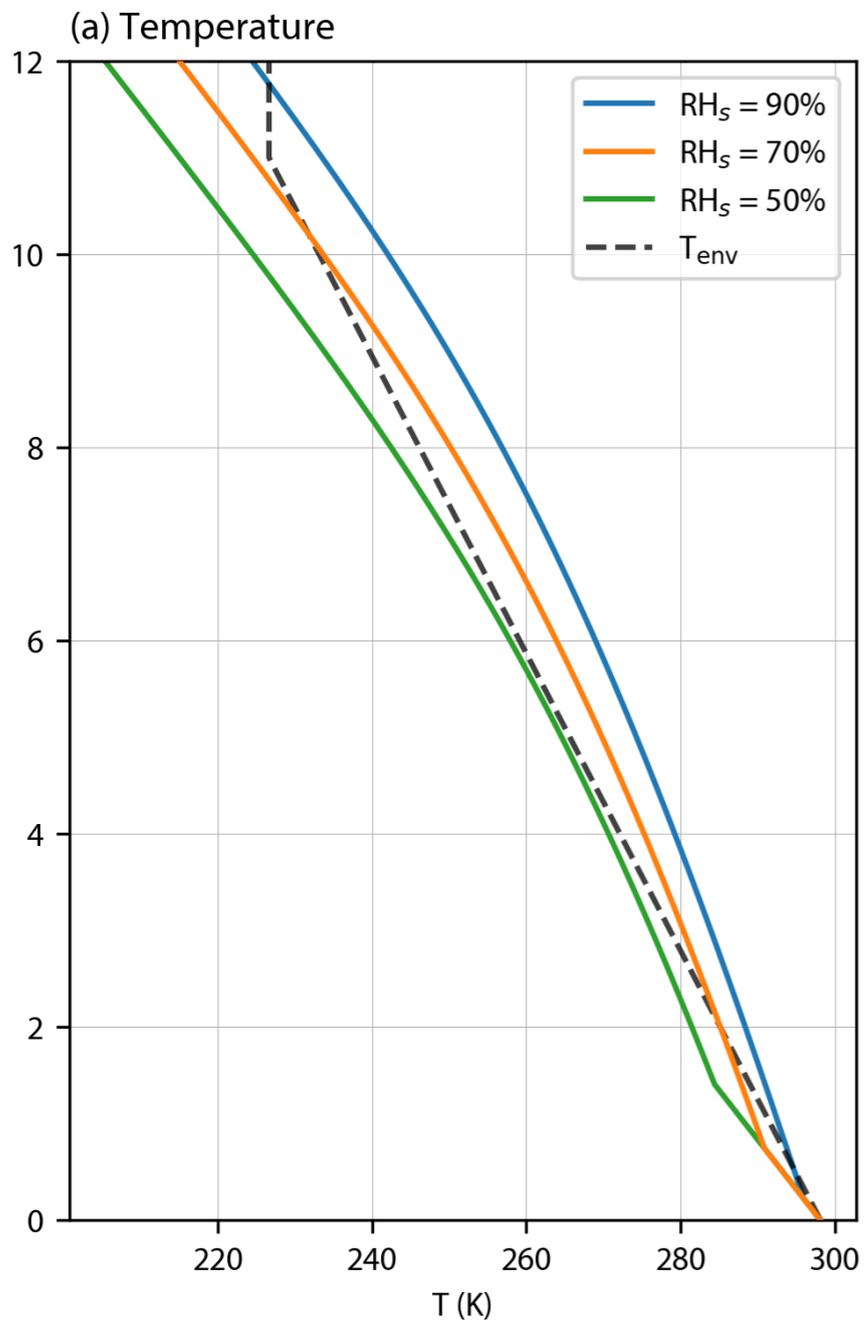
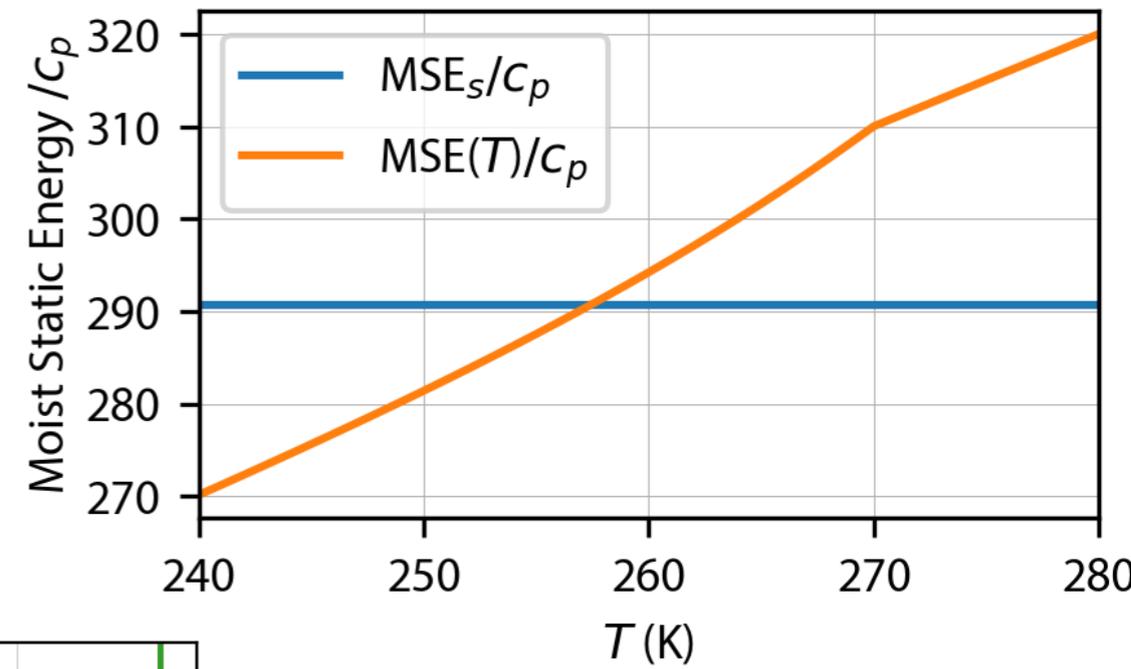
Moist Convection: LCL and LFC



workshop #2
Convection and cloud formation

workshop #2

Convection and cloud formation



- Lift Condensation Level
- Level of Free Convection
- Level of Neutral Buoyancy
- Stable vs unstable conditions

A major source of cloud uncertainty: turbulent entrainment



<https://www.volcanocafe.org/the-wandering-earth-mantle-in-motion/>

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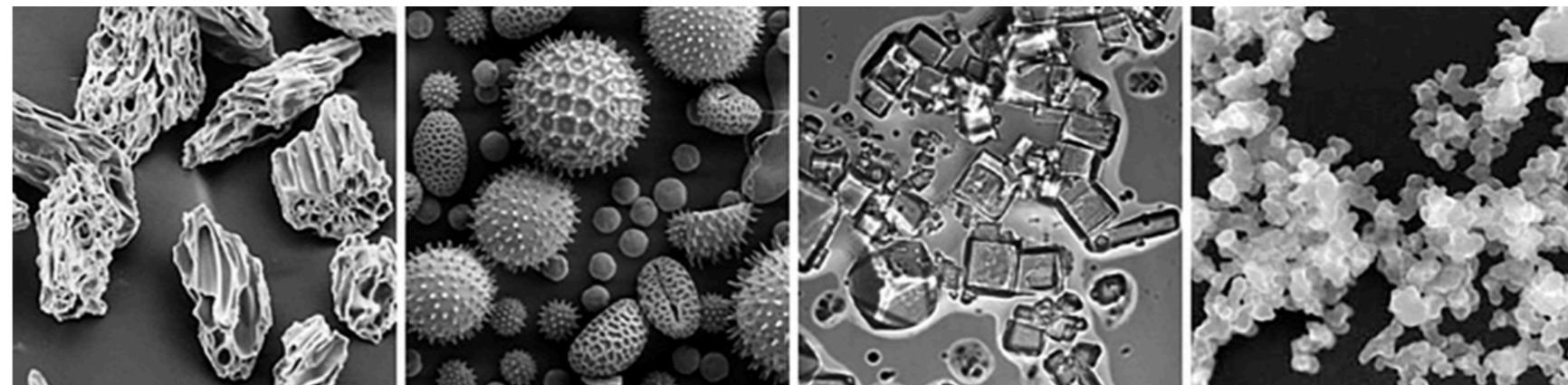
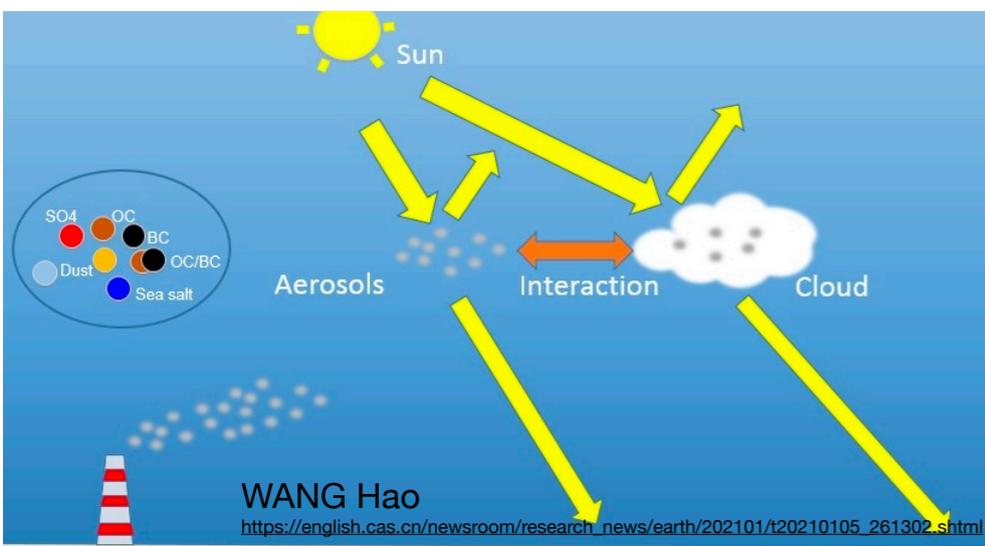


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Another major source of cloud uncertainty: Cloud Microphysics

Microphysics

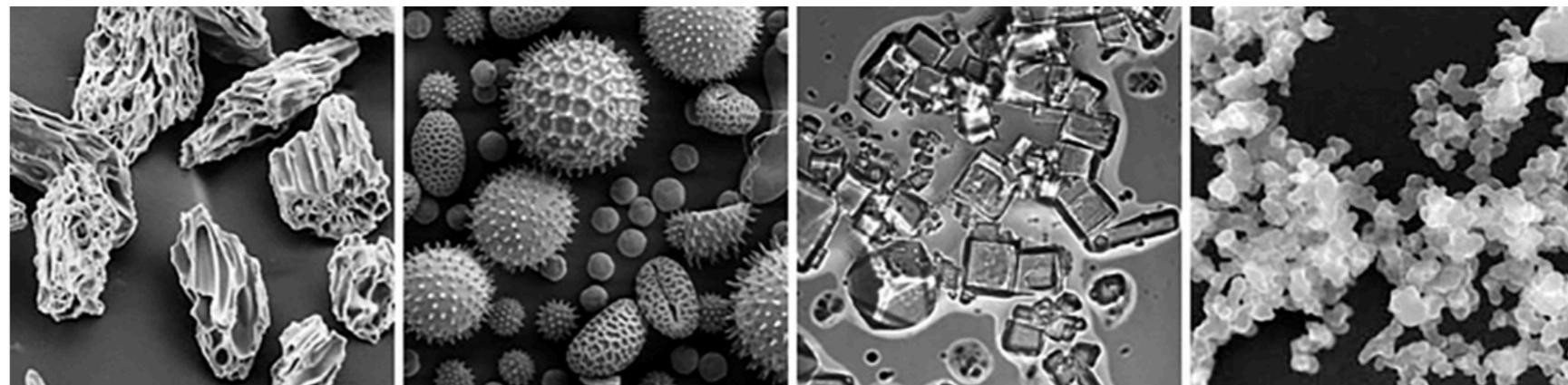
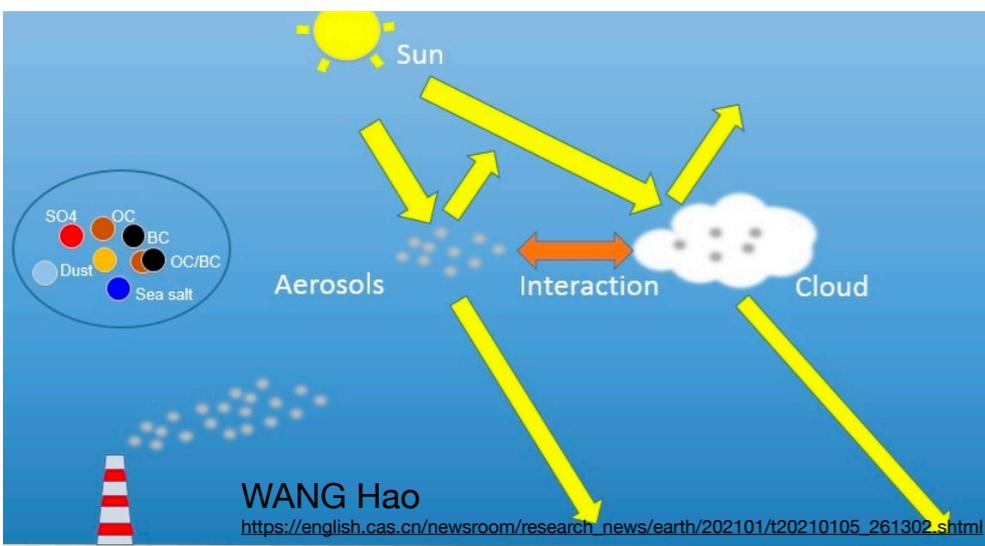
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Scanning electron microscope images (not same scale) of aerosols. volcanic ash, pollen, sea salt, and soot. NASA, from USGS, UMBC, (Chere Petty), Arizona State Univ (Peter Buseck)
<https://scied.ucar.edu/learning-zone/air-quality/aerosols>

Microphysics

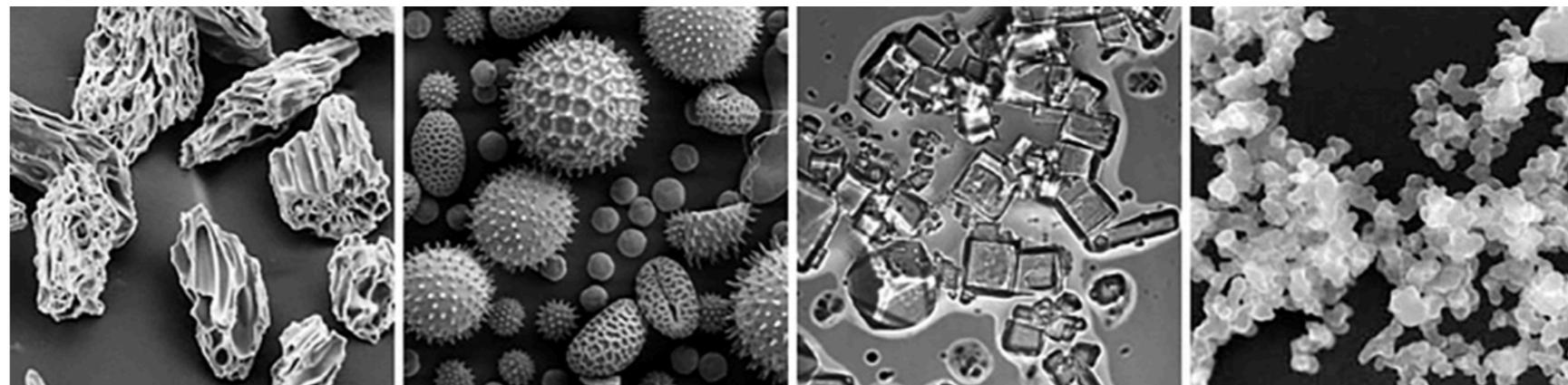
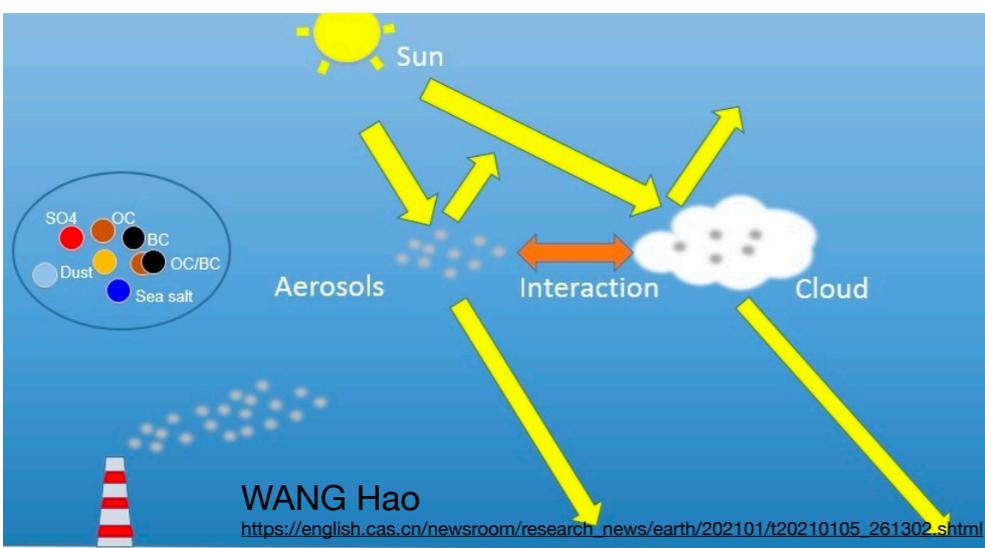
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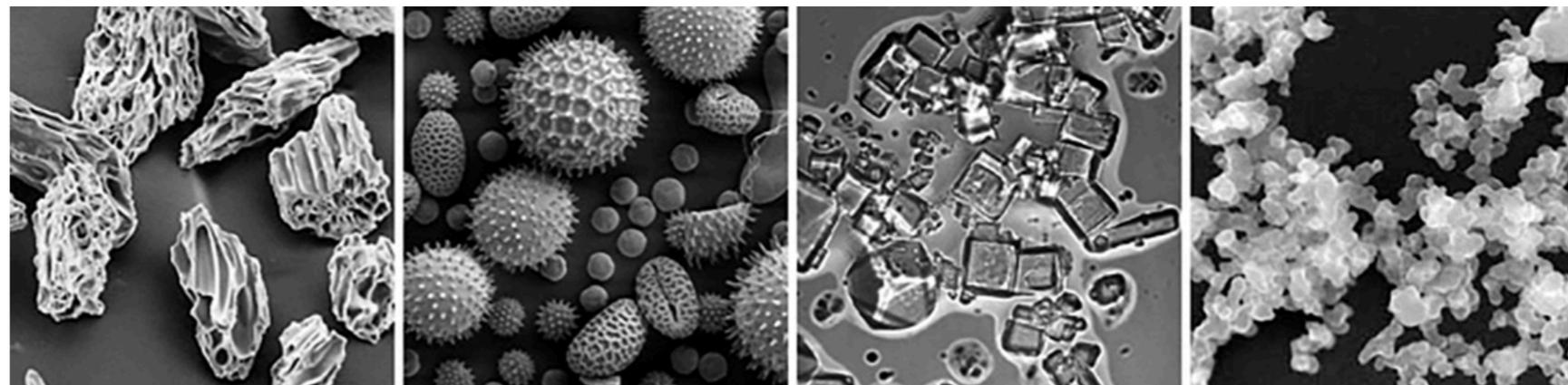
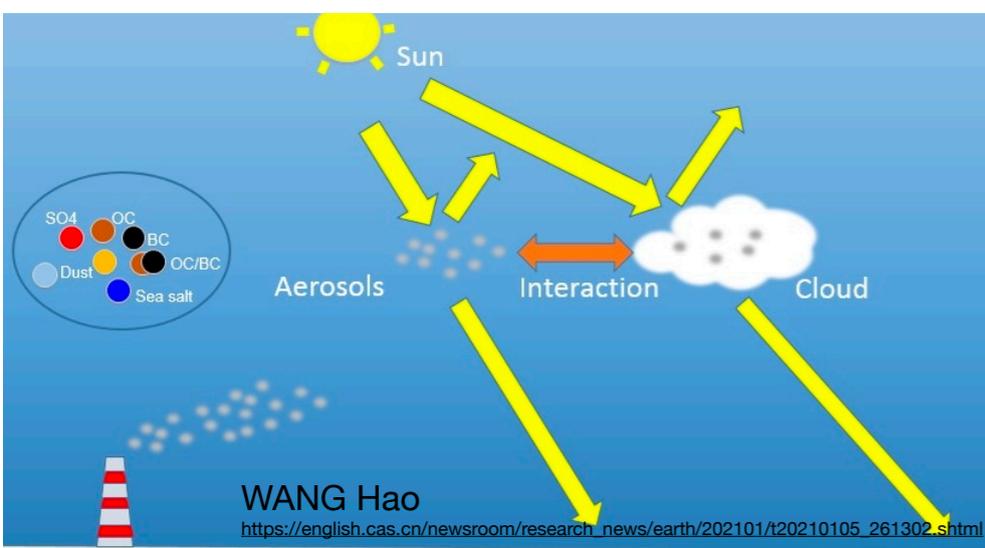
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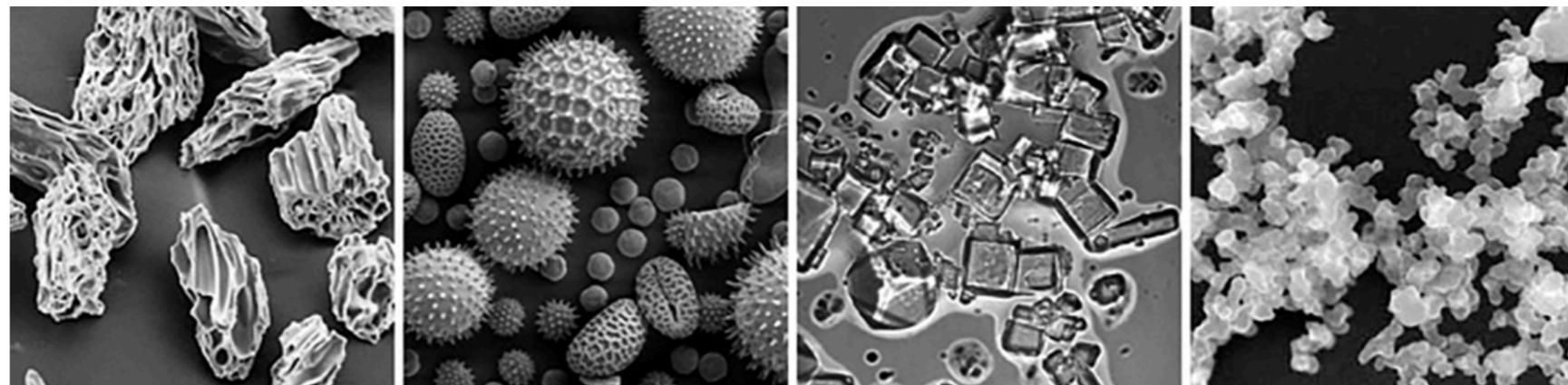
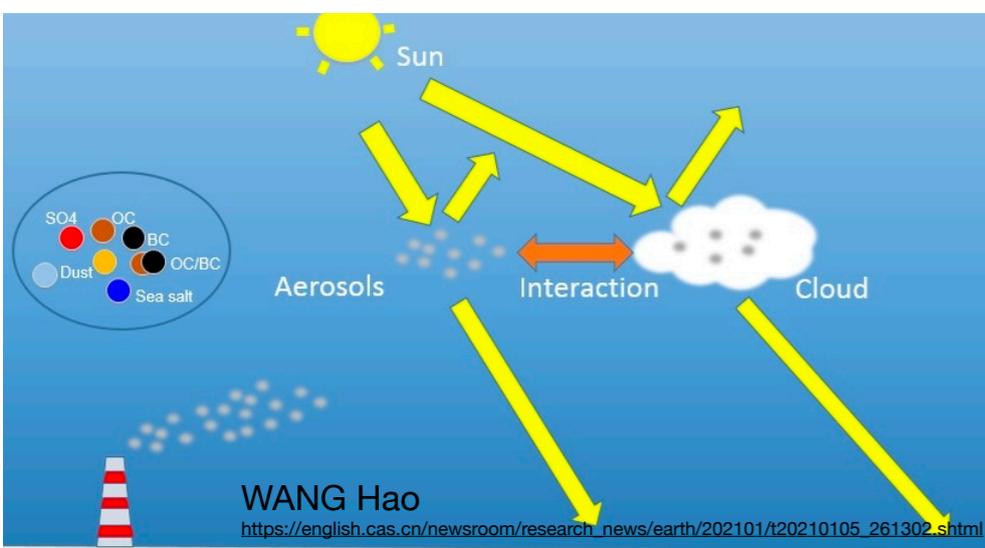
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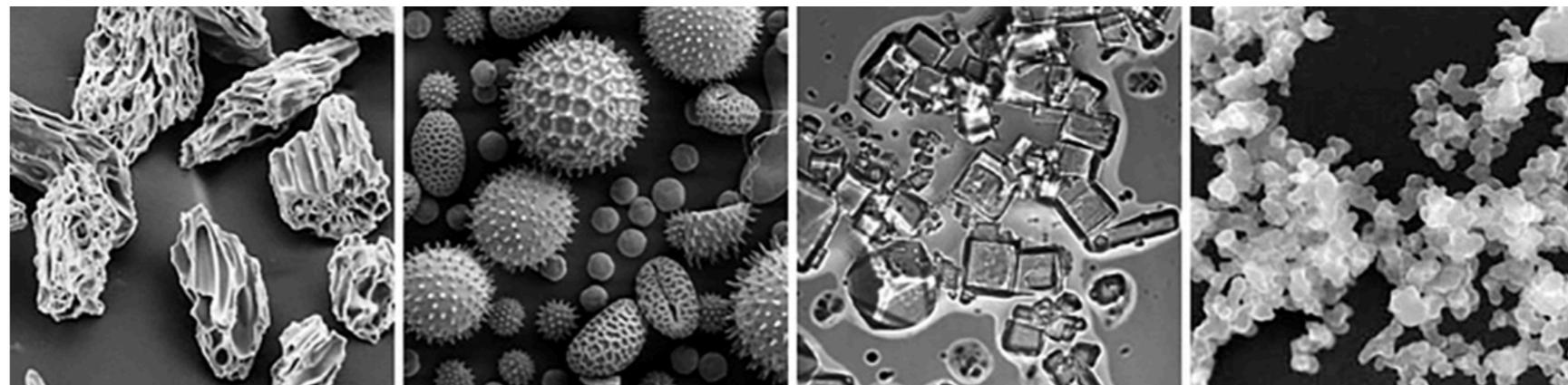
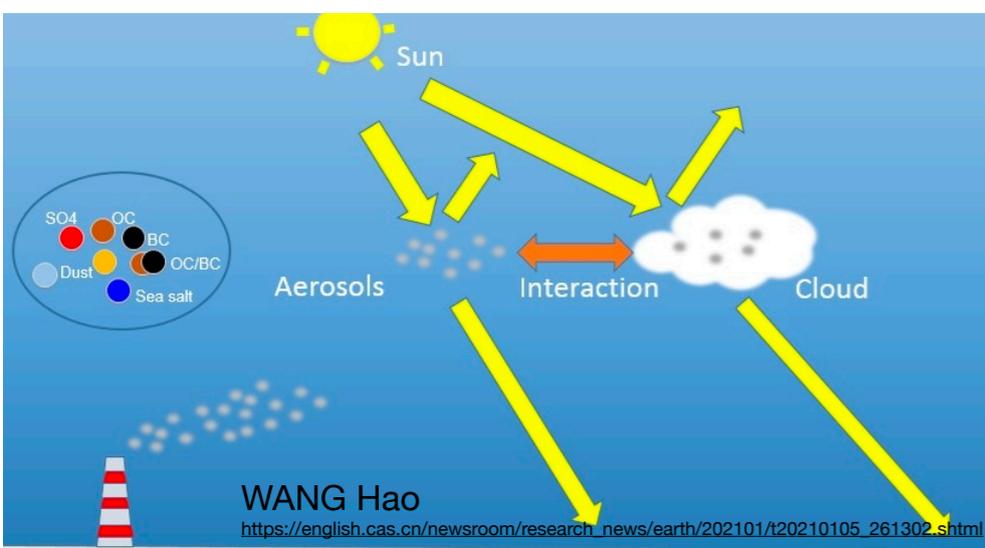
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- ☉ Further uncertainty due to radiative effects: Aerosols absorb/ reflect sunlight (**direct effects**) & affect the number/size of cloud droplets/ice particles, and therefore the radiative effects of clouds (the **indirect effect**).



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Life cycle of a cloud

- Clouds dissipation occurs both via droplets/ ice crystals falling toward the ground and via the continuous evaporation of droplets.
- Timelapse of clouds (2 hours of footage in 2 minutes):



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Note the complex evolution, and the small scale relative to climate model's grid scale, leading to the large uncertainty

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Cloud types: high vs low, water vs ice

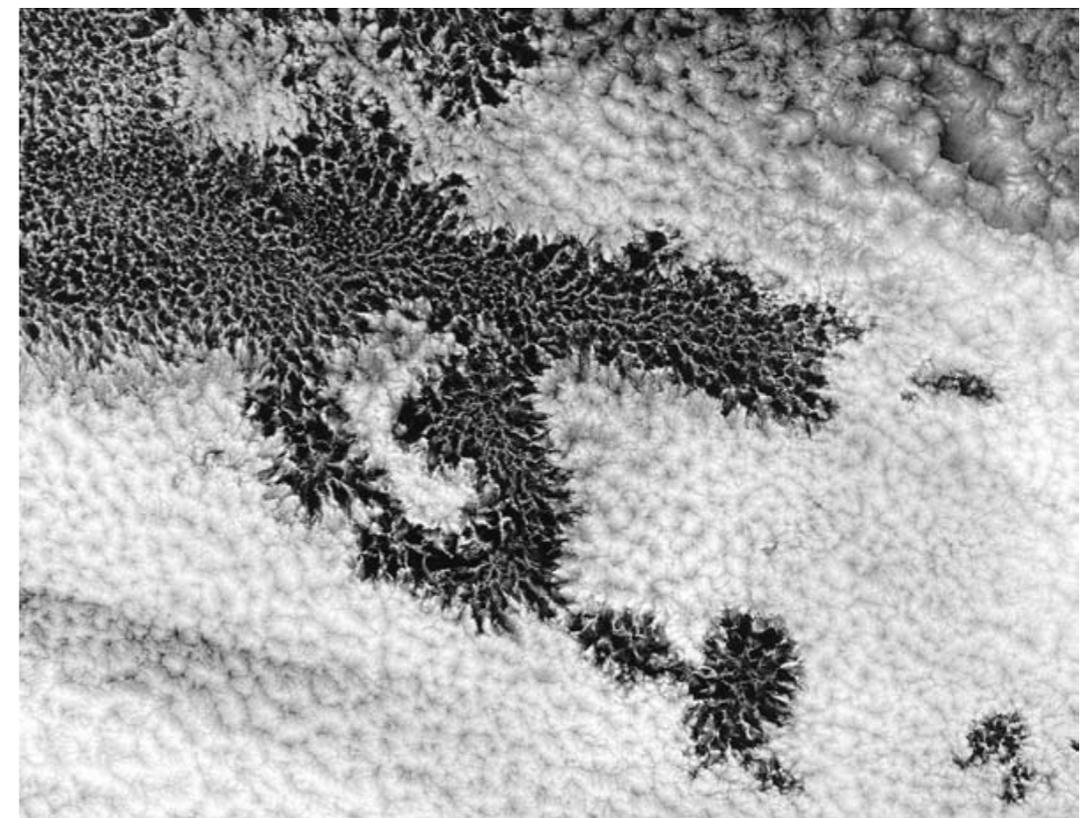
Low Cloud: Stratocumulus

- Cover broad regions over the subtropical oceans.
- Characterized by lines, waves, and cellular structures.
- Radiative cooling from the cloud tops drives mixing with surface air that replenishes the liquid water in these clouds.
- Can often be seen out of an airplane window while flying.
- Large SW albedo, strong cooling effects on climate.



Stratocumulus clouds from a plane

http://www.pilotfriend.com/training/flight_training/met/clouds.htm



Cellular convective structures

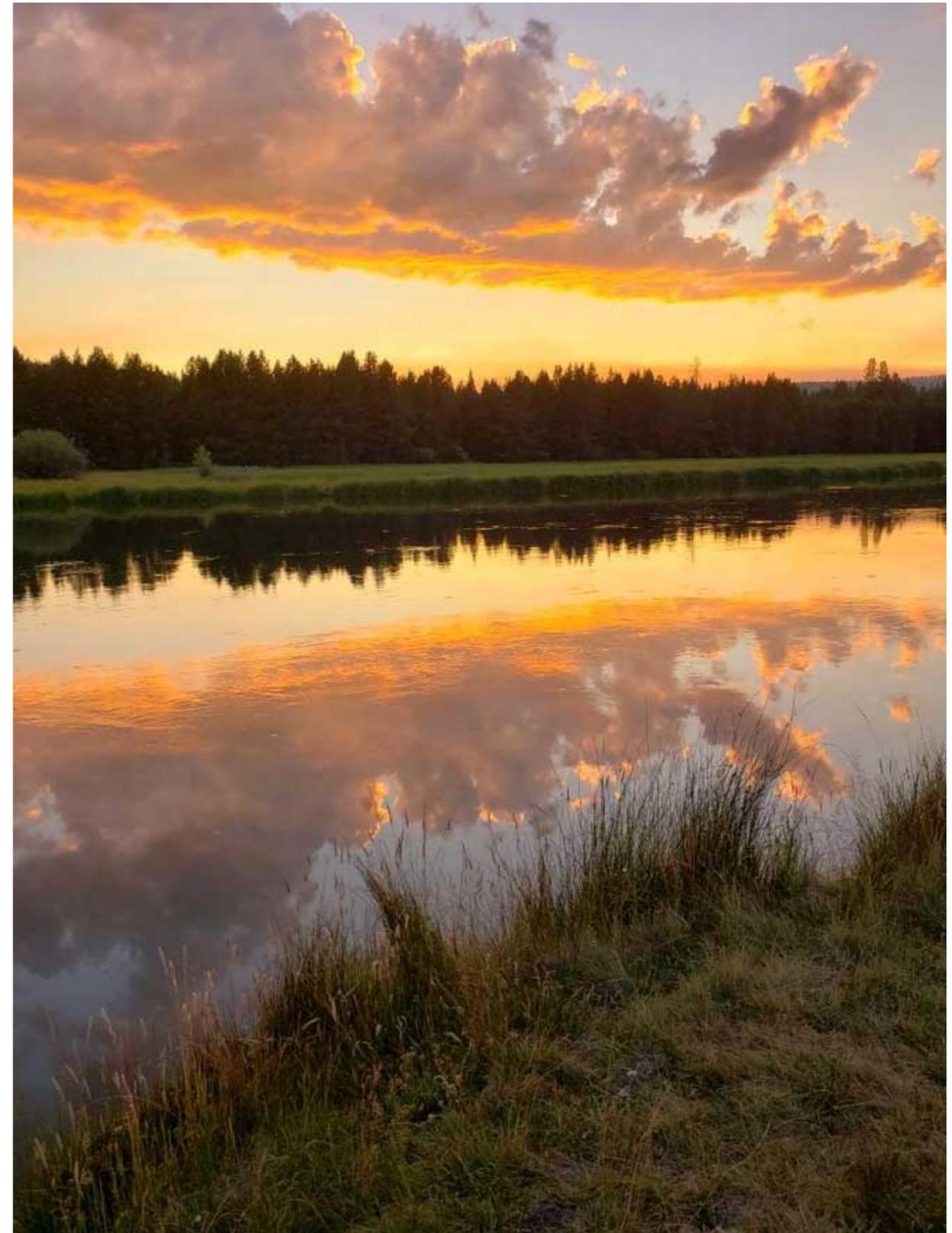
<https://visibleearth.nasa.gov/images/98570/clouds-in-eastern-south-pacific-ocean?size=small>

Low Cloud: Shallow cumulus

- The most familiar type of cloud.
- Low level clouds that do not precipitate.
- Small size and thus small radiative impact, comparatively less important to climate.



<https://cloudappreciationsociety.org/gallery/photo/photo-n-357365>



<https://cloudappreciationsociety.org/gallery/photo/photo-n-359960/>

High Cloud: Deep Cumulus

- Strongly convecting updrafts that may reach up to the tropopause (9–17 km above the surface).
- Often characterized by a flat anvil-like top.
- Most common in tropical regions.
- **Cover a very small fraction of tropical areas** but are important for setting moisture and temperature profiles in the tropics.



<https://cloudappreciationsociety.org/gallery/photo/photo-n-357970>



<https://cloudappreciationsociety.org/gallery/photo/photo-n-358495>

Deep cumulus clouds with associated anvil tops.

High Cloud: Cirrus clouds

- Thin wispy clouds formed of ice crystals.
- Very high in altitude (4–20 km).
- Can form at the outflow of deep cumulus clouds or in warm fronts.
- **Large LW emissivity, strong warming effect on climate.**



<https://cloudappreciationsociety.org/gallery/photo/photo-n-359475>



<https://cloudappreciationsociety.org/gallery/photo/photo-n-285941>

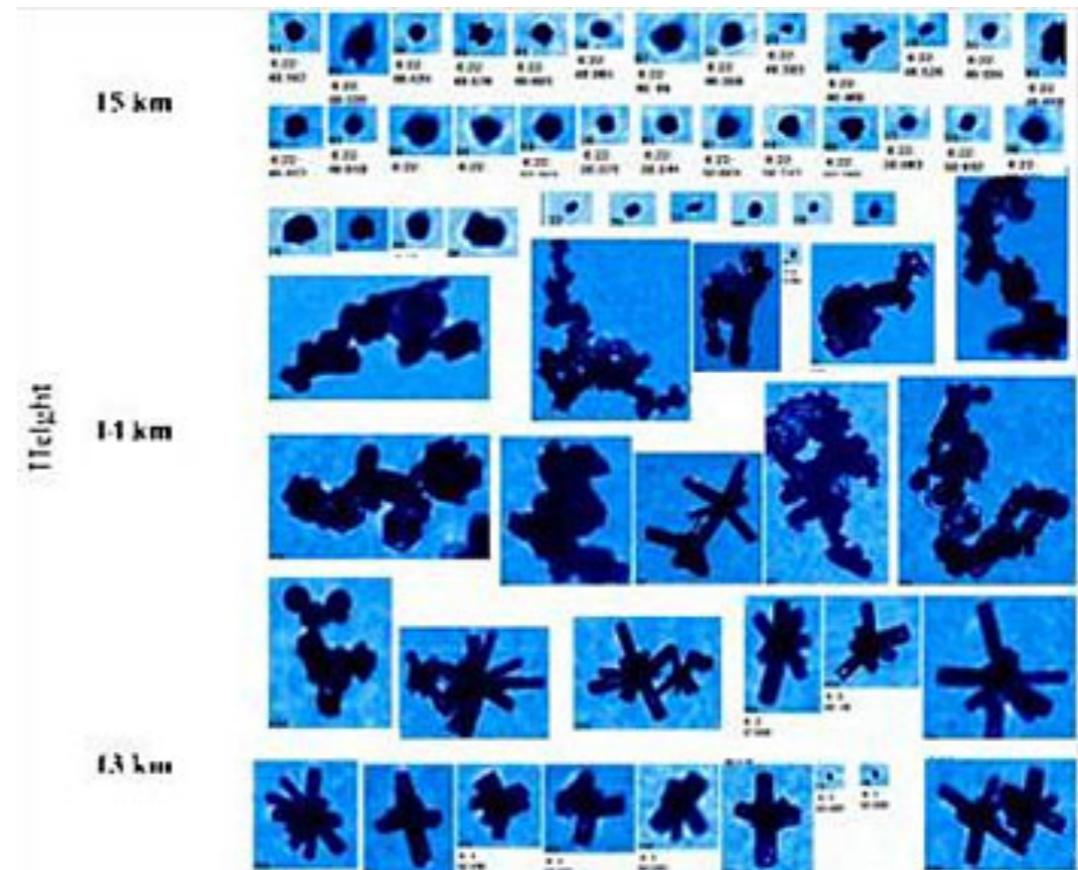
Water vs Ice Clouds

- ❖ Smaller particles yield a larger aggregate cross-section area of cloud particles for the same total water content.
- ❖ Ice clouds tend to be composed of larger particles. Their high altitude and cold environment also lead to small water content.
- ❖ Hence, they are not as good as low water clouds at scattering SW radiation.
- ❖ However, they are still, to a very good approximation, black bodies in the LW.
- ❖ ➔ **Thin ice cirrus clouds, are effective at warming: a strong longwave but little shortwave CRE.**



Sky filled with cirrus clouds.
(Wikipedia commons)

https://en.wikipedia.org/wiki/Cirrus_cloud#/media/File:CirrusField-color.jpg



Cloud Ice Particles (ARM) (Lawson et al 2006)

<https://journals.ametsoc.org/jamc/article/45/11/1505/12640/Microphysical-and-Optical-Properties-of>

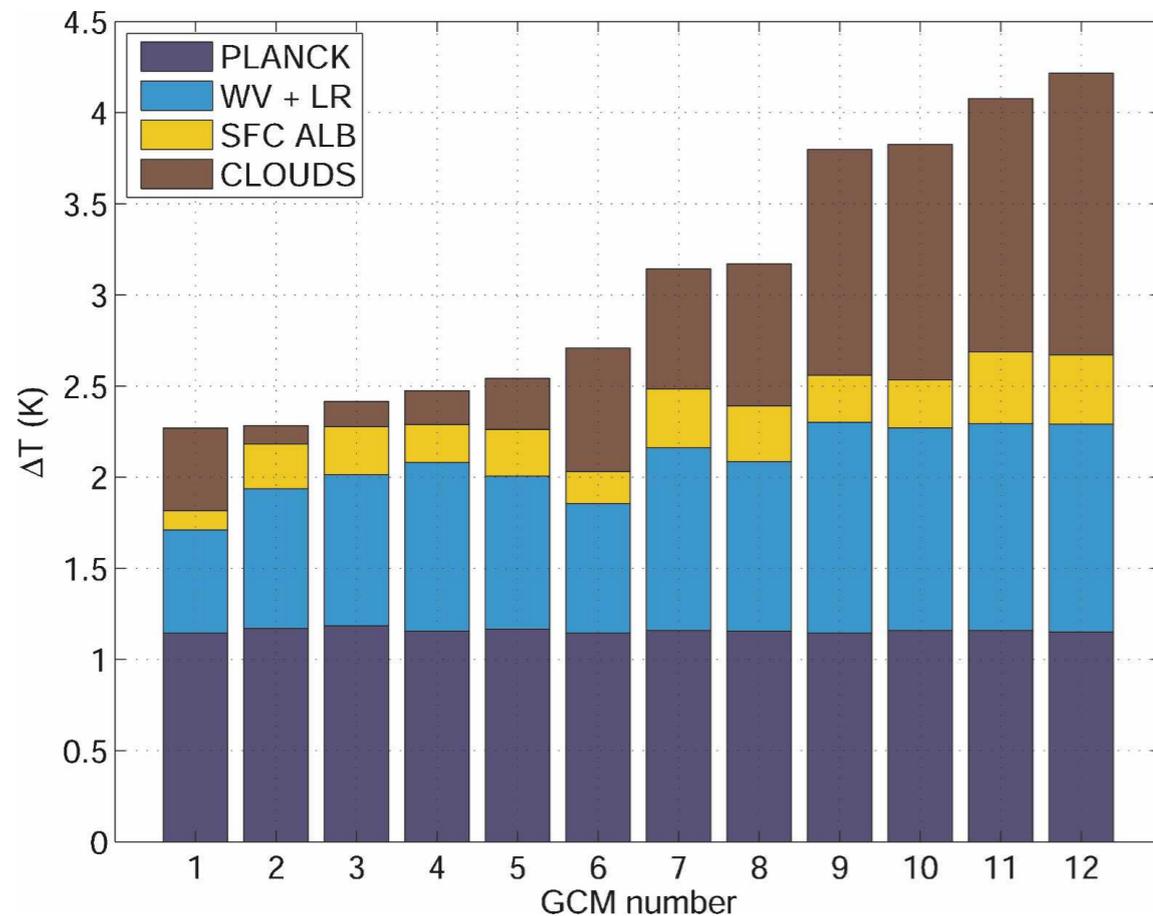
Clouds and climate uncertainty

CRE: Primary Source of Climate Uncertainty

- ★ Cloud feedback on higher CO₂ is generally estimated to be positive, although highly uncertain in magnitude. This uncertainty is partly because CRE is composed of two large competing effects: LW warming/SW cooling.

CRE: Primary Source of Climate Uncertainty

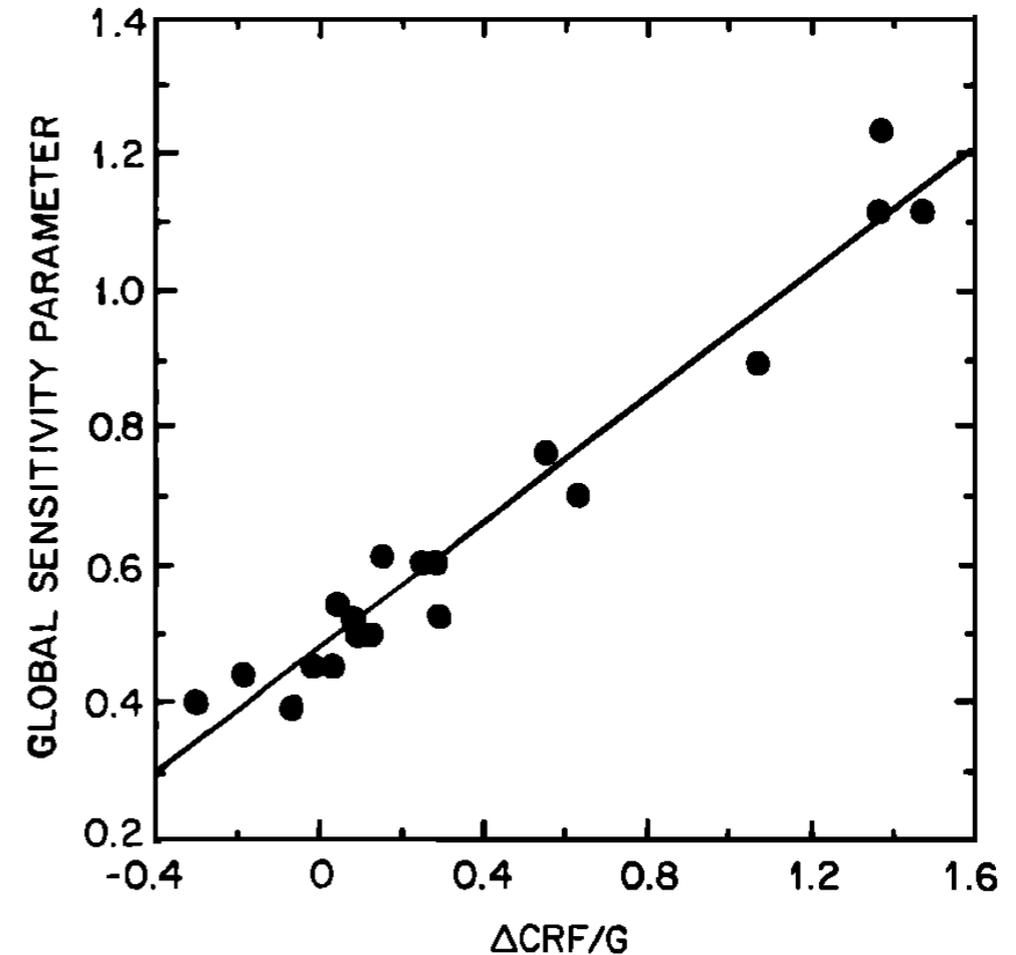
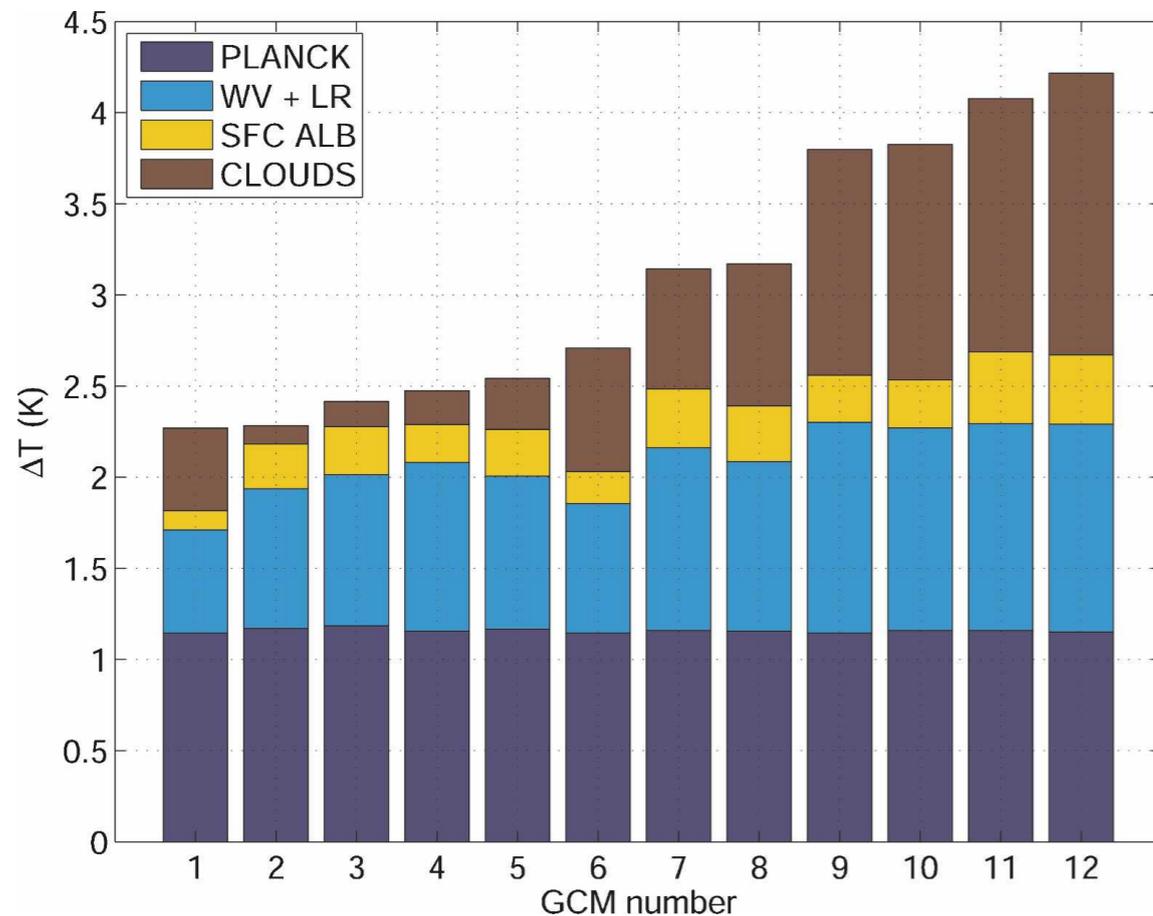
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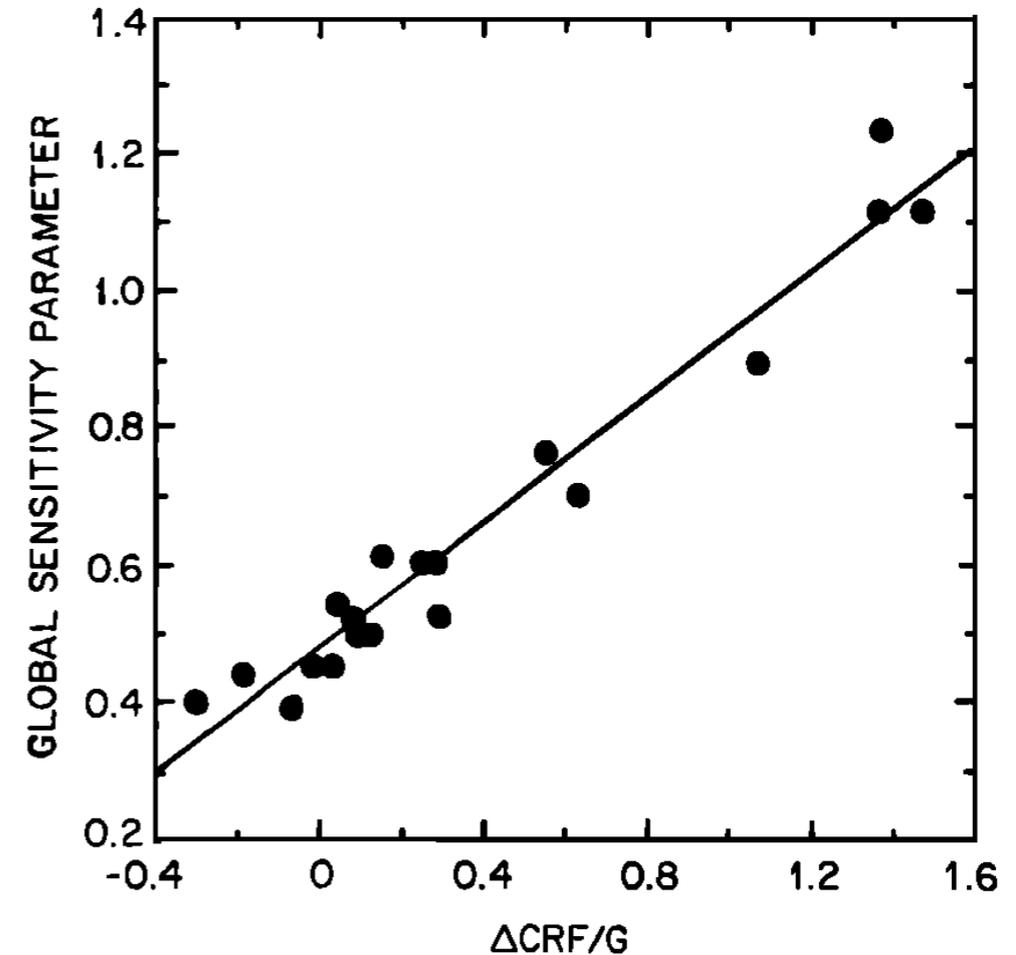
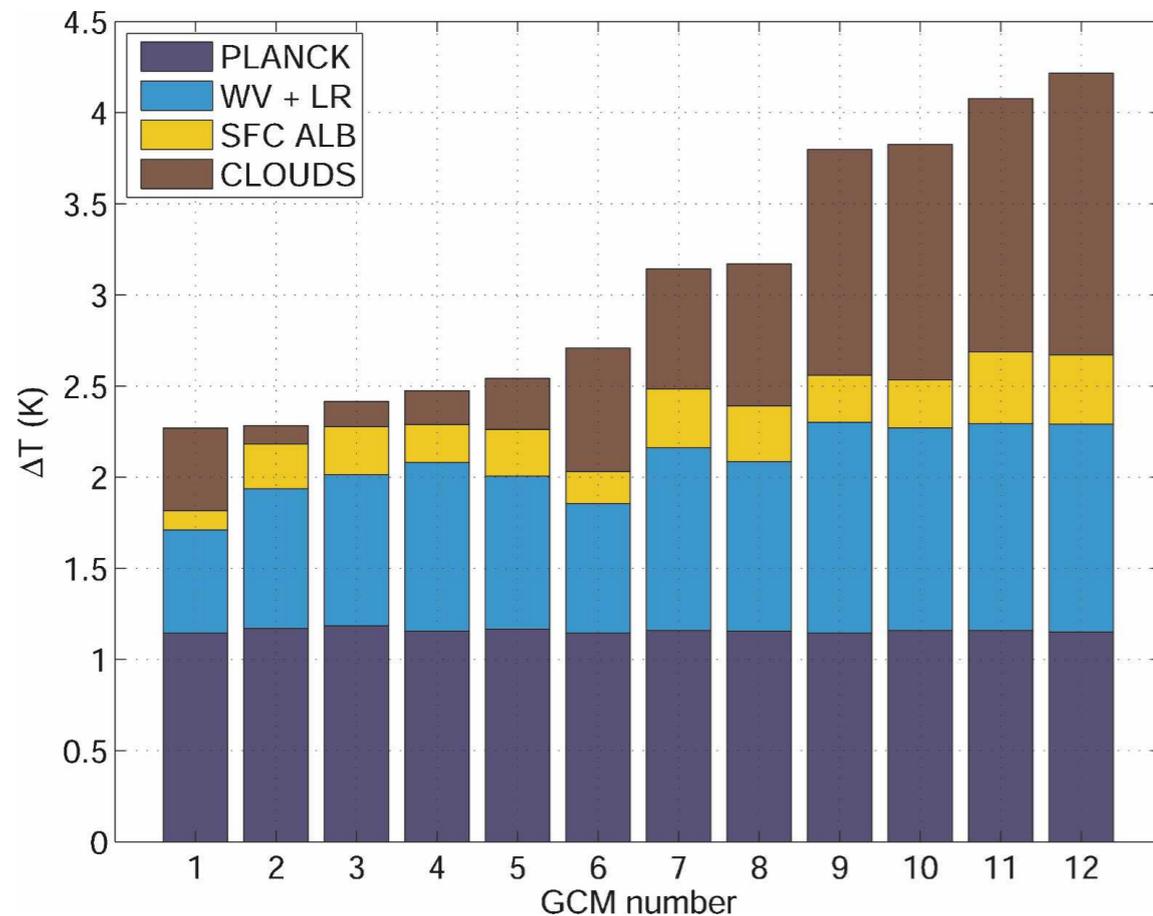


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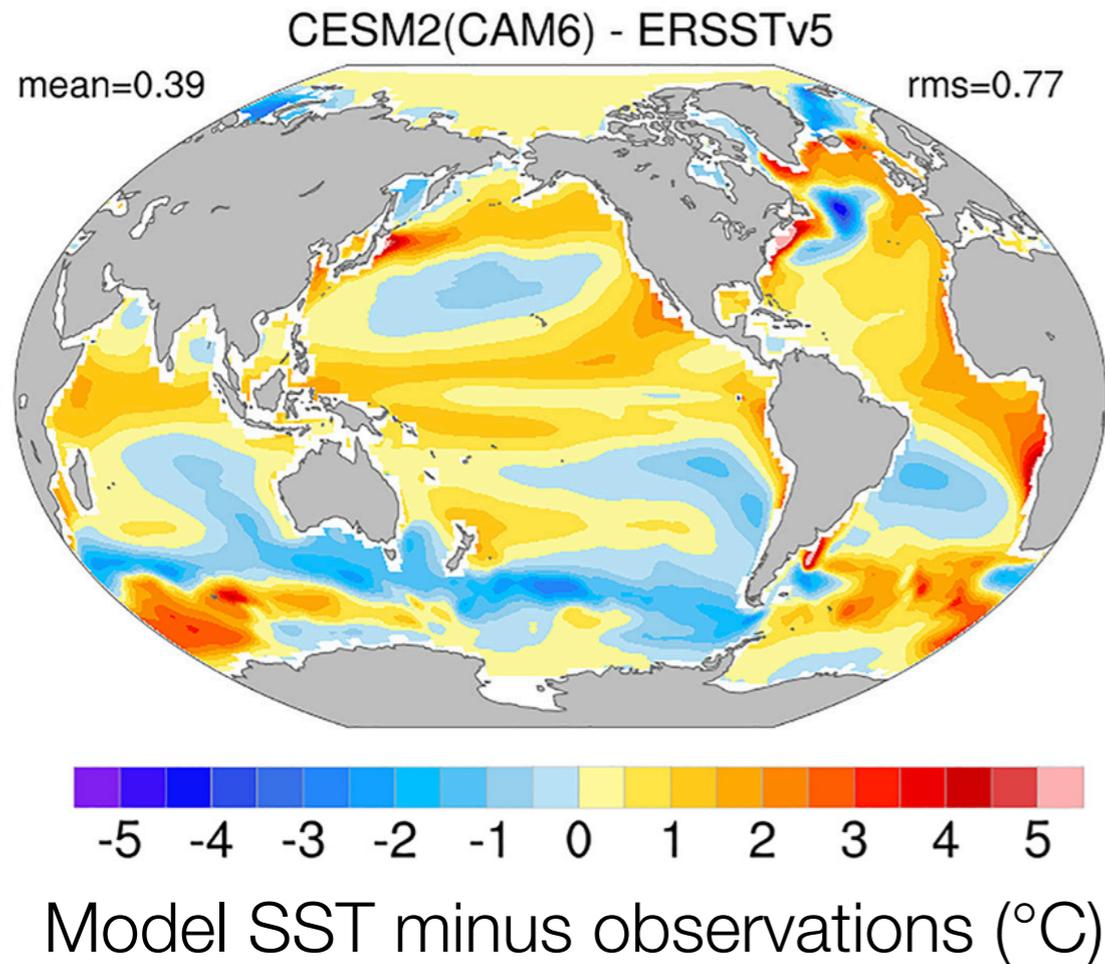


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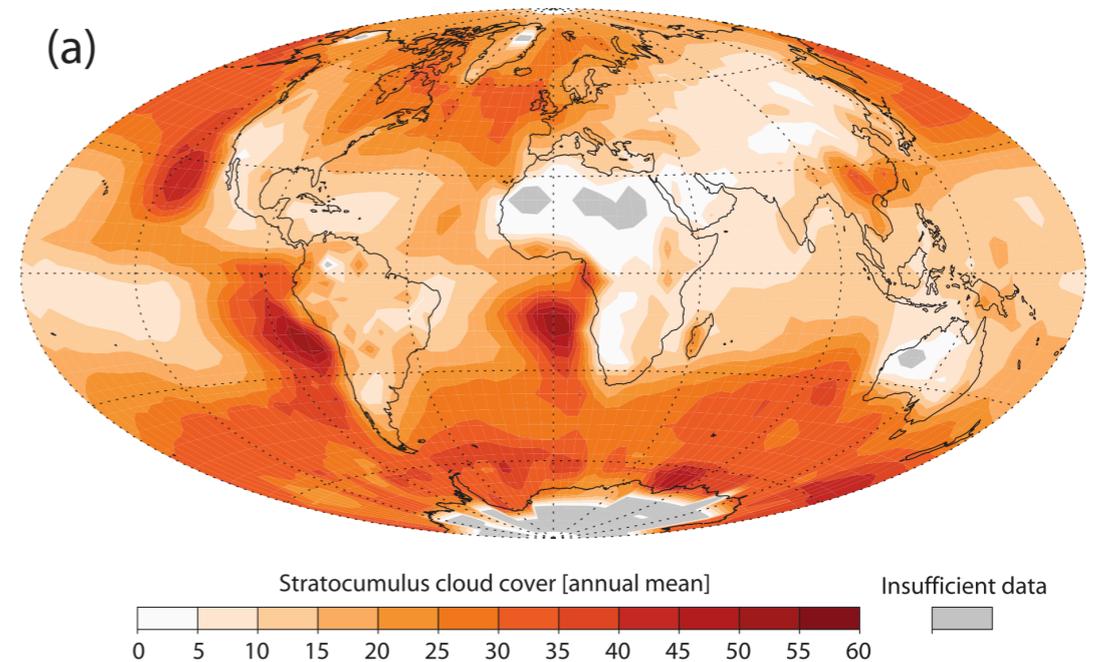
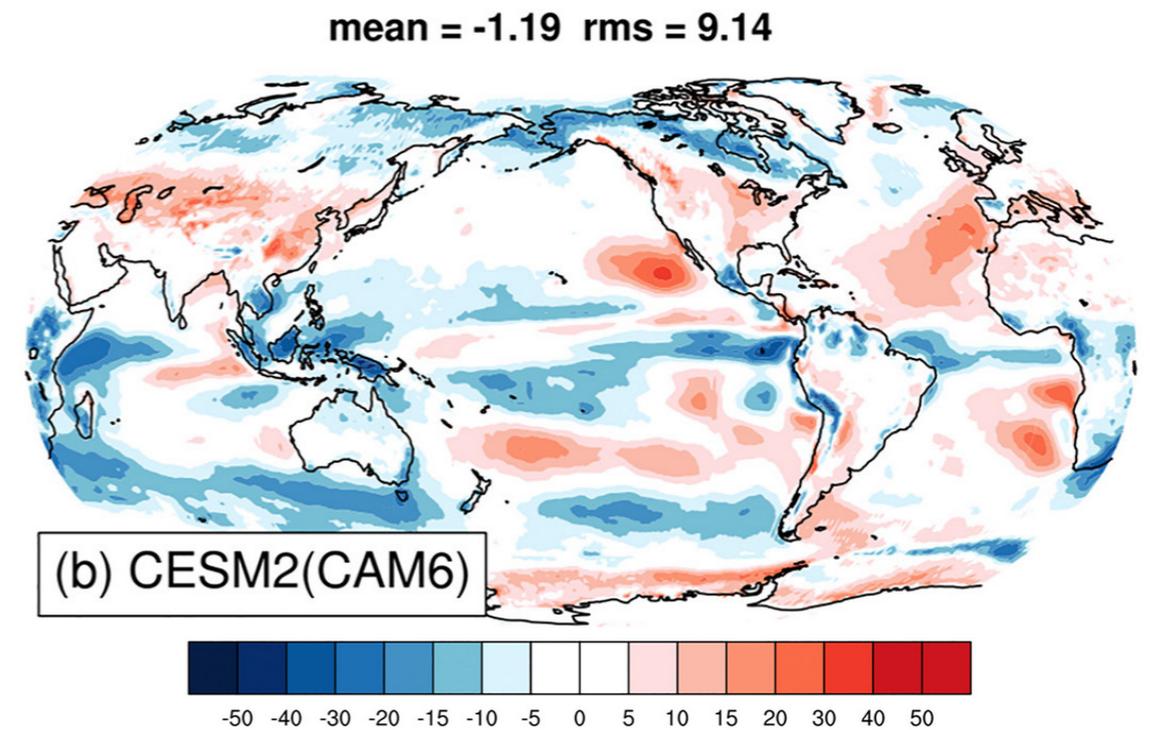
Old models: Global climate sensitivity plotted against CRE/CO₂ Radiative Effect for 19 Models: a strong linear fit between ΔCRE and global climate sensitivity. (Cess et al. 1990)

- ★ **Future change in CRE is the dominant source of difference between models and of uncertainty in climate model prediction.**

Stratocumulus cloud model bias leads to significant SST errors



G. Danabasoglu et al 2020,
<https://agupubs.onlinelibrary.wiley.com/doi/10.1029/2019MS001916>

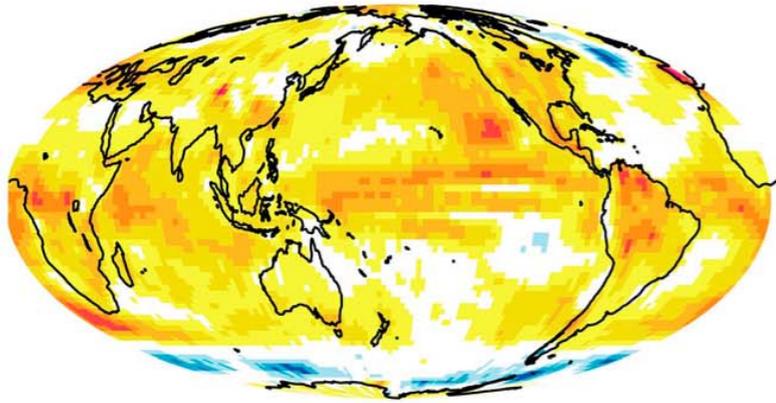


observed stratocumulus cloud fraction (%)

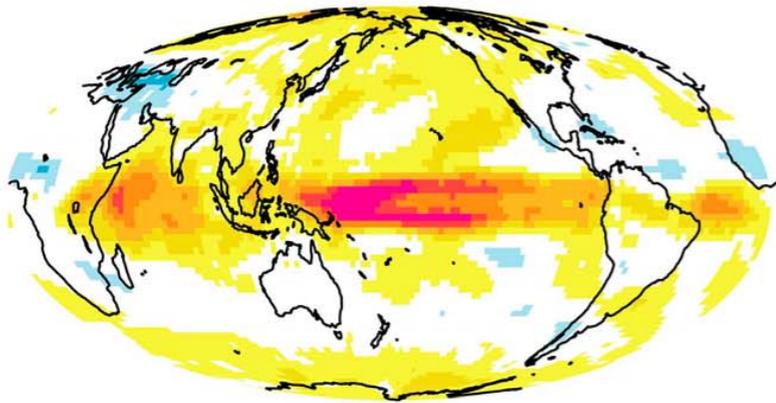
SST error (difference between model and observations) is large, ~2.5C in regions with underestimated stratocumulus cloud cover

Model Disagreement on Cloud Feedbacks

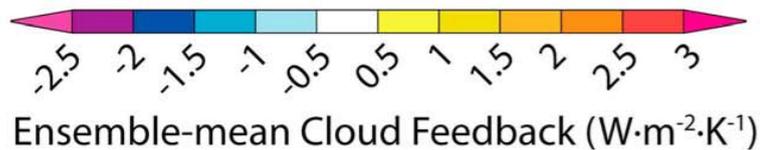
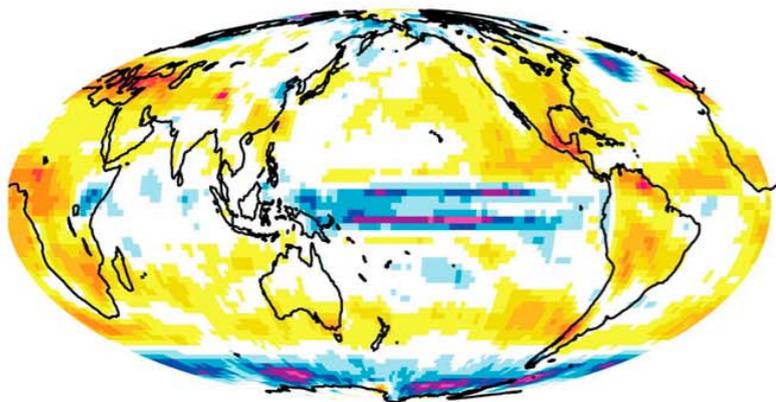
Net Cloud Feedback



LW Cloud Feedback



SW Cloud Feedback

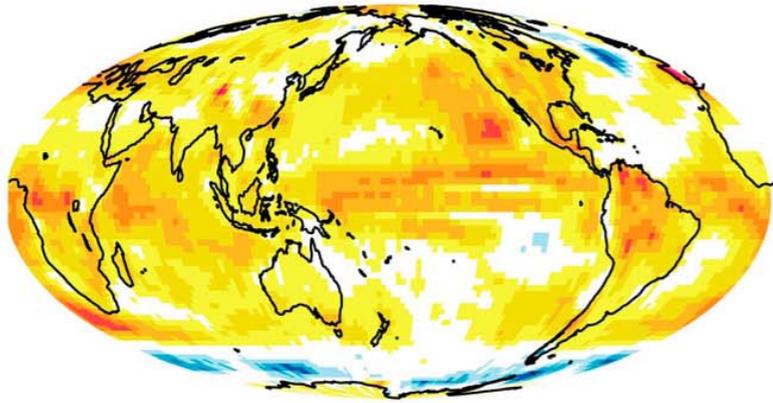


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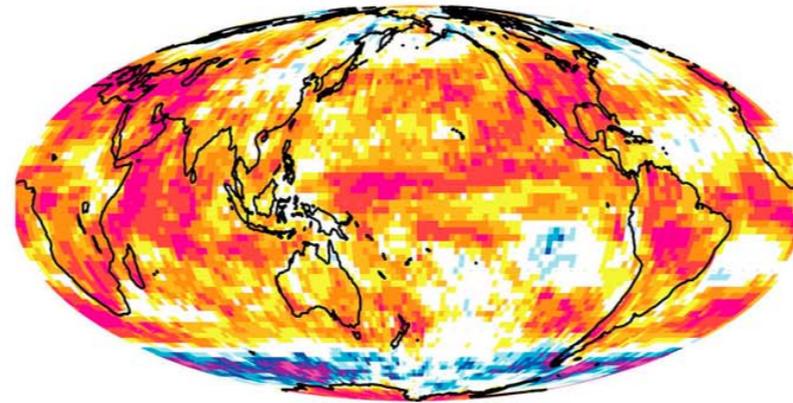
Spatial distribution of cloud feedback values and their uncertainties (Soden et al. 2011)

Model Disagreement on Cloud Feedbacks

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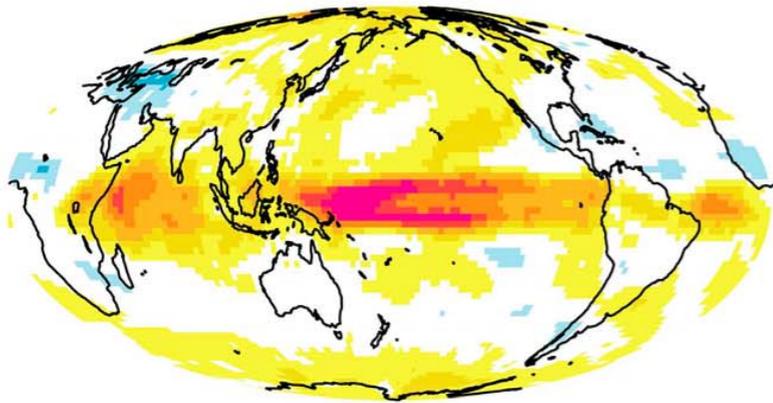
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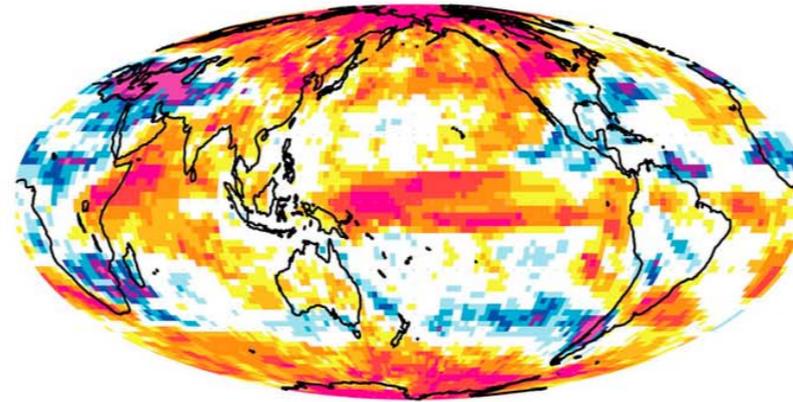
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- Right panel: number of models out of 12 with a positive cloud feedback.

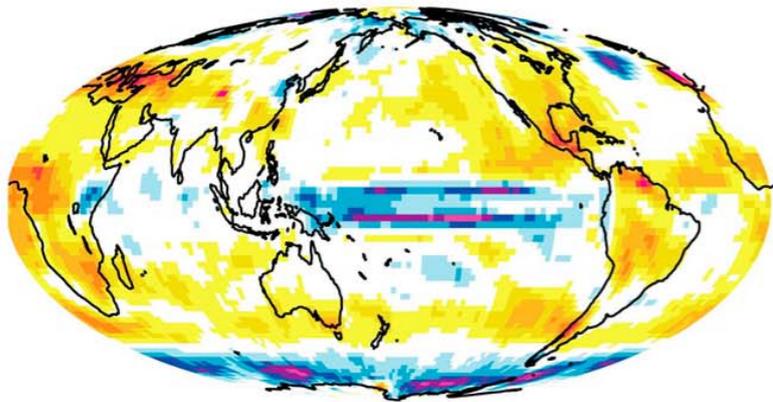
LW Cloud Feedback



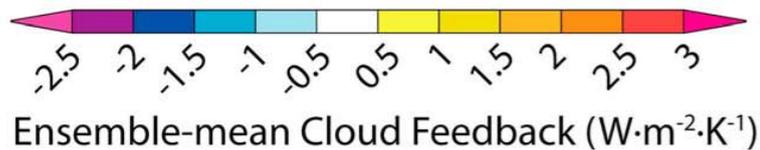
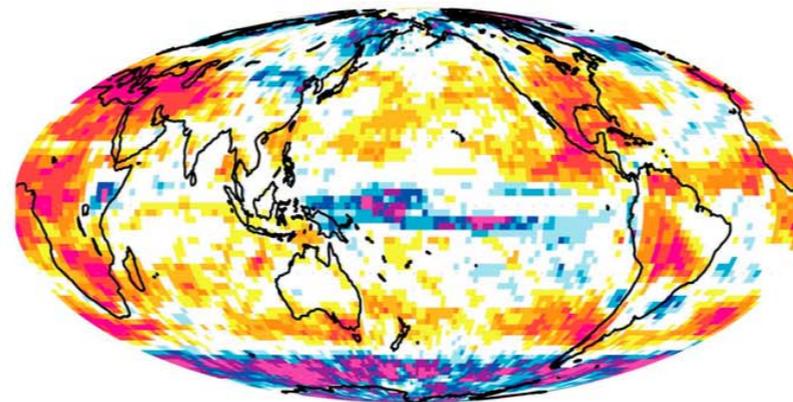
LW Cloud Feedback



SW Cloud Feedback

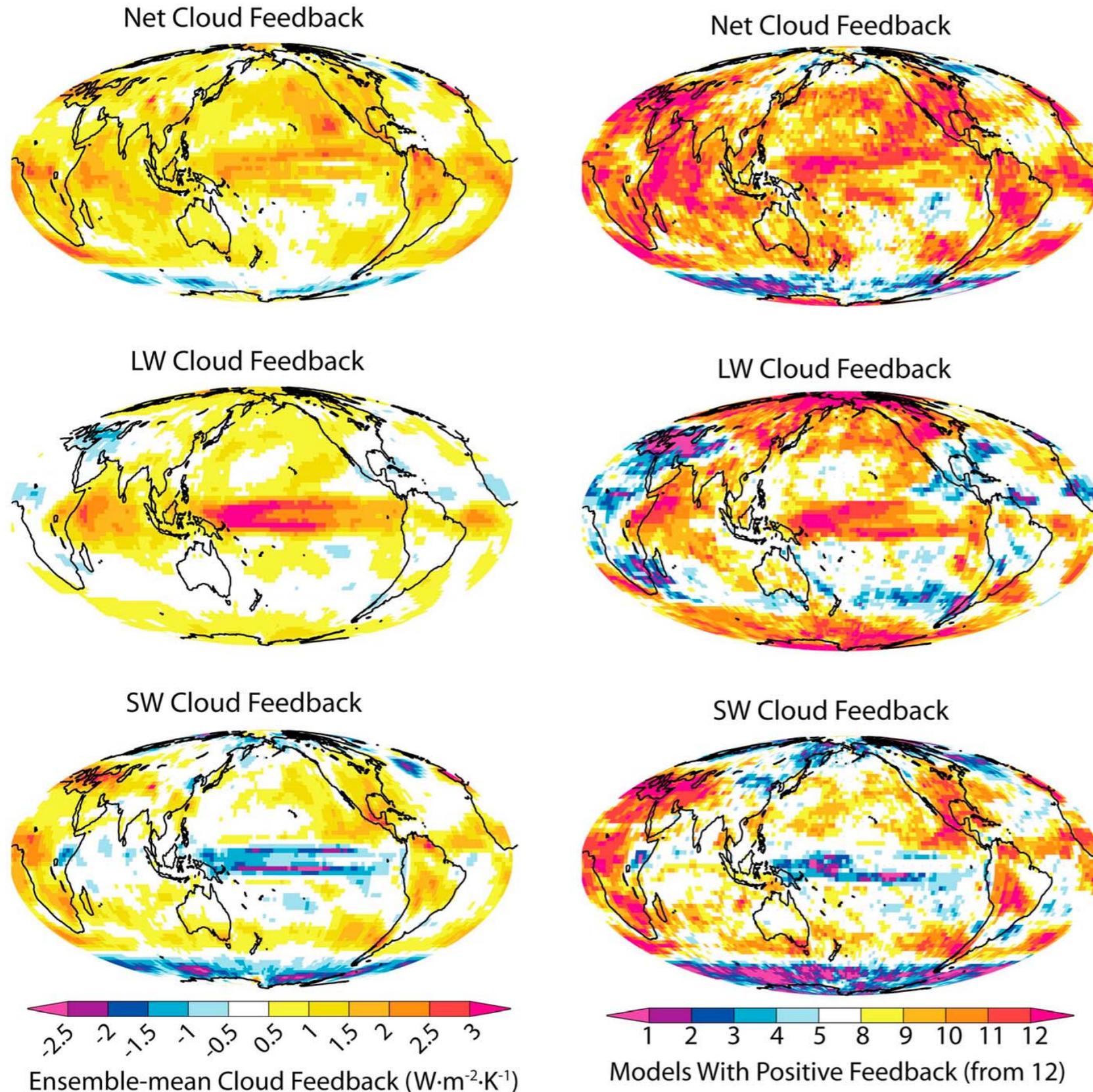


SW Cloud Feedback



Spatial distribution of cloud feedback values and their uncertainties (Soden et al. 2011)

Model Disagreement on Cloud Feedbacks

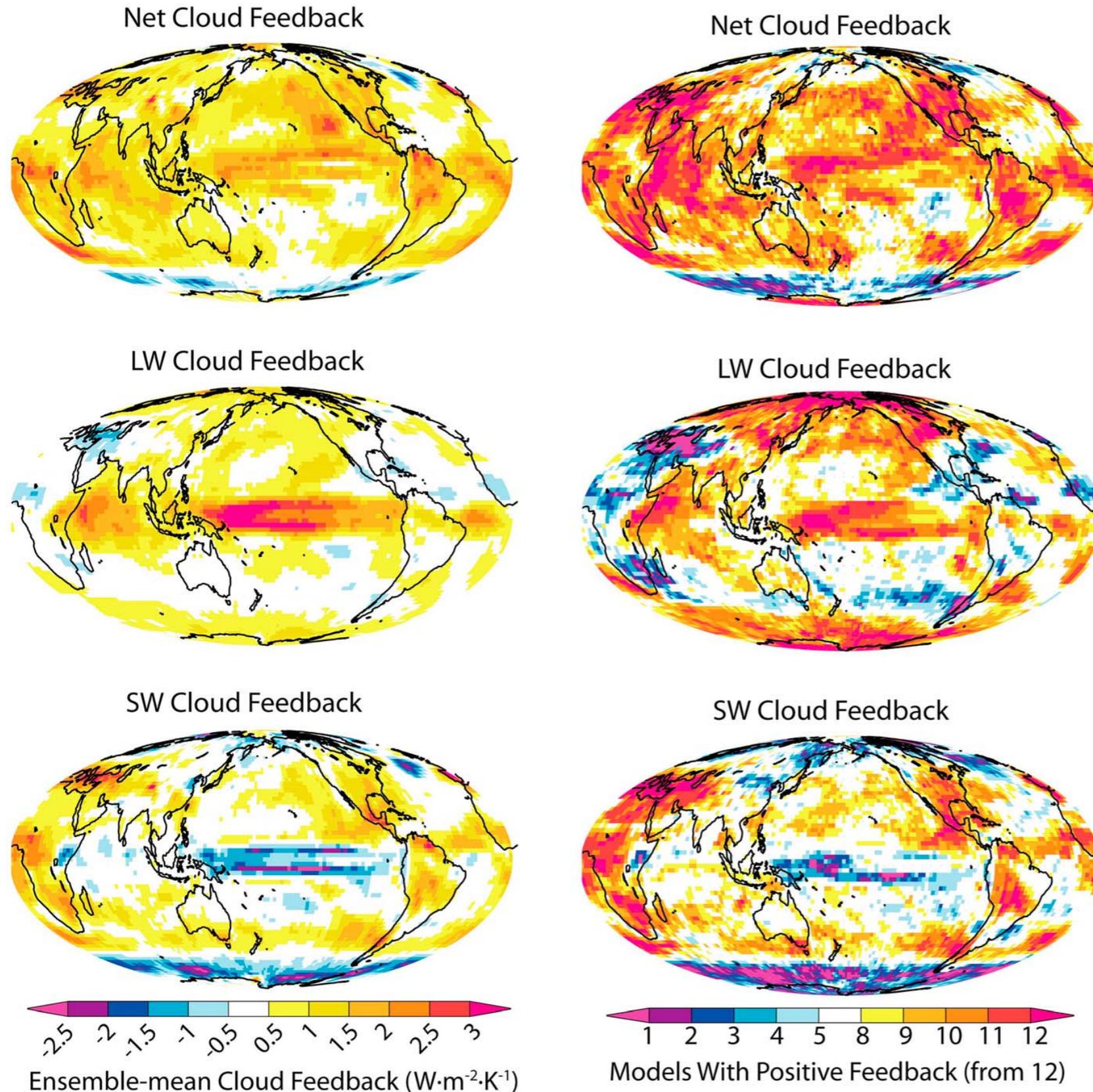


☉ **Clouds drive climate uncertainty, but which cloud types?**

☉ Right panel: number of models out of 12 with a positive cloud feedback.

☉ Right panels: $n < 6$ are areas of highly uncertain feedback.

Model Disagreement on Cloud Feedbacks



☉ **Clouds drive climate uncertainty, but which cloud types?**

☉ Right panel: number of models out of 12 with a positive cloud feedback.

☉ Right panels: $n < 6$ are areas of highly uncertain feedback.

☉ The subtropics, characterized by broad stratocumulus cloud coverage, tend to have the lowest model agreement.

notes section 7.4:
Clouds and climate uncertainty
(see next slides)

notes section 7.4:
Clouds and climate uncertainty

Two-layer energy balance again:

$$C_{\text{surface}} \frac{dT}{dt} = \frac{S_0}{4} (1 - \alpha(T)) + \epsilon(\text{CO}_2, T_a) \sigma T_a^4 - \sigma T^4$$

Surface/upper ocean energy balance

$$C_{\text{atm}} \frac{dT_a}{dt} = \epsilon(\text{CO}_2, T_a) \sigma T^4 - 2\epsilon(\text{CO}_2, T_a) \sigma T_a^4.$$

Atmospheric energy balance

notes section 7.4: Clouds and climate uncertainty

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Atmospheric energy balance

With cloud feedbacks given by:

$$\alpha(T) = \alpha_0 (1 + \Delta_{SW}(T - T_0))$$

Low clouds' albedo depends on surface temperature

$$\epsilon(T_a) = \epsilon_0(\text{CO}_2) (1 + \Delta_{LW}(T_a - T_{a,0}))$$

High clouds' emissivity depends on atmospheric temperature

$$\epsilon_0(\text{CO}_2) = 0.75 + 0.05 \log_2(\text{CO}_2/280).$$

Emissivity dependence on CO₂

notes section 7.4: Clouds and climate uncertainty

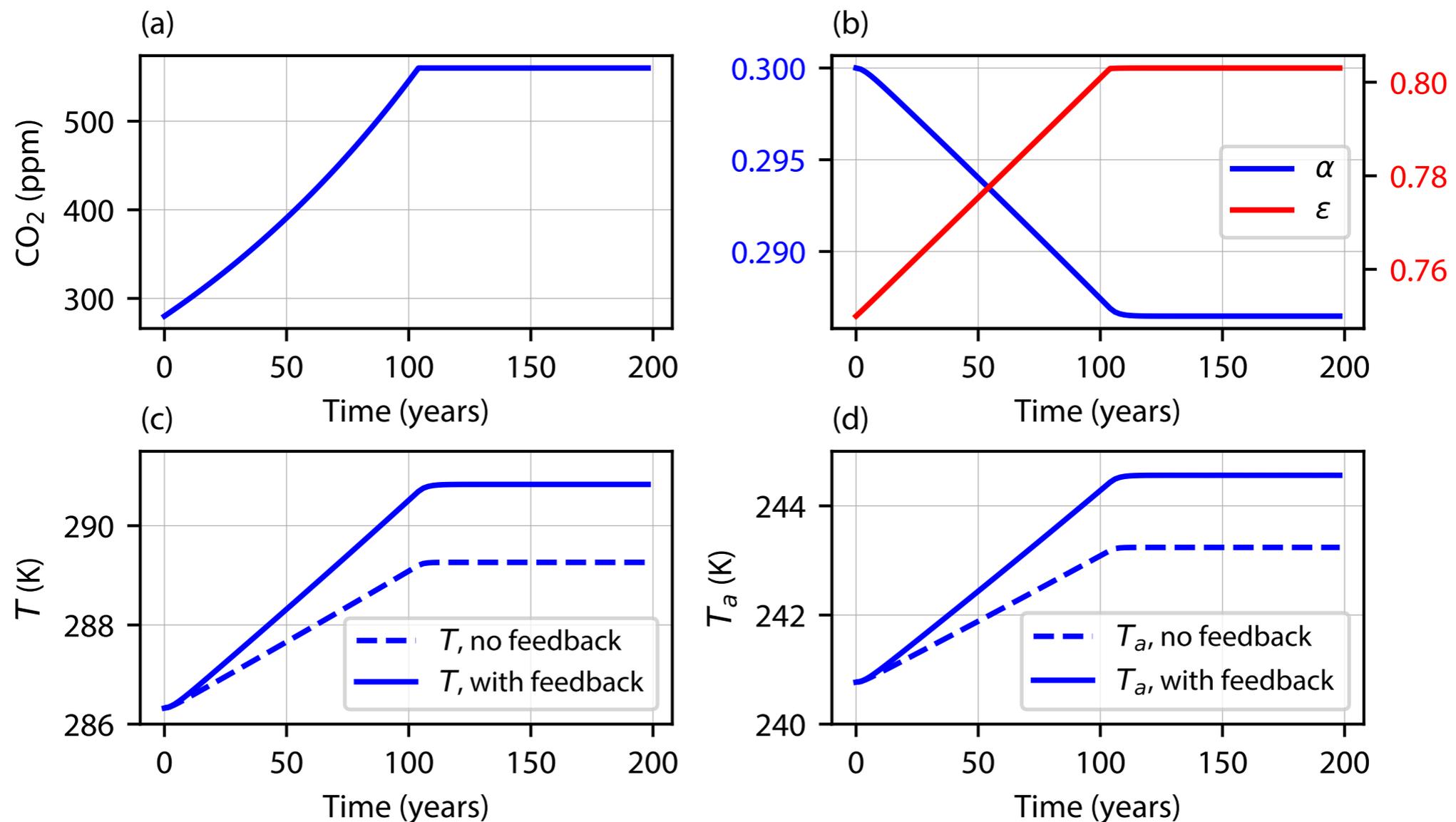


Figure 7.3: Cloud feedbacks in a two-layer energy balance model.

Response of a two-level energy balance model to SW and LW cloud feedbacks. (a) Atmospheric CO₂ as a function of time, representing a doubling scenario. (b) The change to cloud albedo and emissivity resulting from the formulation in eqn (7.3). (c) The surface temperature as a function of time with and without cloud feedbacks. (d) Same, for the atmospheric temperature.

workshop #3:

Cloud feedbacks and climate uncertainty in a simple two-level energy balance model

Summary

- Clouds are characterized by an outsized radiative impact; their shortwave and longwave radiative effect are large and partially offset each other.



stein egil liland: Nordland, Norway,

<https://www.pexels.com/photo/clouds-over-mountains-12035615/>

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- We covered MSE conservation $c_p T(z) + L \min(q^*, q_s) + gz = MSE_s$ & used it to understand convection (LCL, LFC, LNB), cloud formation & the atmospheric lapse rate.



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- ☉ **Clouds have been, & continue to be the most important yet poorly understood aspect of climate's response to CO₂.**

- ☉ & they are beautiful and interesting!


The End